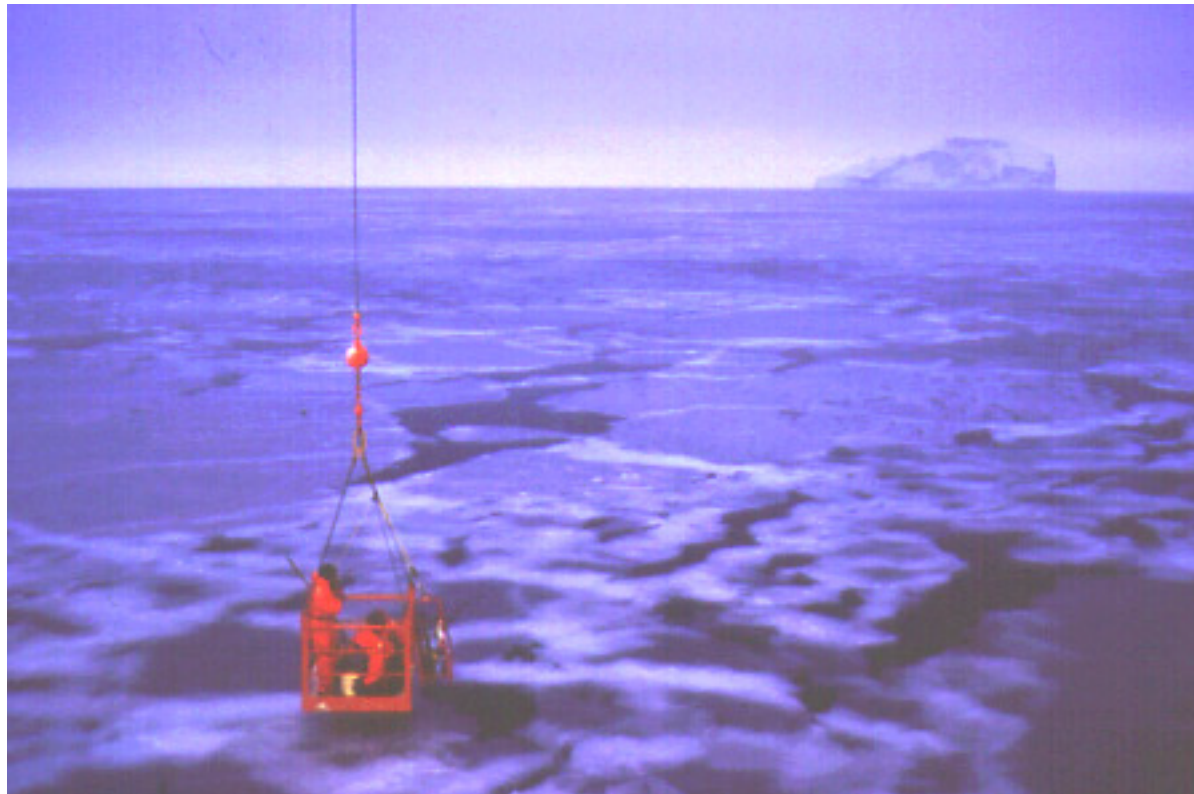


Sea ice processes in Antarctic polynyas

Kenneth M. Golden

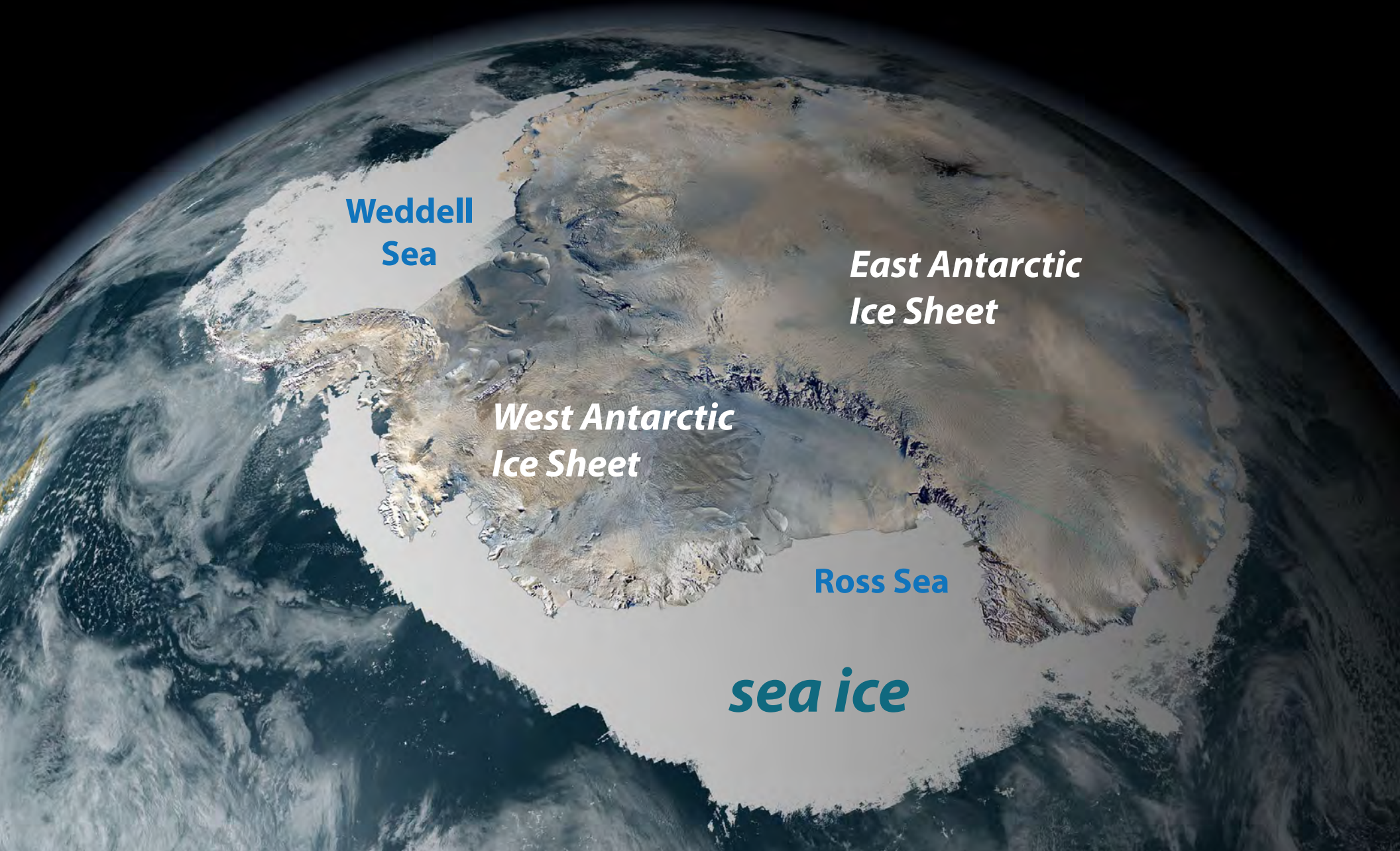
Department of Mathematics
University of Utah



Mertz Glacier Polynya, July 1999

ANTARCTICA

southern cryosphere



**Weddell
Sea**

***East Antarctic
Ice Sheet***

***West Antarctic
Ice Sheet***

Ross Sea

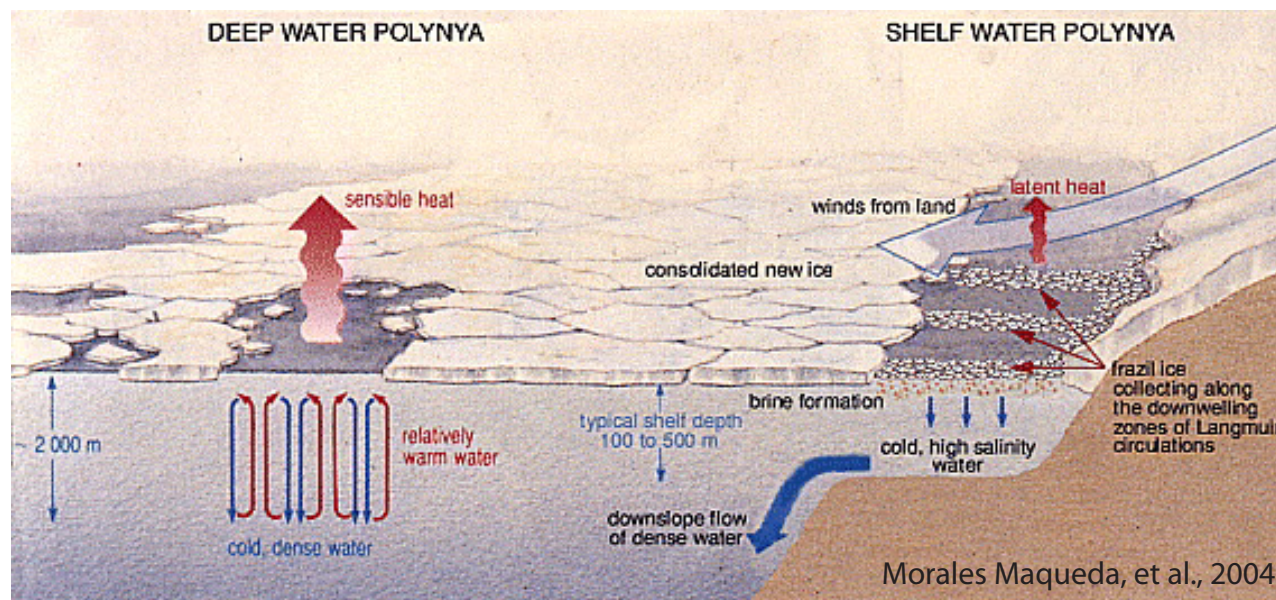
sea ice

Polynyas

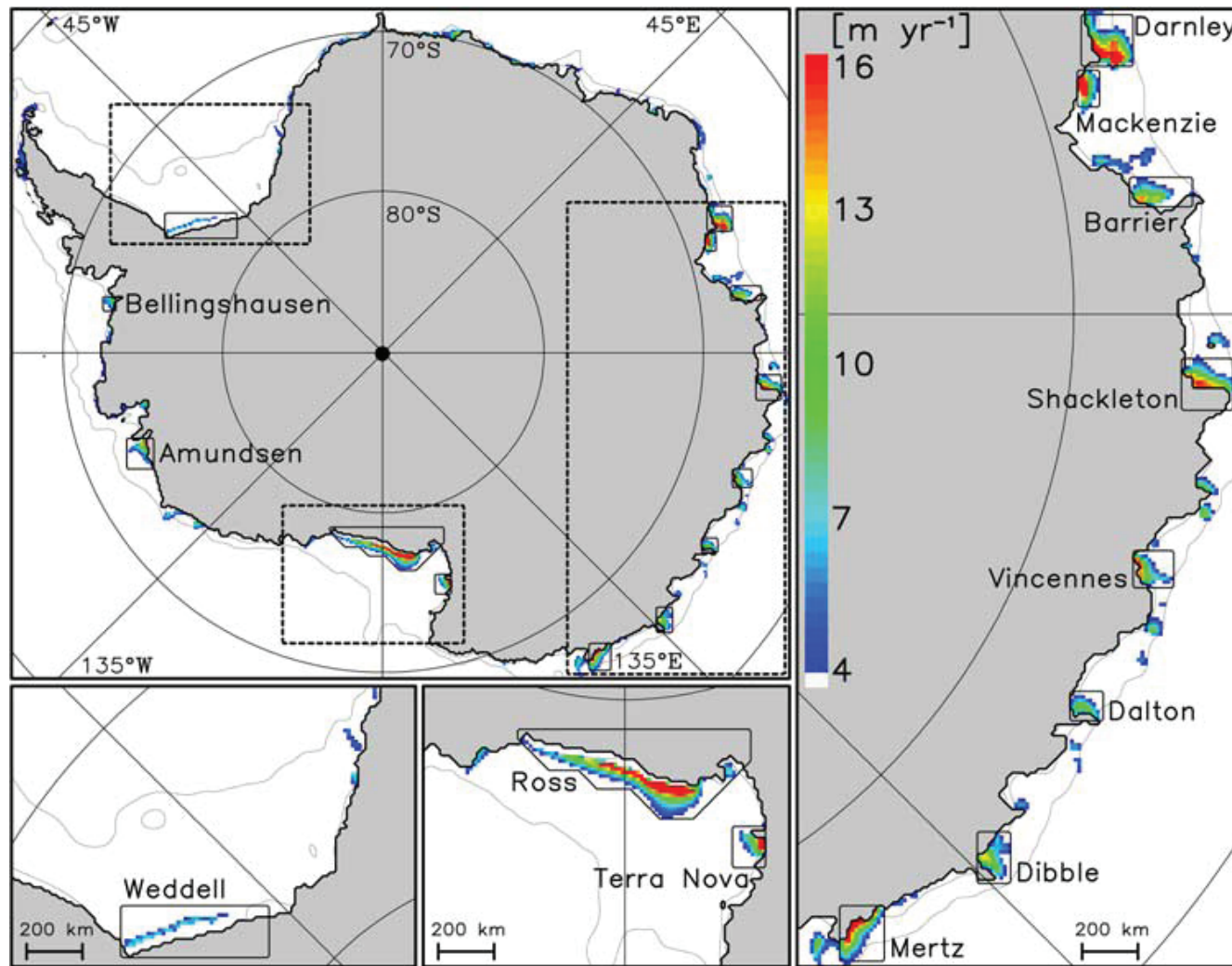
Size: 100 m - 1000 km

Two mechanisms can contribute to keeping polynyas open:

1. **Latent heat (or coastal) polynyas:** **Mertz Glacier Polynya**
Sea ice grows in open-water and is continually removed by winds and currents (e.g. katabatic winds)
 - latent heat released to the ocean during ice formation perpetuates the process
2. **Sensible heat (or open-ocean) polynyas:** **Weddell Polynya**
Upwelling warm waters, vertical heat diffusion, or convection may provide enough oceanic heat flux to maintain ice-free region



Antarctic coastal polynyas = ice factories



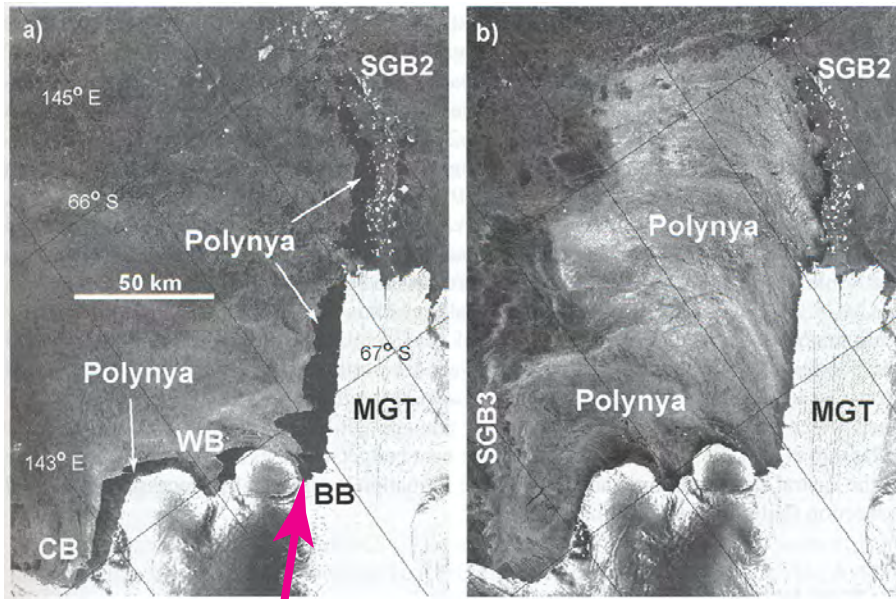
around 10% of Southern Ocean sea ice is produced in the major Antarctic coastal polynyas
ice production in Ross Ice Shelf Polynya decreased by about 30% from the 1990's to the 2000's
(caused by atmospheric warming or decreased polynya size from calving icebergs)

candidate for causing recent freshening of AABW

Tamura, Ohshima, Nihashi, GRL 2008

polynyas ice factories

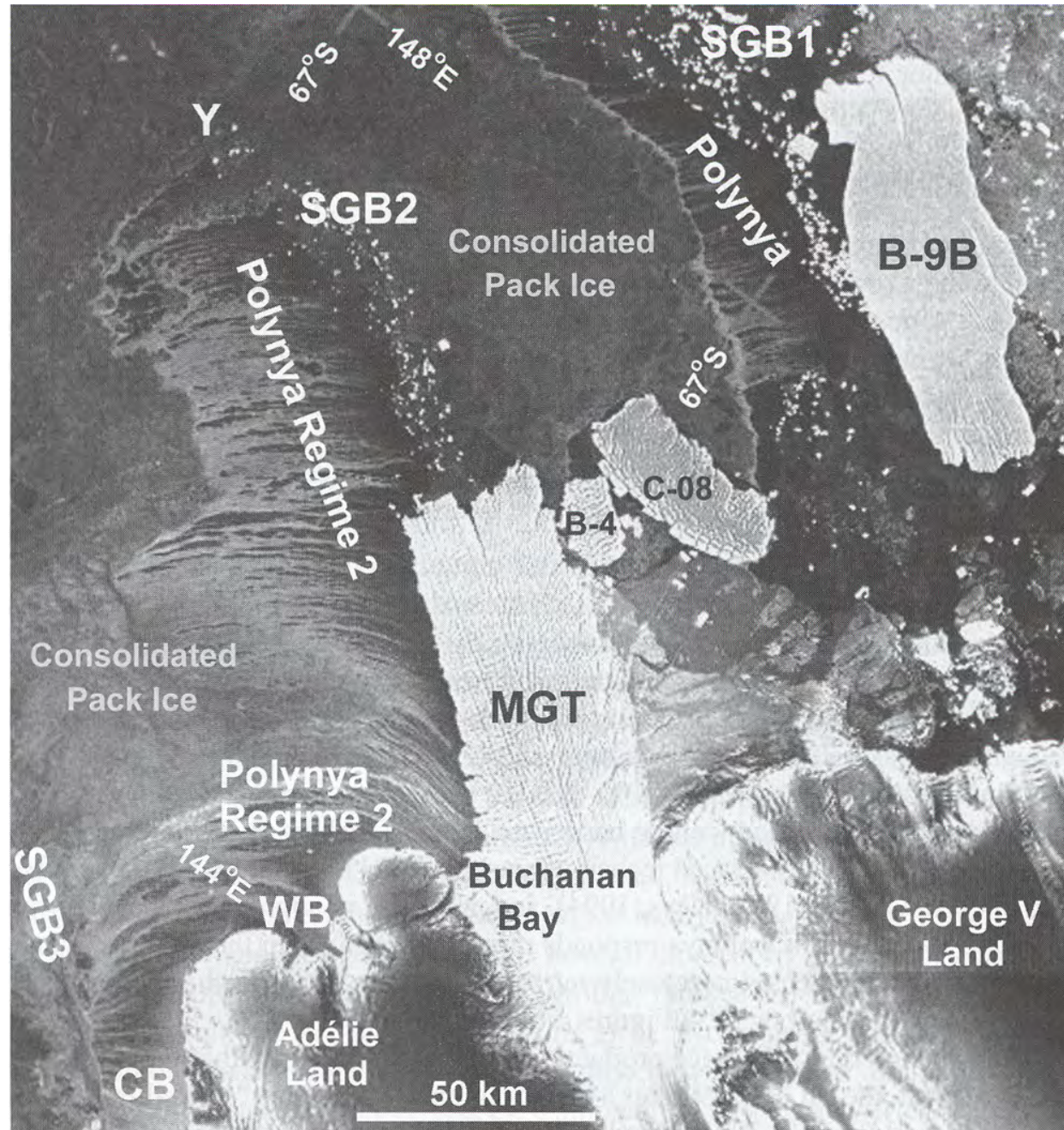
Mertz Glacier Polynya, located in East Antarctica, covers only 0.001% of the overall Antarctic sea ice zone at its maximum winter extent, but is responsible for 1% of the total sea ice production in the Southern Ocean.

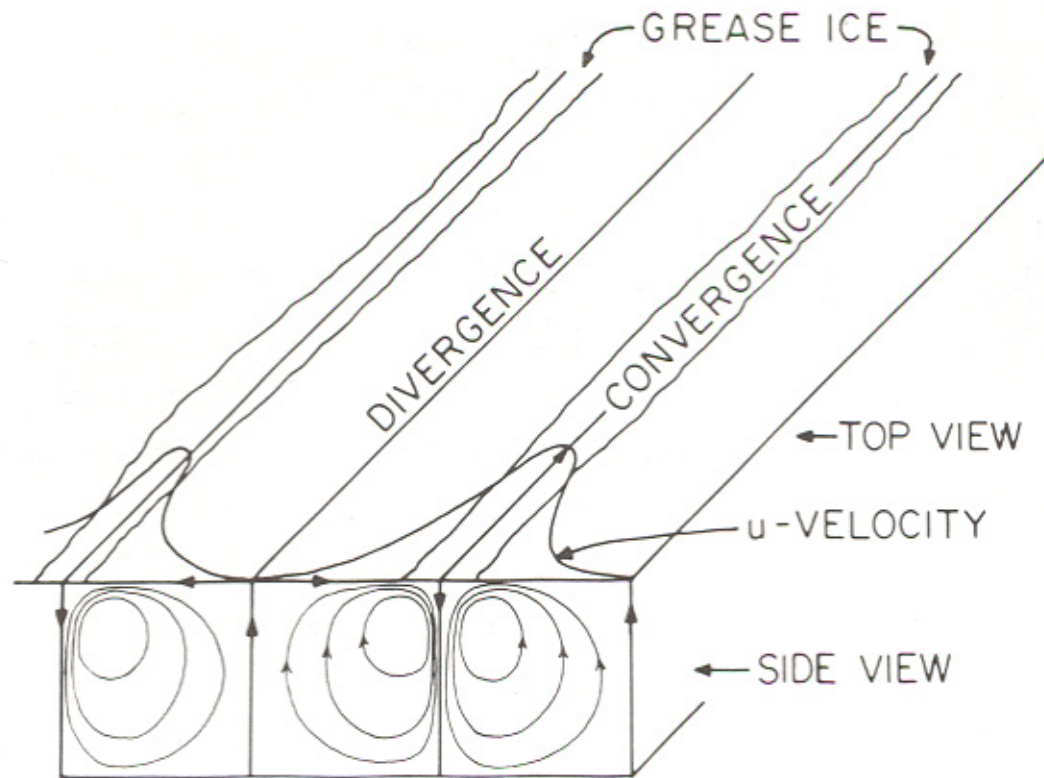


Buchanan Bay

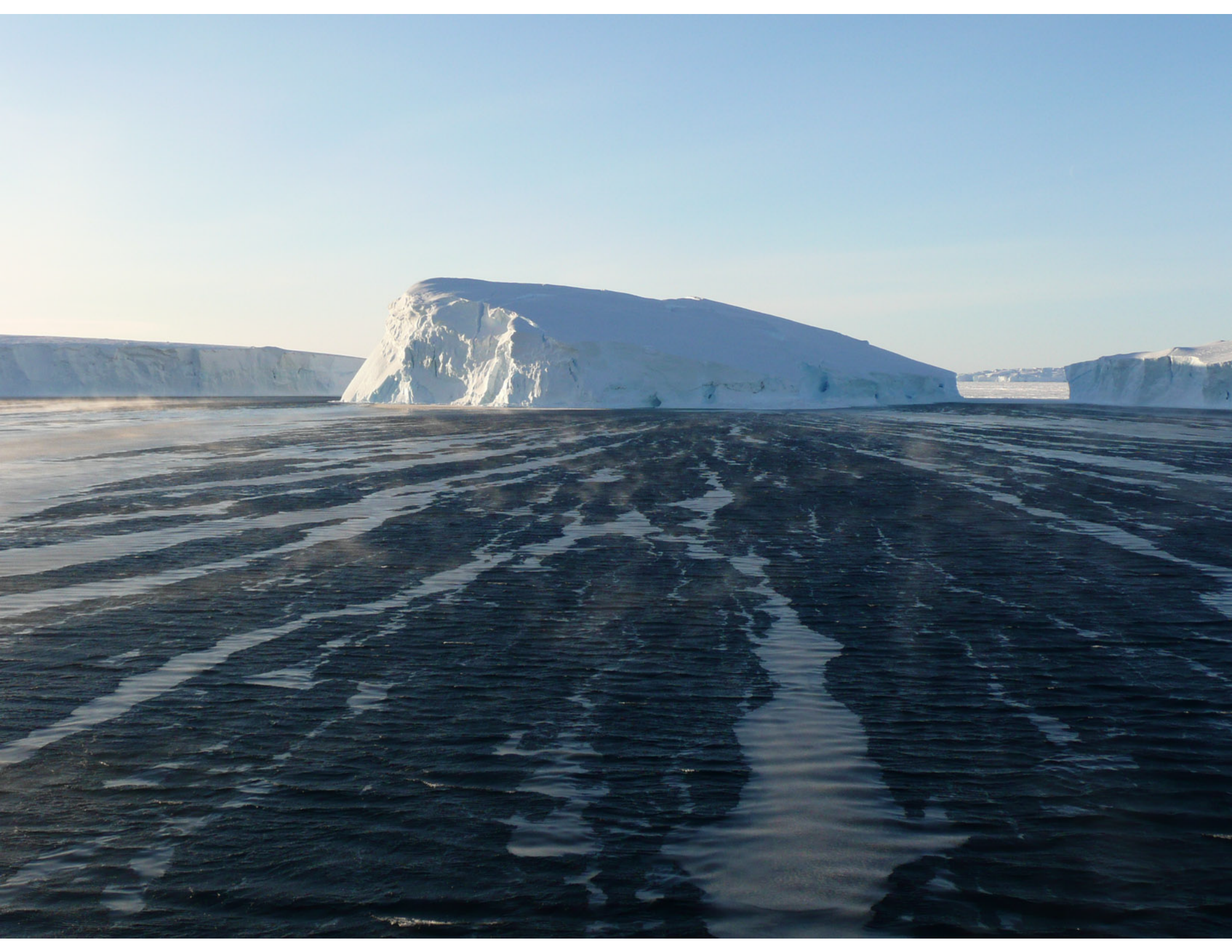


Mertz Glacier Polynya -- third largest Antarctic sea ice producer

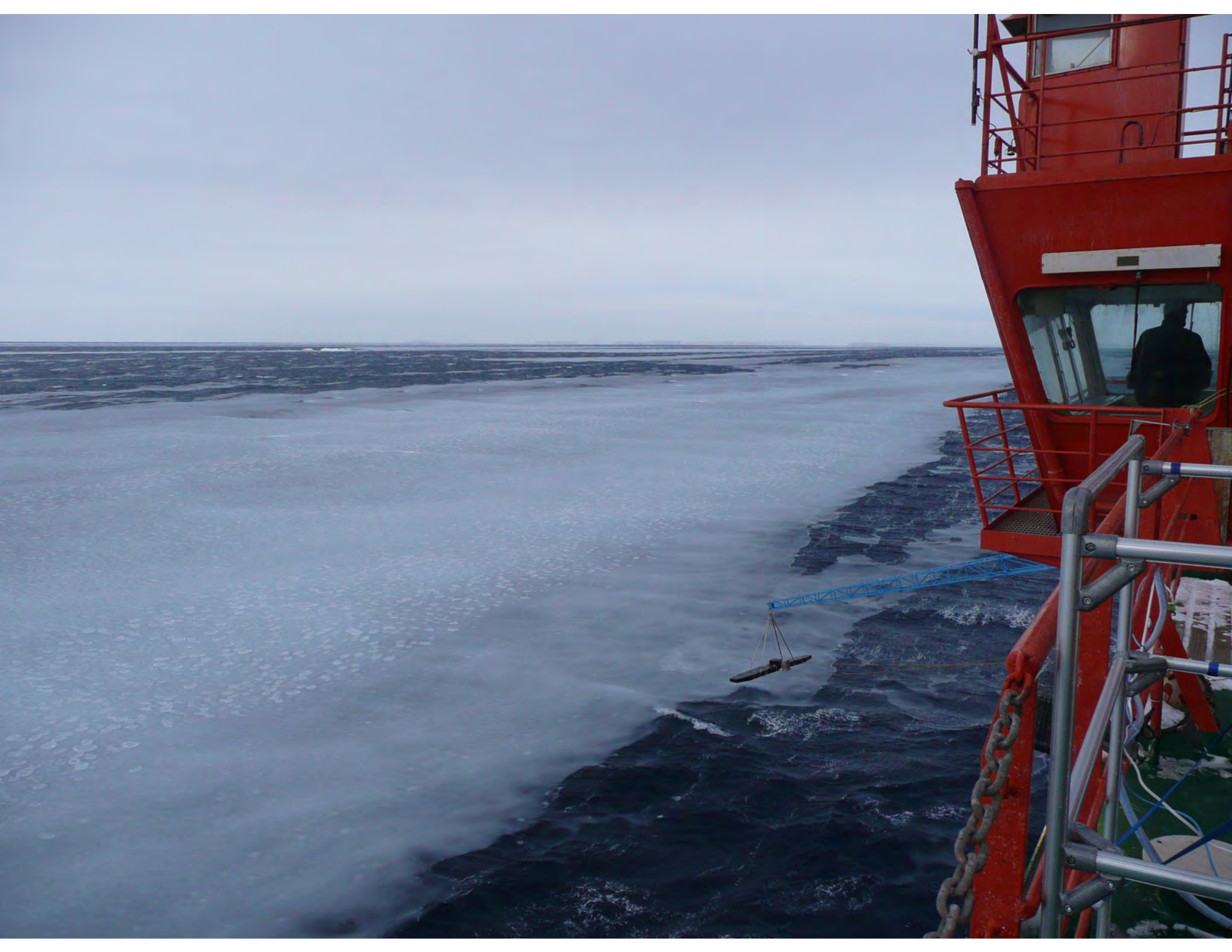




effect of Langmuir circulation
on grease and pancake ice

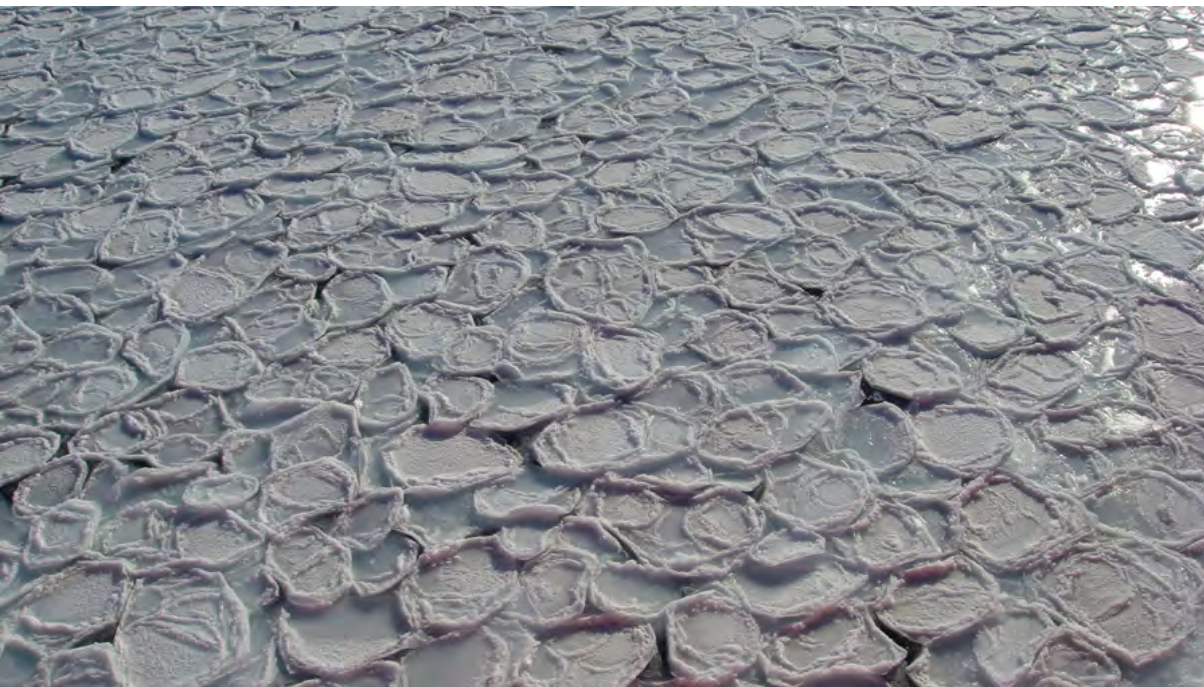


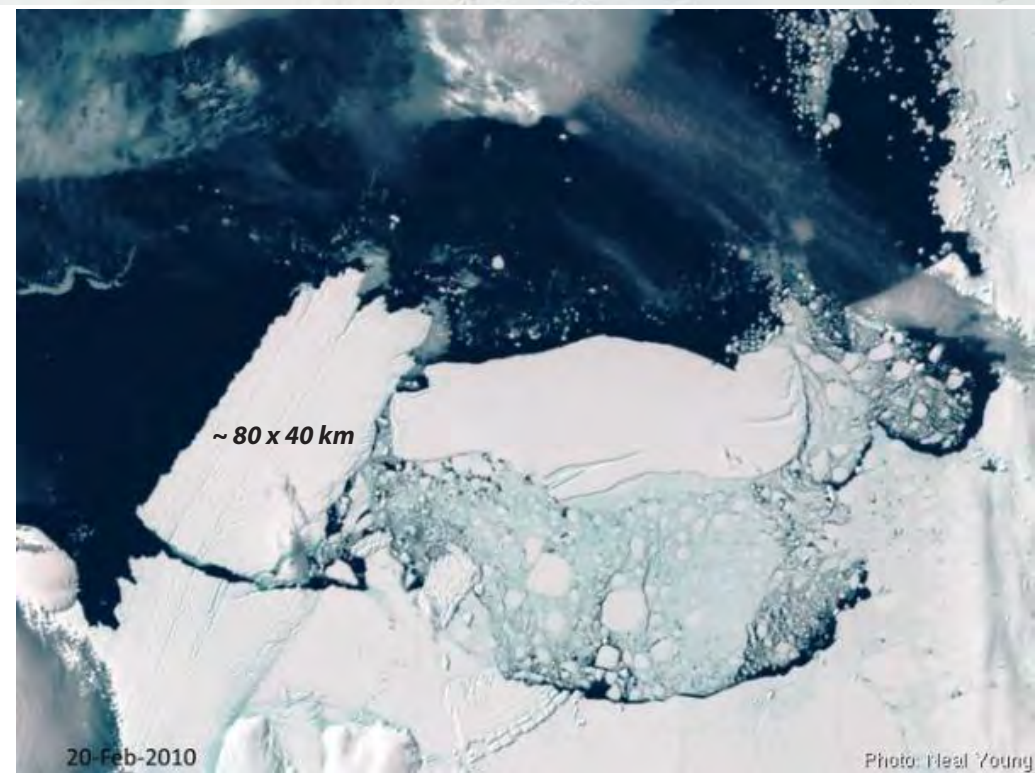
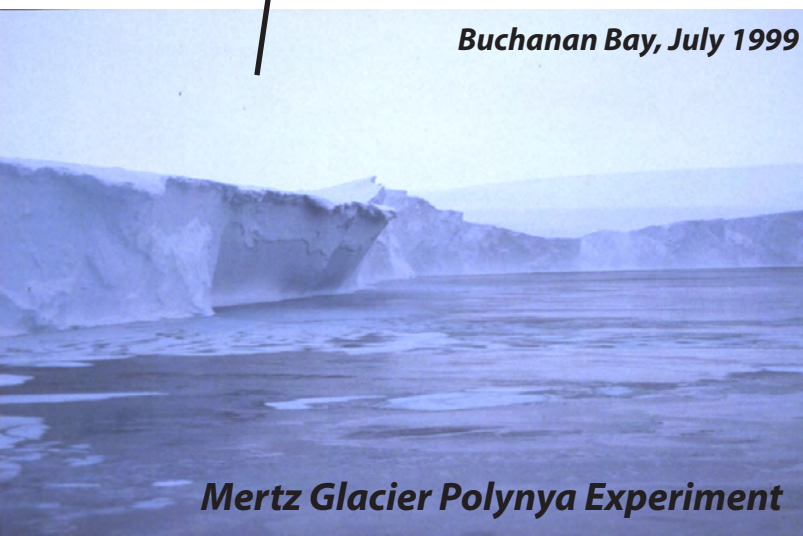
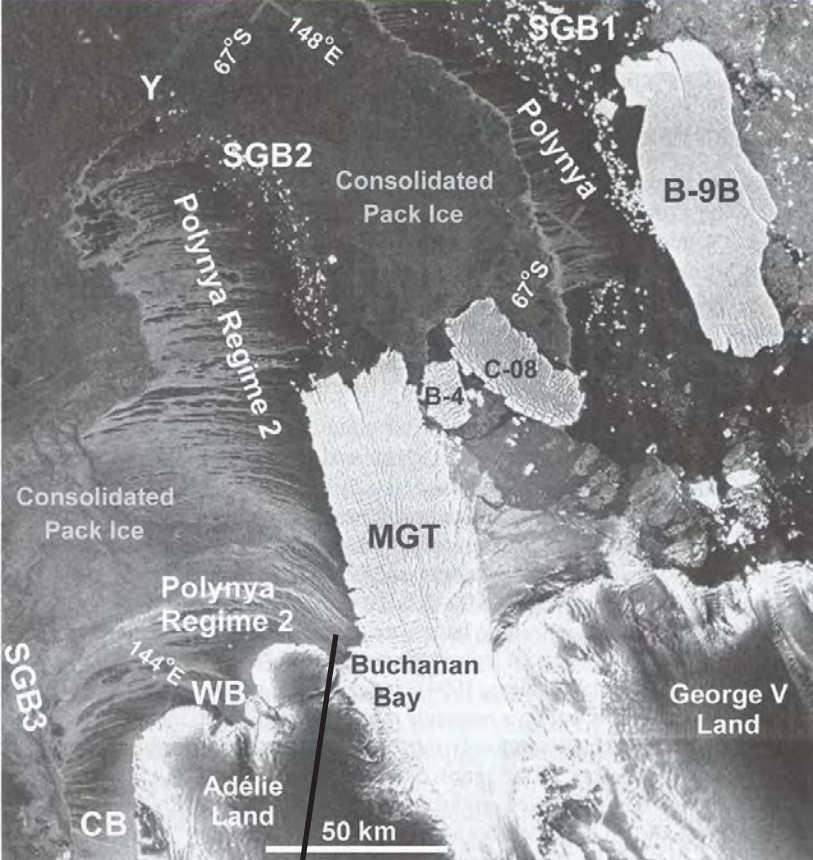




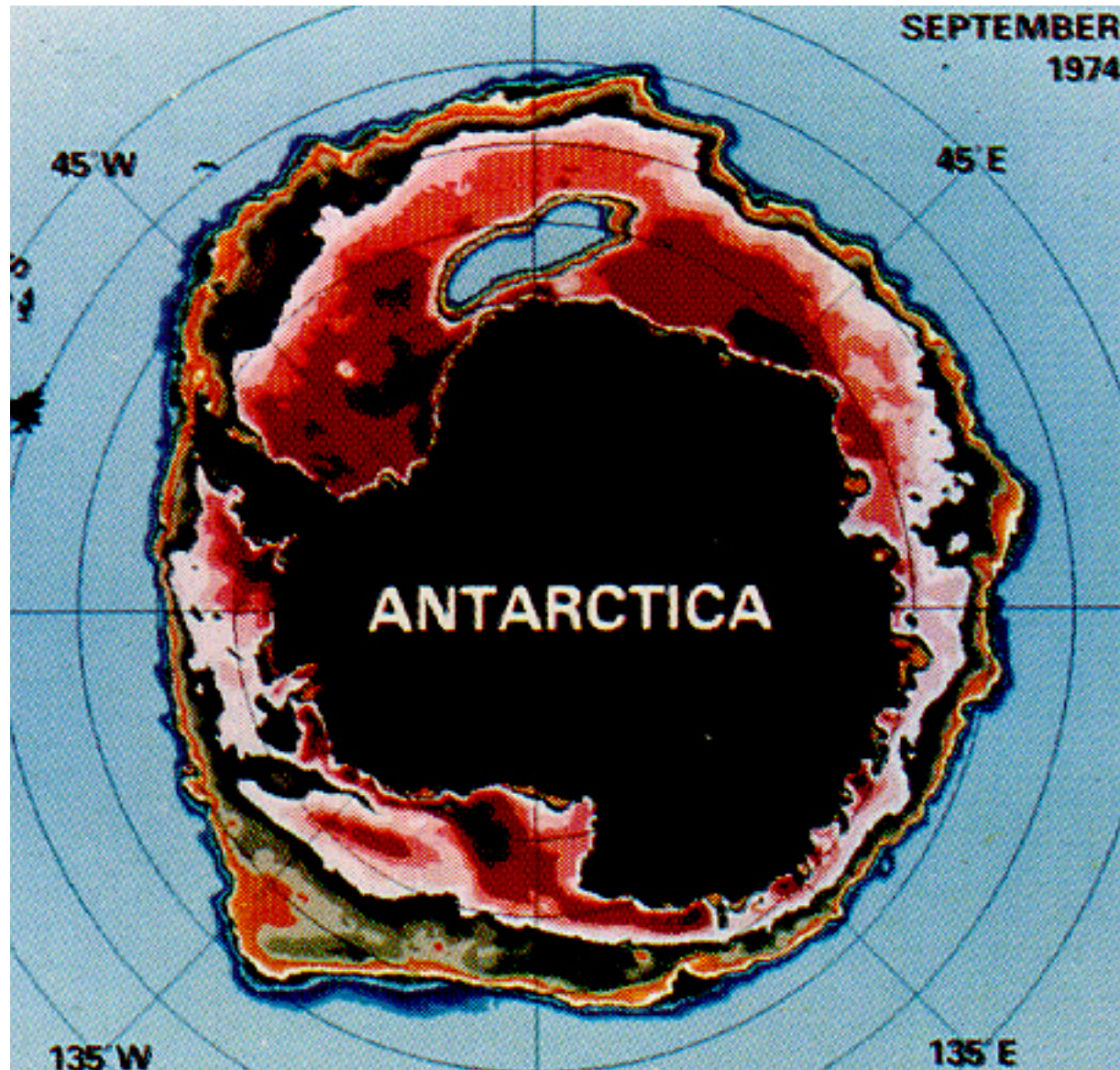


pancake ice



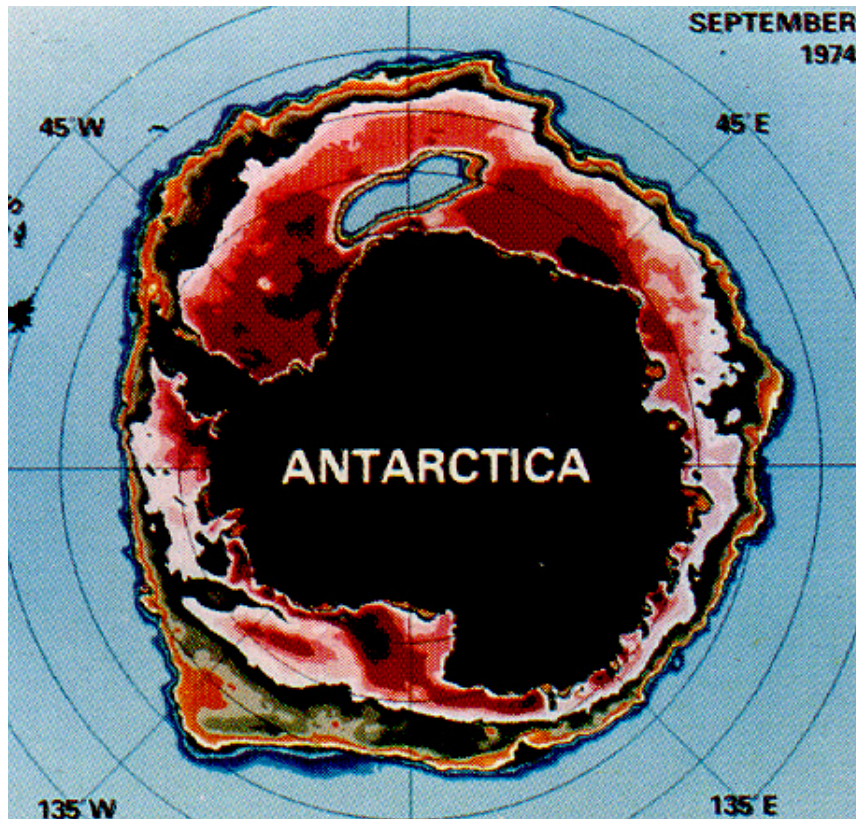


Weddell Polynya



Antarctic Zone Flux Experiment (ANZFLUX) 1994

Antarctic Zone Flux Experiment (ANZFLUX) 1994



*dynamic equilibrium
of sea ice thickness*

snow loading during storms

surface flooding ->
snow-ice formation

**controlled by
ice permeability**

snow-ice growth a key process in Antarctic

may become more important in Arctic with
thinning ice and increased precipitation

Ackley, Lytle, Golden, Darling, Kuehn, 1995
Maksym and Jeffries, 2001

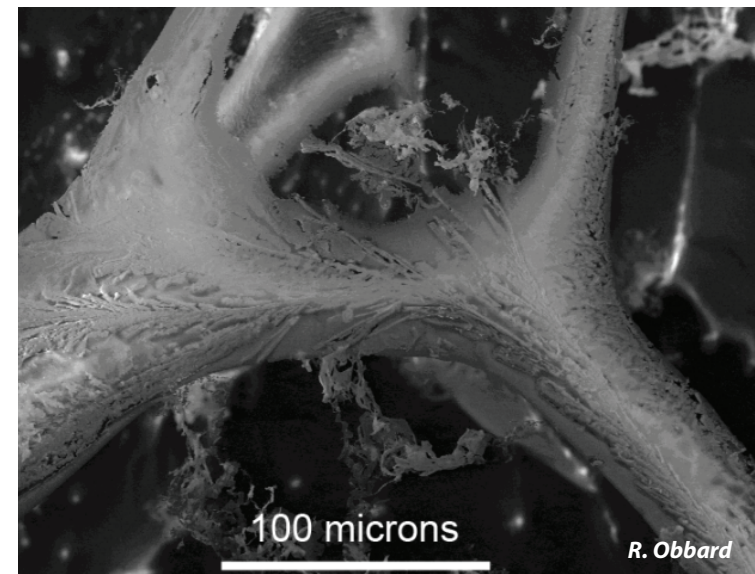
Maksym and Markus, 2008
Maksym and Golden, 2012



*sea ice may appear to be a
barren, impermeable cap ...*



brine inclusions in sea ice (mm)



micro - brine channel (SEM)

***sea ice is a
porous composite***

pure ice with brine, air, and salt inclusions

brine channels (cm)



horizontal section

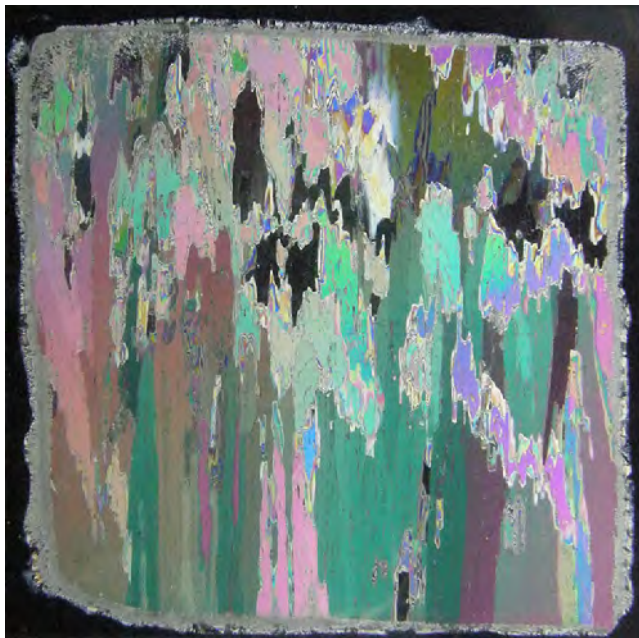


vertical section

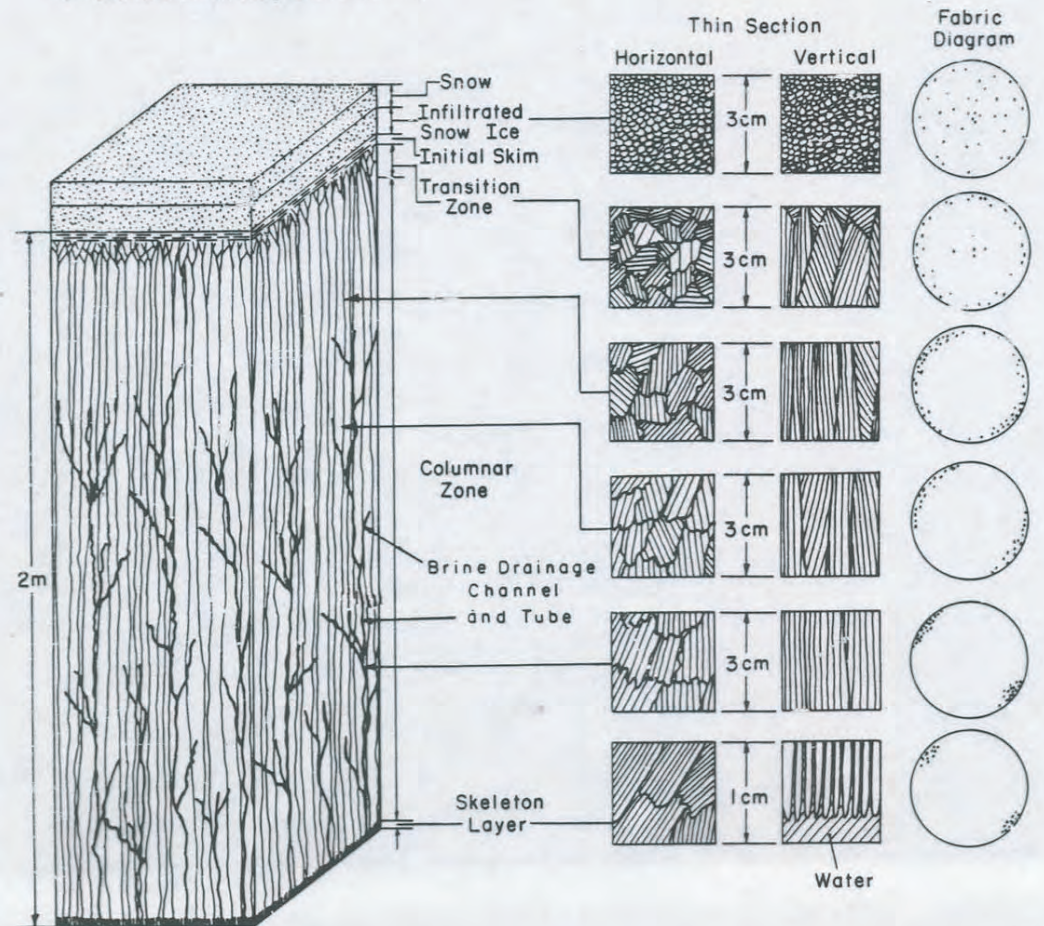
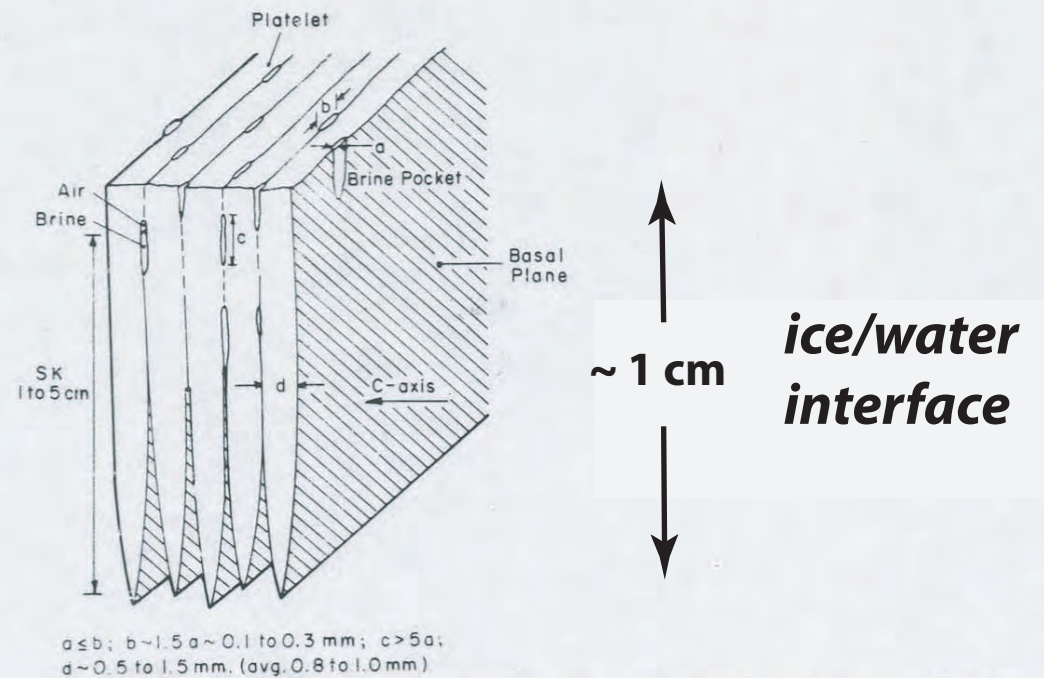
cross-sections of sea ice structure

$$T_{freeze} = -1.8^{\circ}\text{C}$$

crystallographic texture



vertical thin section



fluid flow through porous sea ice mediates key processes in polar climate and ecosystems:

evolution of Arctic melt ponds and sea ice albedo

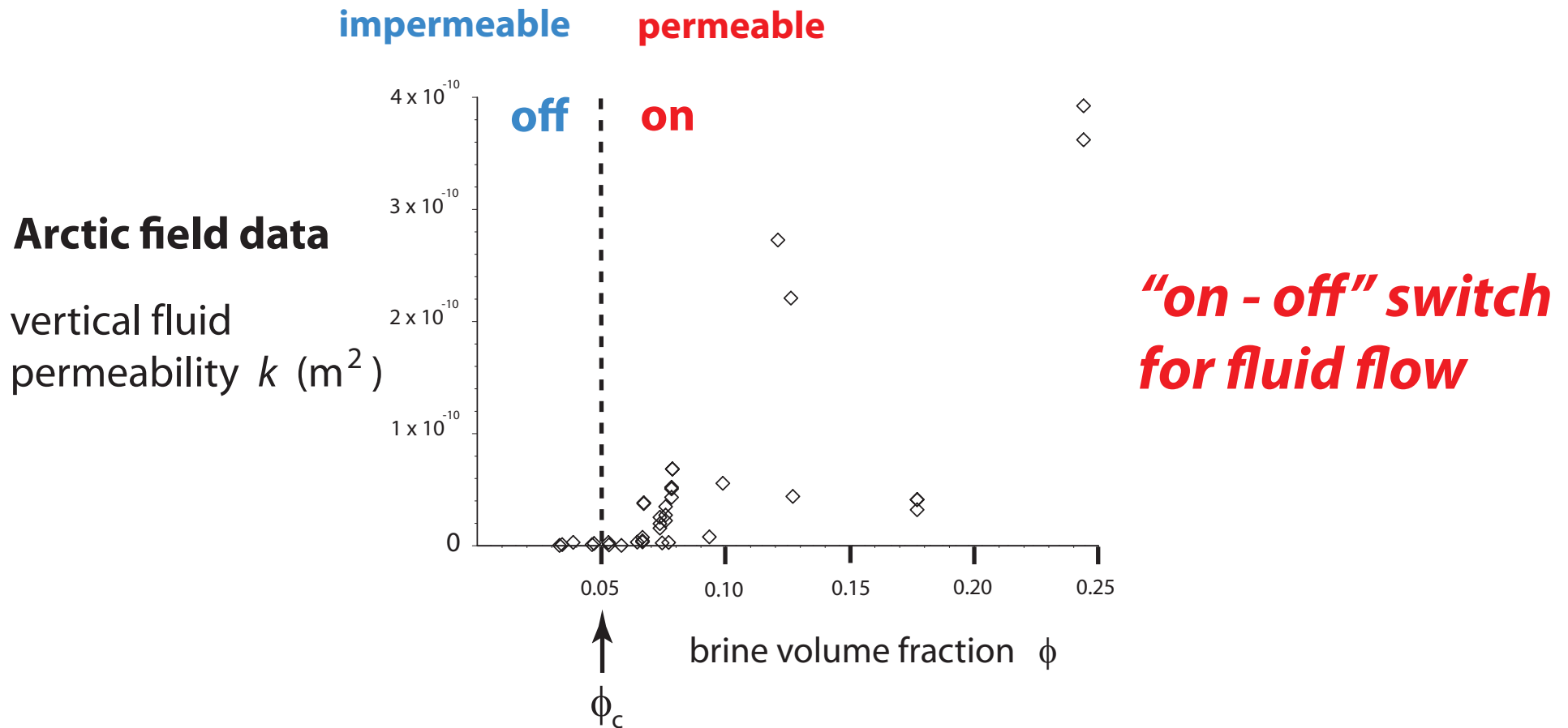


nutrient flux for algal communities



- *formation and melting of sea ice*
- *drainage of brine and melt water*
- *ocean-ice-atmosphere exchanges of heat, brine, CO₂*
- *growth and decline of microbial communities*

Critical behavior of fluid transport in sea ice



critical brine volume fraction $\phi_c \approx 5\%$ \longleftrightarrow $T_c \approx -5^\circ \text{C}$, $S \approx 5 \text{ ppt}$

RULE OF FIVES

Golden, Ackley, Lytle Science 1998

Golden, Eicken, Heaton, Miner, Pringle, Zhu Geophys. Res. Lett. 2007

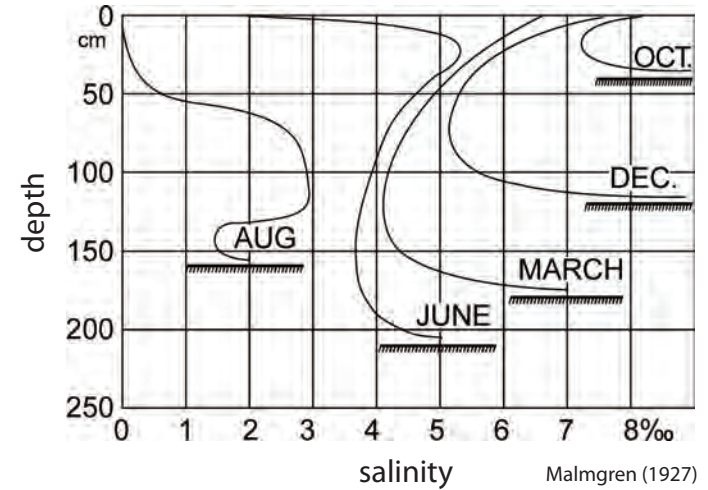
Pringle, Miner, Eicken, Golden J. Geophys. Res. 2009

rule of fives constrains:

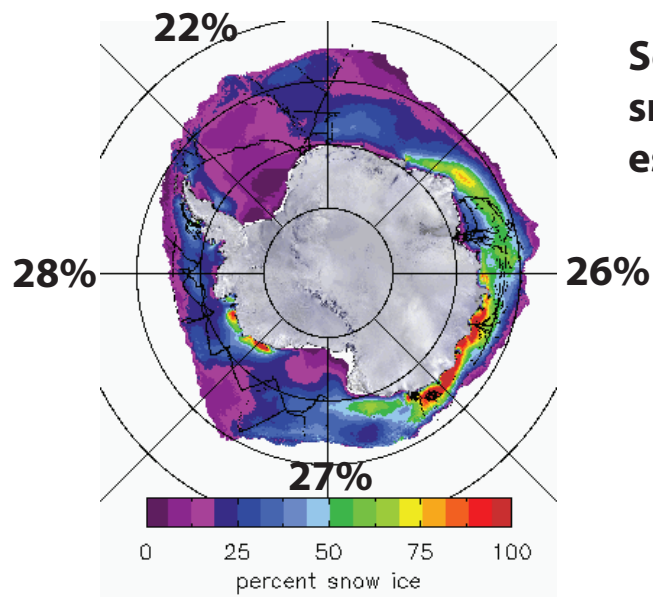
Antarctic surface flooding and snow-ice formation



evolution of salinity profiles



currently assumed constant in climate models



September
snow-ice
estimates

convection - enhanced thermal conductivity

Lytle and Ackley, 1996

Trodahl, et. al., 2000, 2001

Wang, Zhu, Golden, 2012



Antarctic snow-to-ice conversion from passive microwave imagery

T. Maksym and T. Markus, 2008



sea ice algal communities

D. Thomas 2004

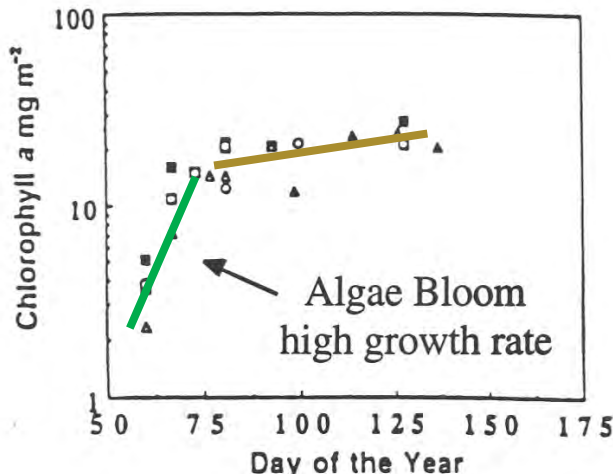
nutrient replenishment
controlled by ice permeability

biological activity turns on
or off according to
rule of fives

Golden, Ackley, Lytle Science 1998

Fritsen, Lytle, Ackley, Sullivan Science 1994

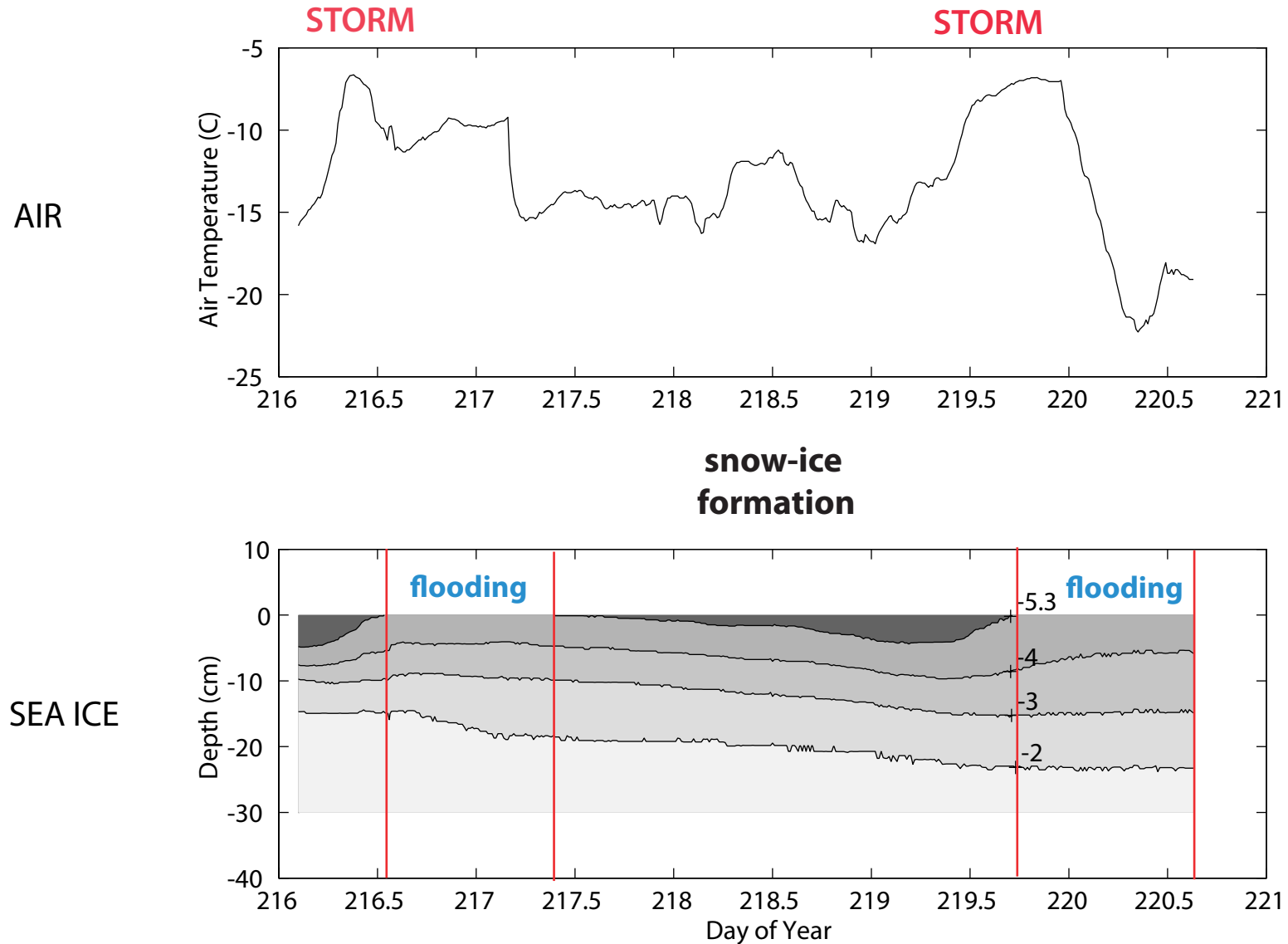
critical behavior of microbial activity



Convection-fueled algae bloom
Ice Station Weddell

ANZFLUX drift camp

snow loading, surface flooding and subsequent snow - ice formation

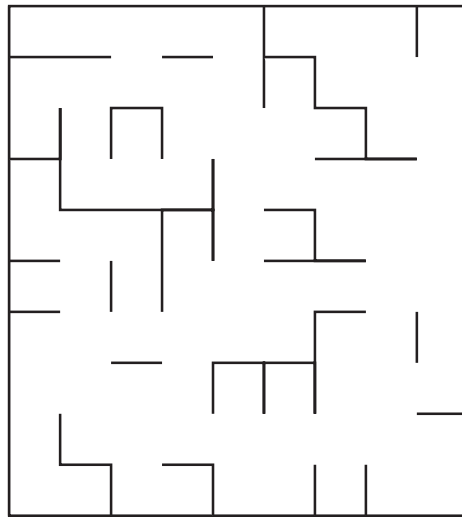


**theoretical models explaining the
rule of fives and fluid flow properties**

percolation theory

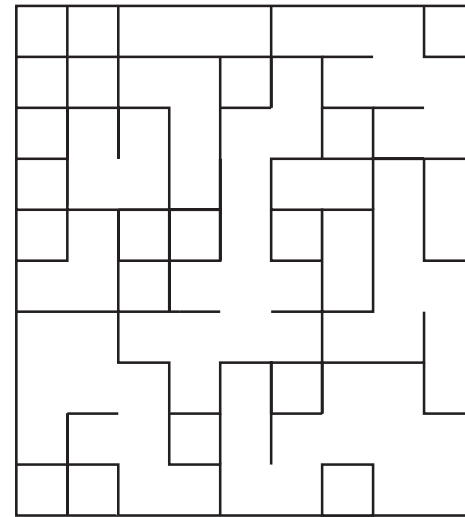
mathematical theory of connectedness

impermeable



$$p = 1/3$$

permeable



$$p = 2/3$$

a bond is *open* with probability p
closed with probability $1-p$

percolation threshold

$$p_c = 1/2 \quad \text{for } d = 2$$

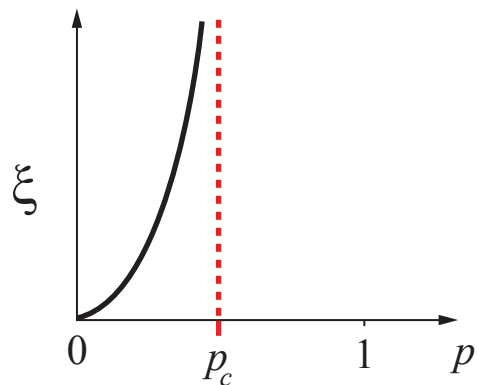
***first appearance
of infinite cluster***

order parameters in percolation theory

geometry

correlation length

characteristic scale
of connectedness

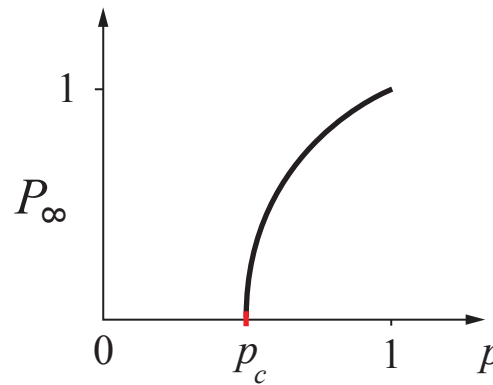


$$\xi(p) \sim |p - p_c|^{-\nu}$$

$$p \rightarrow p_c$$

infinite cluster density

probability the origin
belongs to infinite cluster

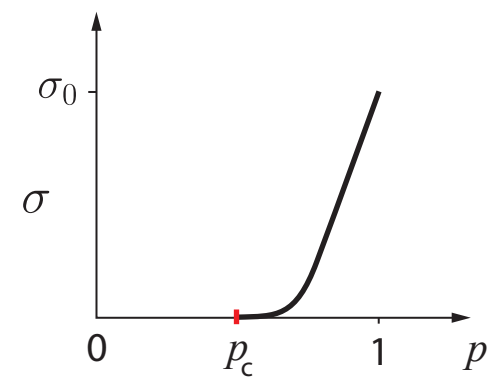


$$P_\infty(p) \sim (p - p_c)^\beta$$

$$p \rightarrow p_c^+$$

transport

effective conductivity
or fluid permeability



$$\sigma(p) \sim \sigma_0 (p - p_c)^t$$

$$p \rightarrow p_c^+$$

UNIVERSAL critical exponents for lattices -- depend only on dimension

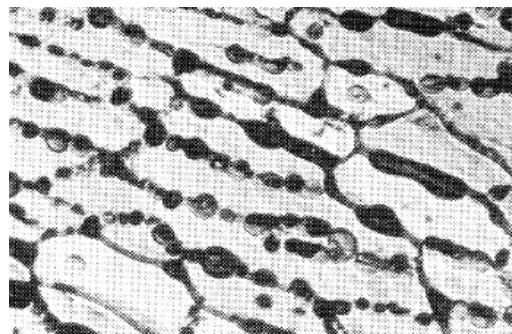
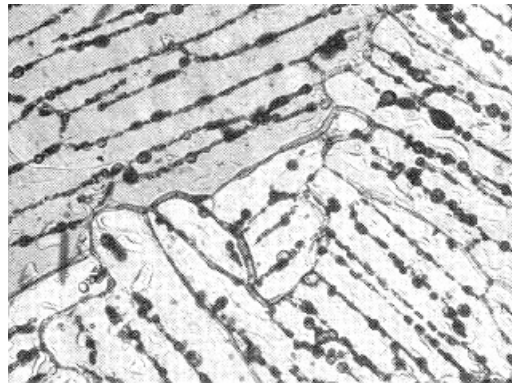
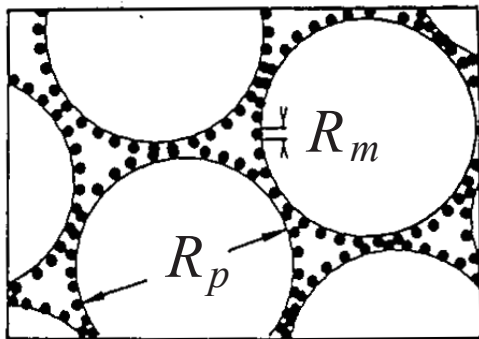
($1 \leq t \leq 2$, Golden, *Phys. Rev. Lett.* 1990 ; *Comm. Math. Phys.* 1992)

non-universal behavior in continuum

Continuum percolation model for stealthy materials applied to sea ice microstructure explains **Rule of Fives** and Antarctic data on **ice production** and **algal growth**

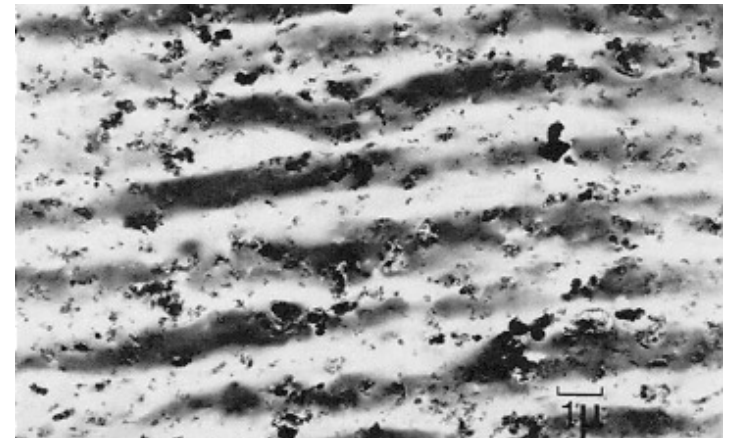
$$\phi_c \approx 5 \%$$

Golden, Ackley, Lytle, *Science*, 1998



compressed
powder

sea ice



microstructure of radar
absorbing composite



***rigorous bounds
percolation theory
hierarchical model
network model***

field data

X-ray tomography for
brine inclusions

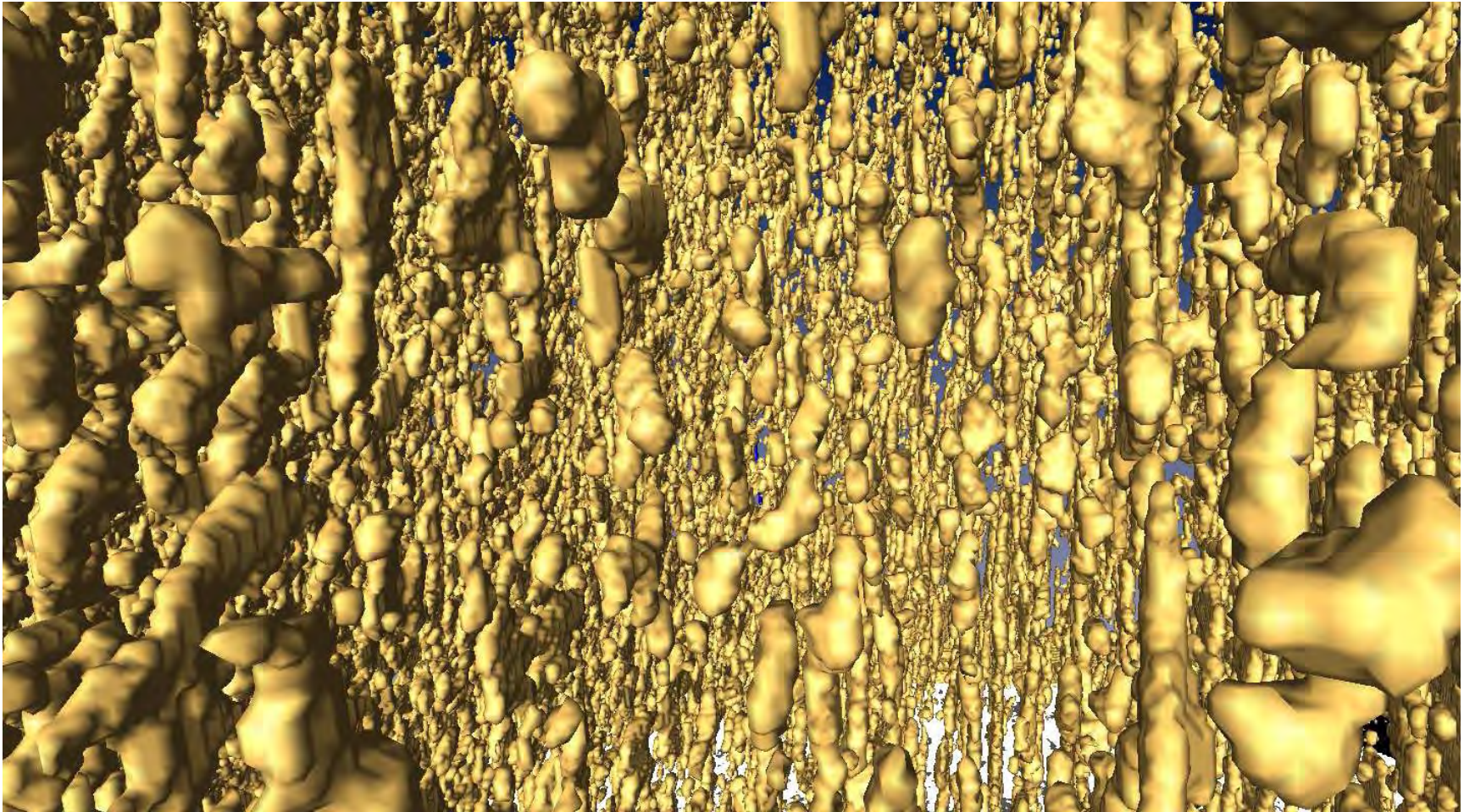
***unprecedented look
at thermal evolution
of brine phase and
its connectivity***

micro-scale
controls
macro-scale
processes

A unified approach to understanding permeability in sea ice • Solving the mystery of
booming sand dunes • Entering into the “greenhouse century”: A case study from Switzerland

X-ray computed tomography of brine inclusions in sea ice

~ 1 cm across

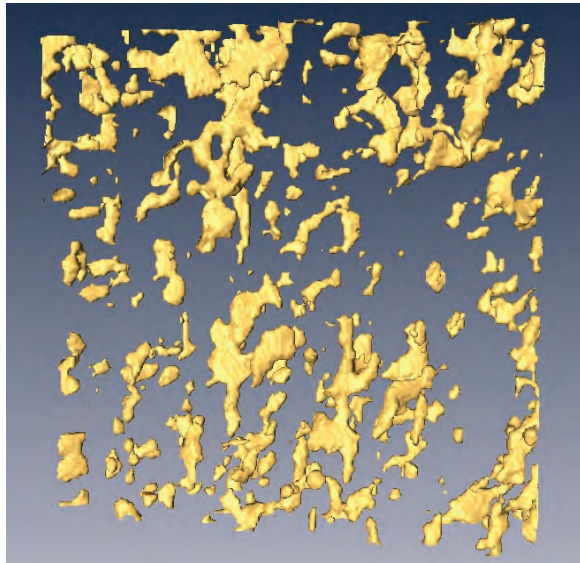


brine volume fraction $\phi = 5.7 \%$ $T = -8^{\circ}\text{C}$

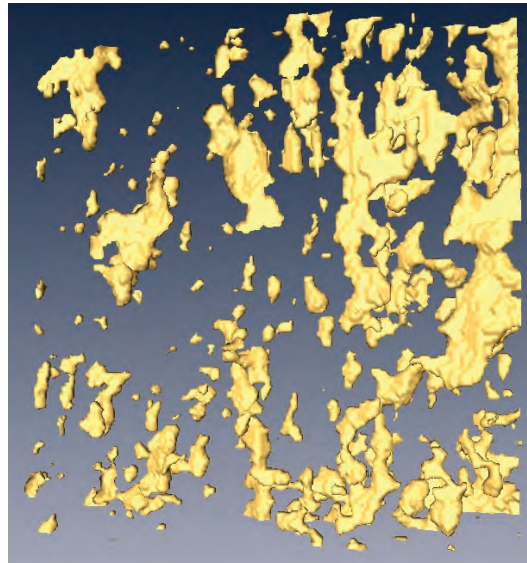
Golden, Eicken, Heaton, Miner, Pringle, Zhu, *Geophys. Res. Lett.* 2007

brine connectivity (over cm scale)

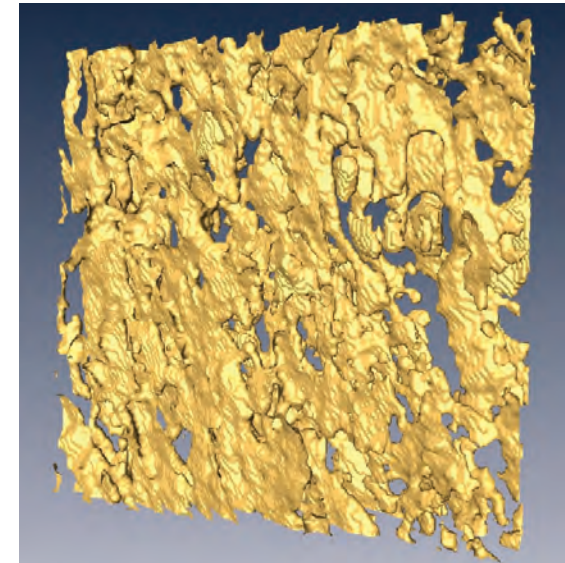
8 x 8 x 2 mm



-15 °C, $\phi = 0.033$



-6 °C, $\phi = 0.075$



-3 °C, $\phi = 0.143$

X-ray tomography confirms percolation threshold

3-D images
pores and throats



3-D graph
nodes and edges

analyze graph connectivity as function of temperature and sample size

- ***use finite size scaling techniques to confirm rule of fives***
- ***order parameter data from a natural material***

lattice and continuum percolation theories yield:

$$k(\phi) = k_0 (\phi - \phi_c)^2$$

critical
exponent
 t

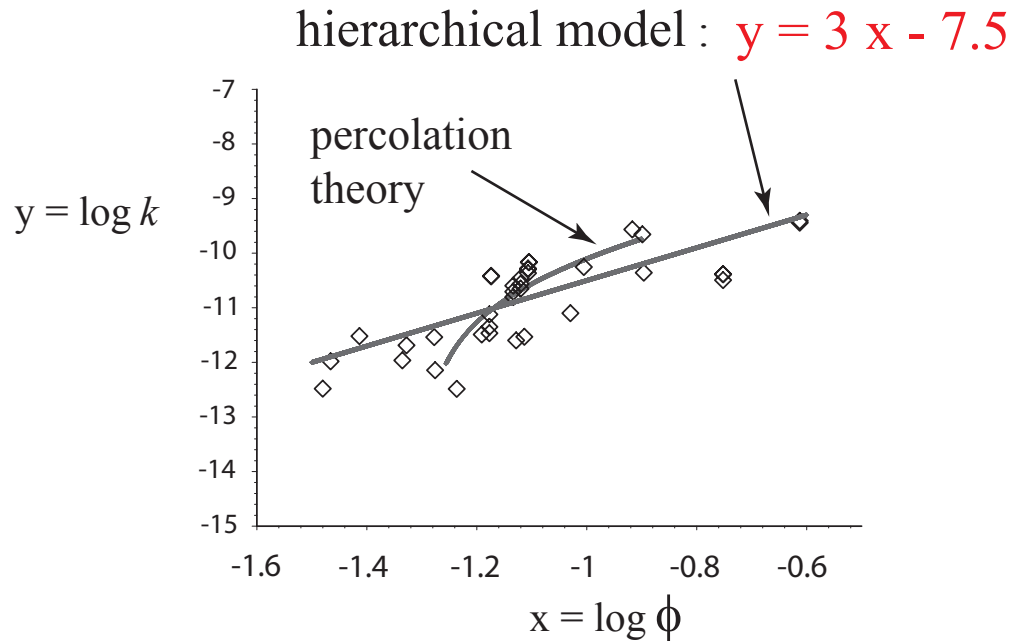
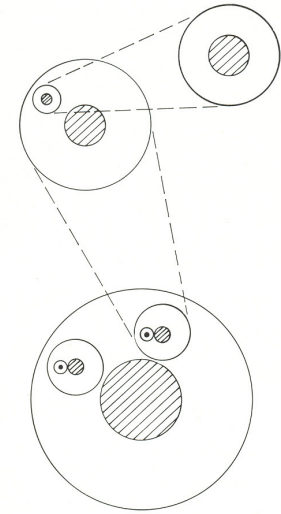
←

$$k_0 = 3 \times 10^{-8} \text{ m}^2$$

- exponent is **UNIVERSAL** lattice value $t \approx 2.0$ from general structure of brine inclusion distribution function (-- other saline ice?)
- **sedimentary rocks** like sandstones also exhibit universality
- **critical path analysis** -- developed for electronic hopping conduction -- yields scaling factor k_0
- no free parameters - microstructural input only

hierarchical and network models

**brine-coated
spherical ice grains**



$$k(\phi) = k_0 \phi^3$$

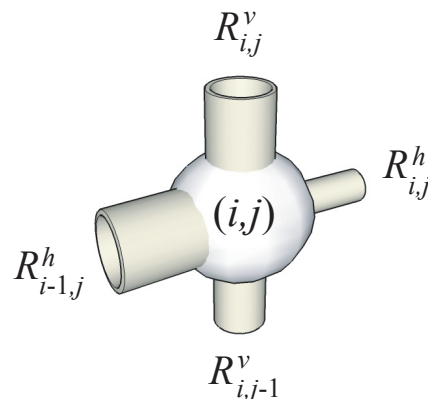
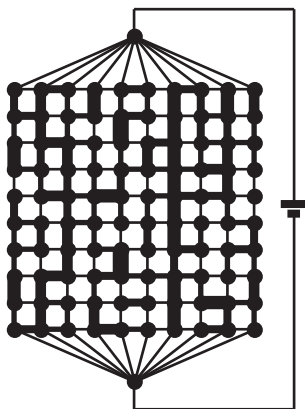
**self-similar model
used for porous rocks**

Sen, Scala, Cohen 1981

Sheng 1990

Wong, Koplick, Tomanic 1984

statistical best fit of data: $y = 3.05x - 7.50$



random pipe network with radii chosen from measured inclusion distributions, solved with fast multigrid method

Zhu, Jabini, Golden, Eicken, Morris, *Annals of Glaciology*, 2006

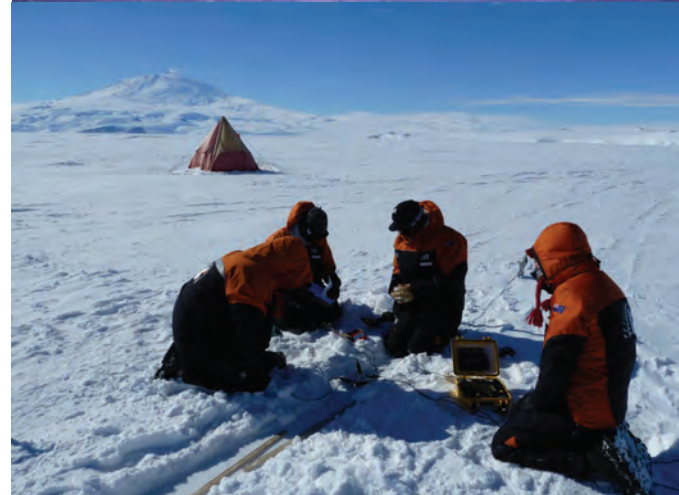
Golden et al., *Geophysical Research Letters*, 2007

Zhu, Golden, Gully and Sampson, *Physica B*, 2010

develop electromagnetic methods of monitoring fluid transport and microstructure

extensive measurements of fluid and
electrical transport properties of sea ice:

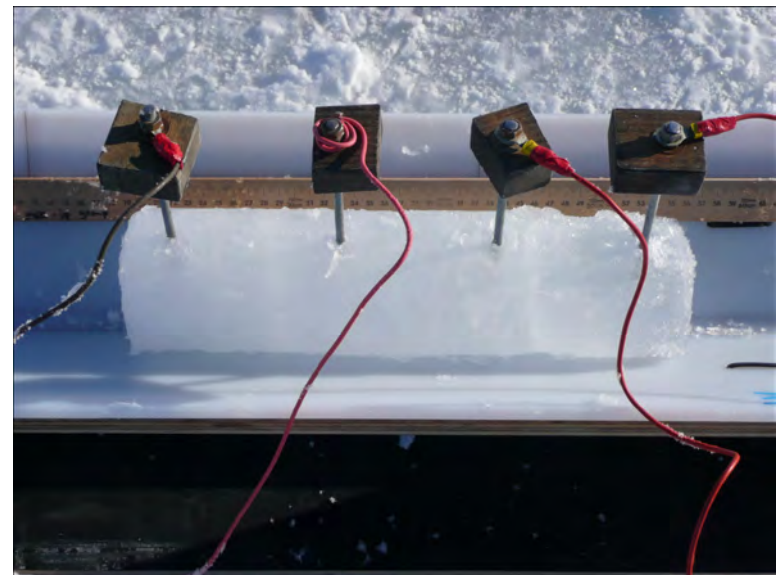
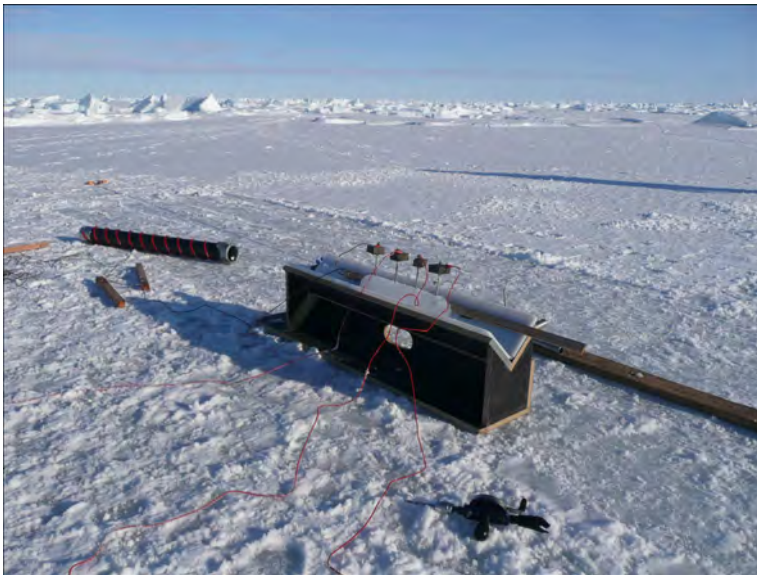
2007	Antarctic	SIPEX
2010	Arctic	Barrow AK
2010	Antarctic	McMurdo Sound
2011	Arctic	Barrow AK
2012	Arctic	Barrow AK
2012	Antarctic	SIPEX II



electrical measurements



Wenner array



vertical conductivity

Zhu, Golden, Gully, Sampson *Physica B* 2010

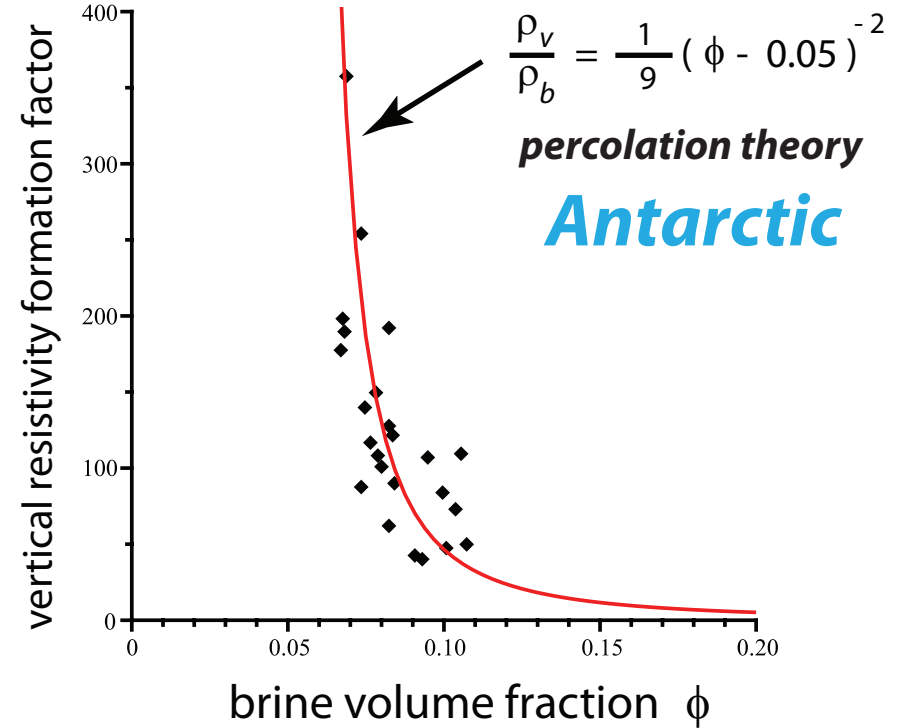
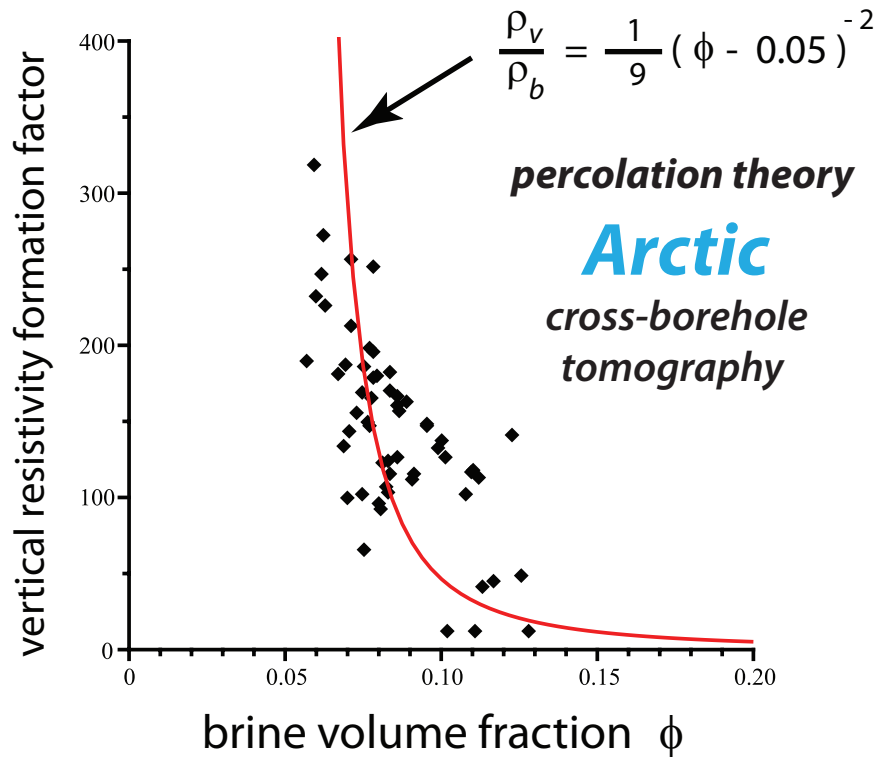
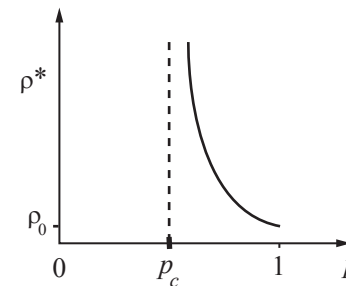
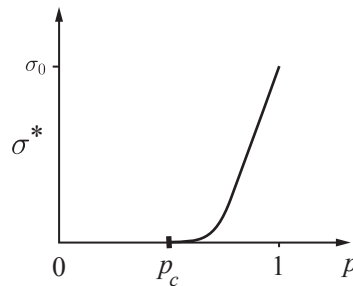
Sampson, Golden, Gully, Worby *Deep Sea Research* 2011

critical behavior of electrical transport in sea ice

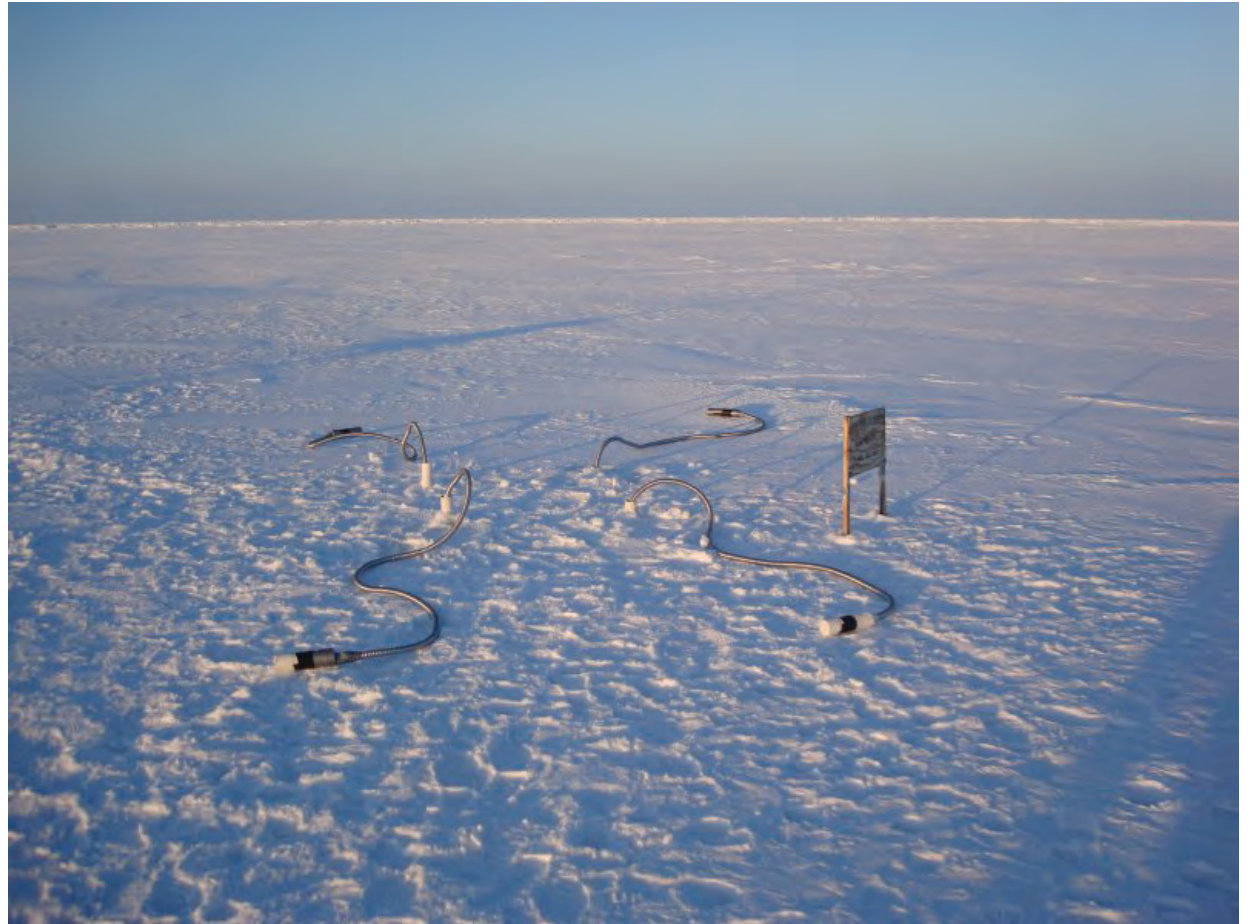
electrical signature of the on-off switch for fluid flow

same universal critical exponent as for fluid permeability

studied for over 50 years but no previous observations or theory of critical behavior

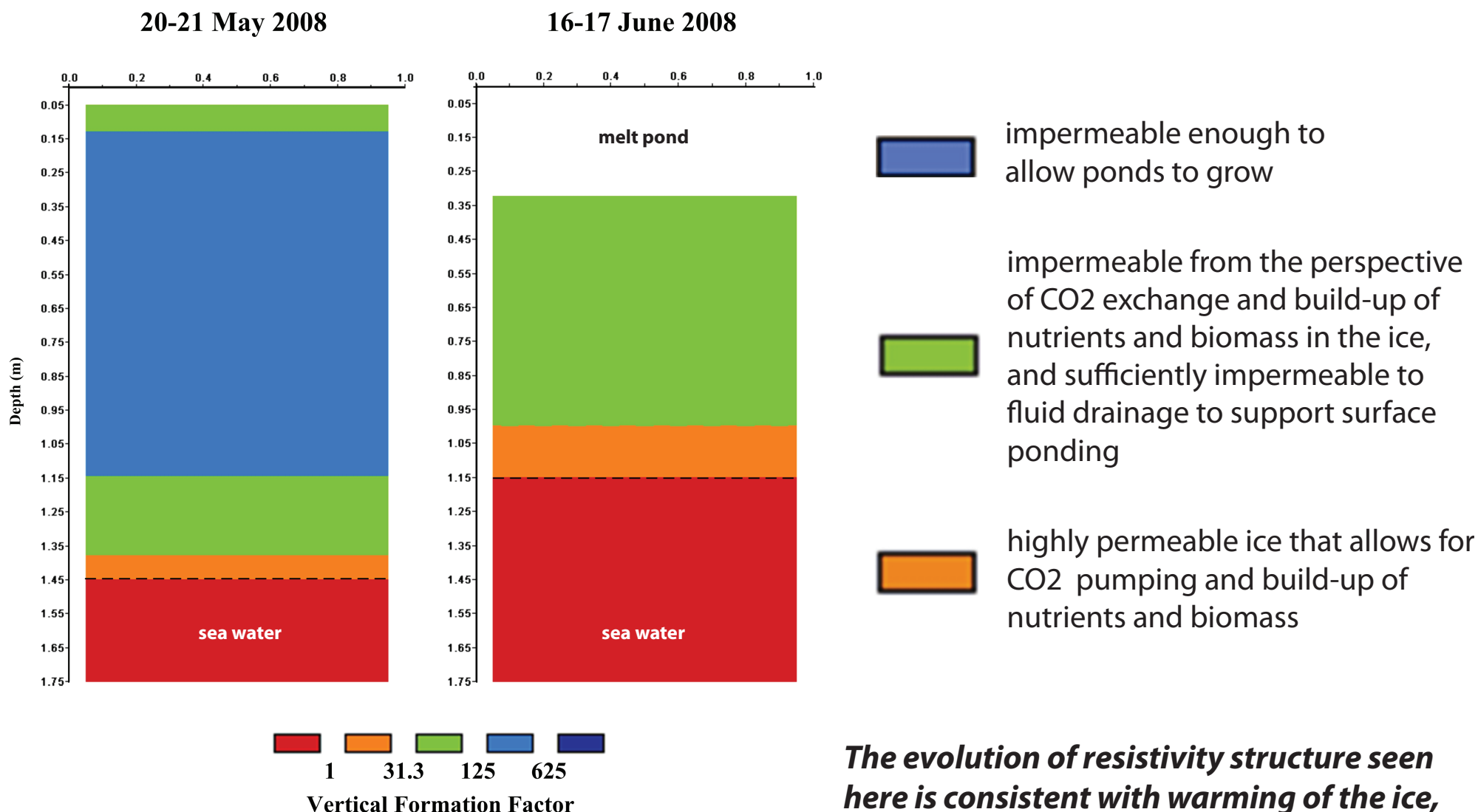


cross borehole tomography



***Ingham, Jones, Buchanan
Victoria University, Wellington, NZ***

Cross-borehole tomographic reconstructions of the vertical resistivity formation factor for Arctic sea ice before and after melt pond formation

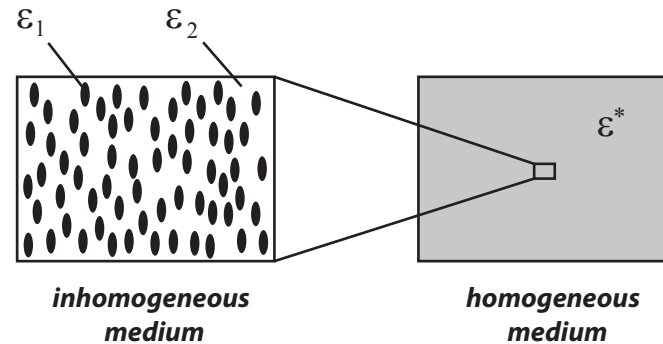


The evolution of resistivity structure seen here is consistent with warming of the ice, thus increasing the fluid permeability.

multiscale homogenization

Theory of Effective Electromagnetic Behavior of Composites

analytic continuation method



Forward Homogenization Bergman (1978), Milton (1979), Golden and Papanicolaou (1983)

composite geometry
(spectral measure μ) \longrightarrow ϵ^*

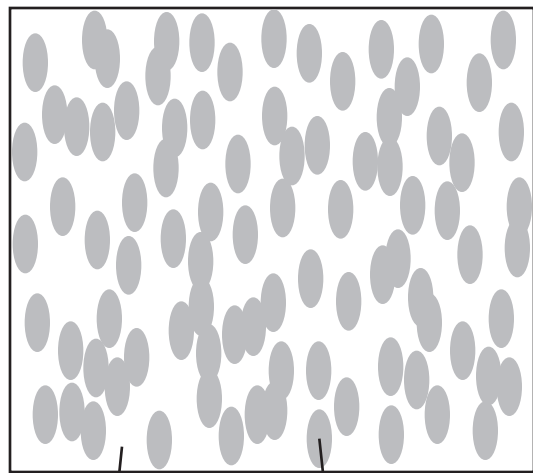
integral representations, rigorous bounds, approximations, etc.

Inverse Homogenization Cherkaev and Golden (1998), Day and Thorpe (1999), Cherkaev (2001)
(McPhedran, McKenzie, and Milton, 1982)

ϵ^* \longrightarrow **composite geometry**
(spectral measure μ)

recover brine volume fraction, connectivity, etc.

Effective complex permittivity of a two phase composite in the quasistatic (long wavelength) limit



ϵ_1

ϵ_2

ϵ^*

$$D = \epsilon E$$

$$\nabla \cdot D = 0$$

$$\nabla \times E = 0$$

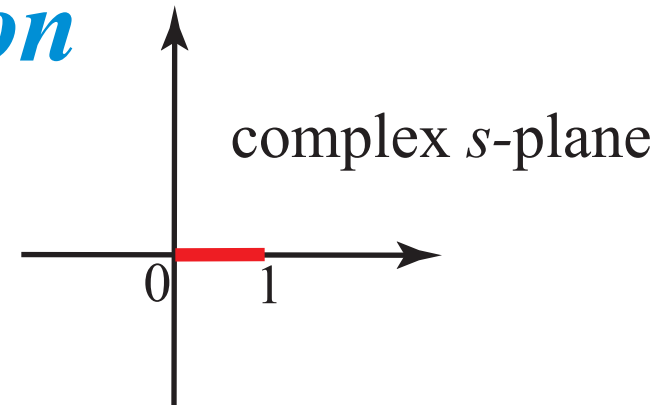
$$\langle D \rangle = \epsilon^* \langle E \rangle$$

p_1, p_2 = volume fractions of
the components

$$\epsilon^* = \epsilon^* \left(\frac{\epsilon_1}{\epsilon_2}, \text{ composite geometry } \right)$$

Stieltjes integral representation

$$F(s) = 1 - \frac{\epsilon^*}{\epsilon_2}, \quad s = \frac{1}{1 - \epsilon_1 / \epsilon_2}$$



$$F(s) = \int_0^1 \frac{d\mu(z)}{s - z}, \quad \mu$$

- spectral measure of self adjoint operator $\Gamma\chi$
- mass = p_1
- higher moments depend on n -point correlations

representation *separates*

GEOMETRY μ from
medium parameters in \mathcal{S}

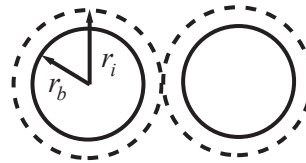
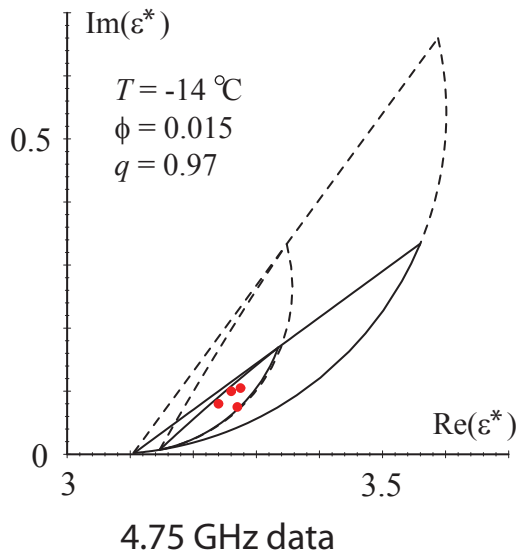
$$E = (s + \Gamma\chi)^{-1} e_k$$

$$\Gamma = \nabla(-\Delta)^{-1}\nabla.$$

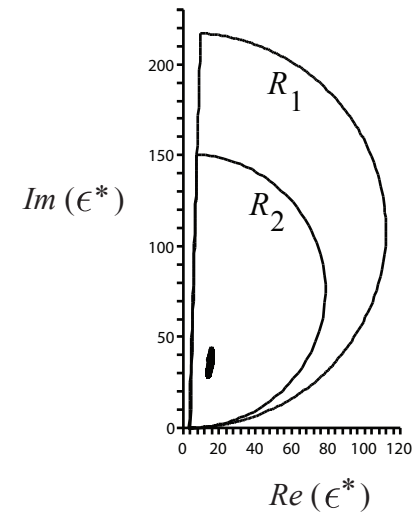
χ = indicator function
of medium 1

forward and inverse bounds for sea ice

matrix particle bounds



Golden 1997



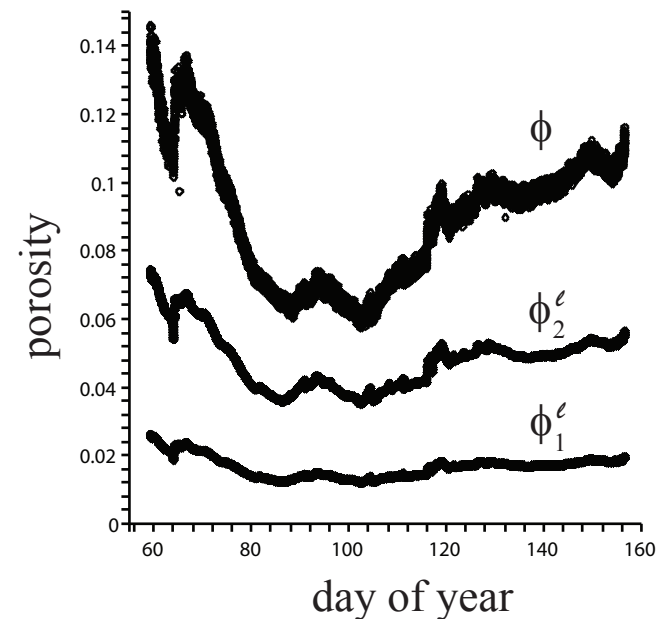
50 MHz capacitance probe data taken near Barrow, AK

inverse bounds and microstructural recovery

Gully, Backstrom, Eicken, Golden, Physica B, 2007

polycrystalline bounds

Gully, Lin, Cherkaev, Golden, 2012

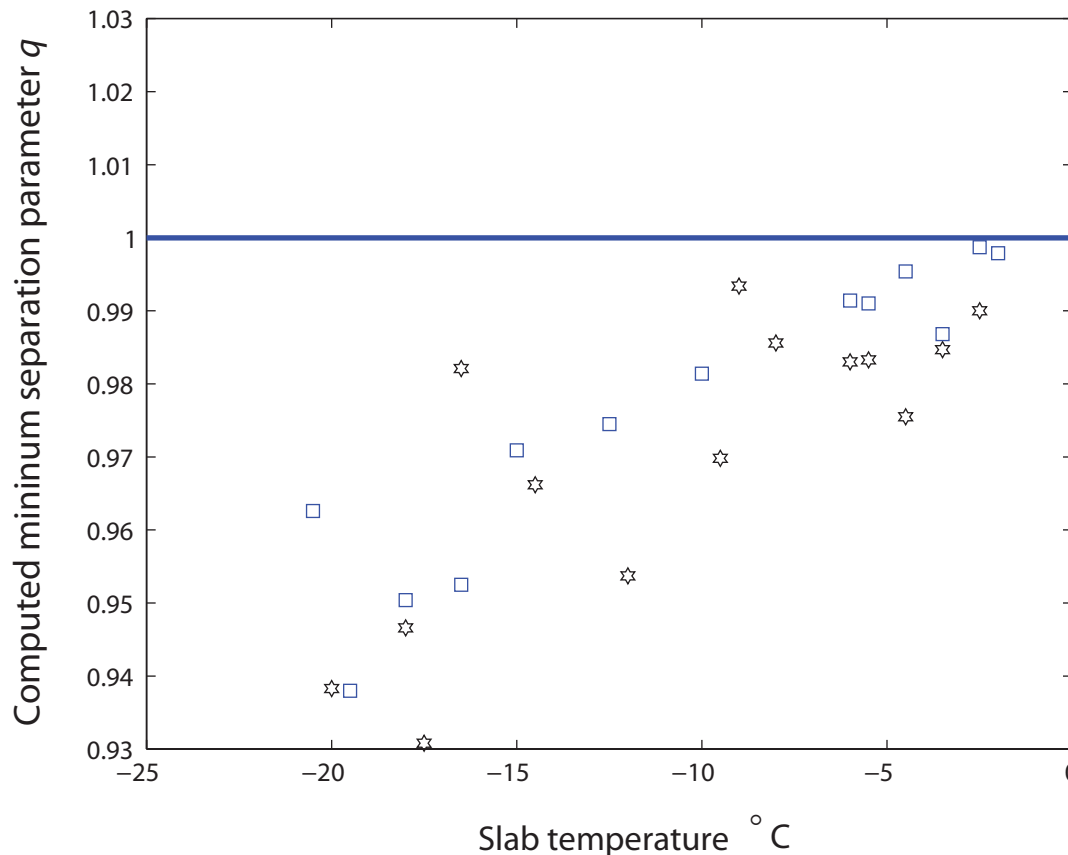


Recovery of inclusion separations in strongly heterogeneous composites from effective property measurements

Chris Orum, Elena Cherkaev, Ken Golden, Proc. Roy. Soc. A, 2012

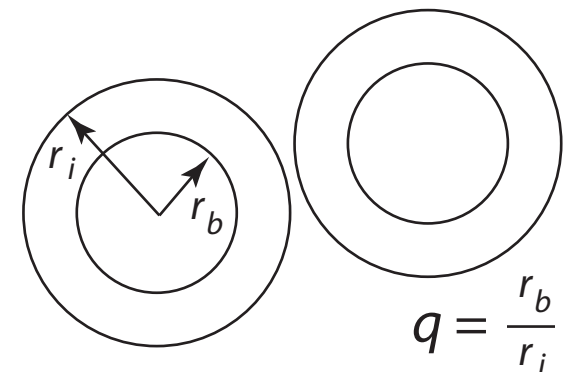
matrix particle composites (O. Bruno, 1991)

reduced spectral inversion -- construct algebraic curves which bound admissible region in (p,q) -space, q = separation parameter < 1

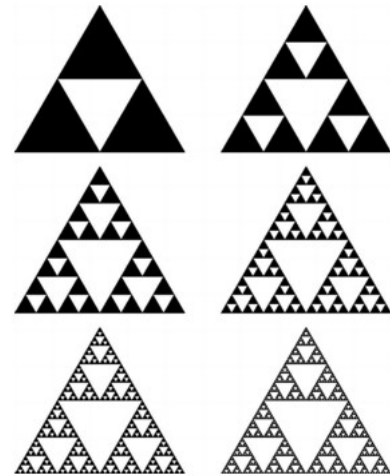


***rigorous inverse bound
on spectral gap***

***inversion for brine inclusion
separations in sea ice from
measurements of effective
complex permittivity***

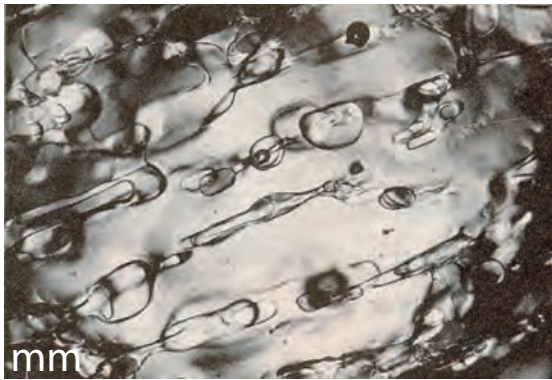


fractals and multiscale structure



sea ice displays *multiscale* structure over 10 orders of magnitude

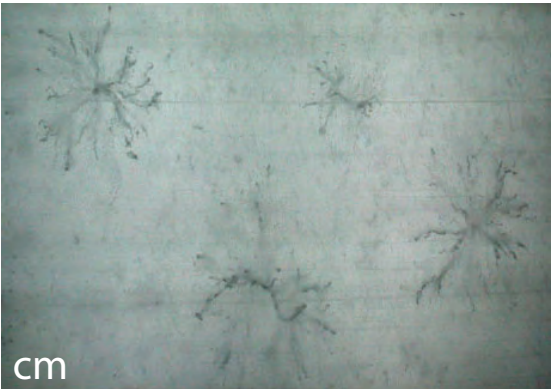
0.1 millimeter



brine inclusions



polycrystals



horizontal

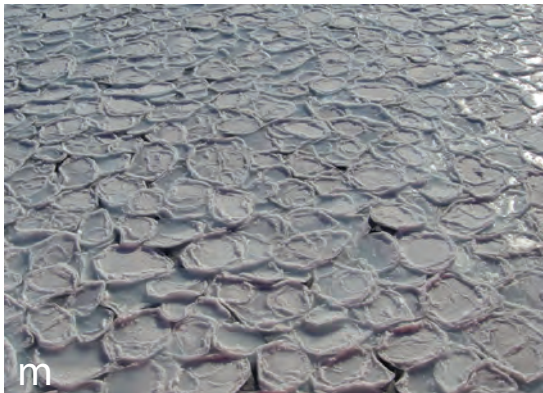


brine channels



vertical

1 meter

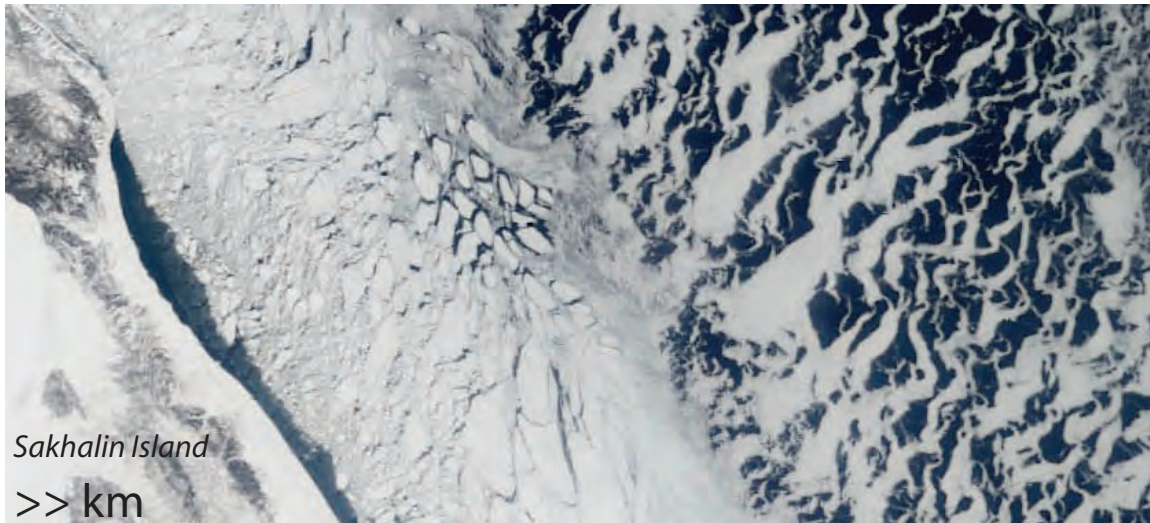
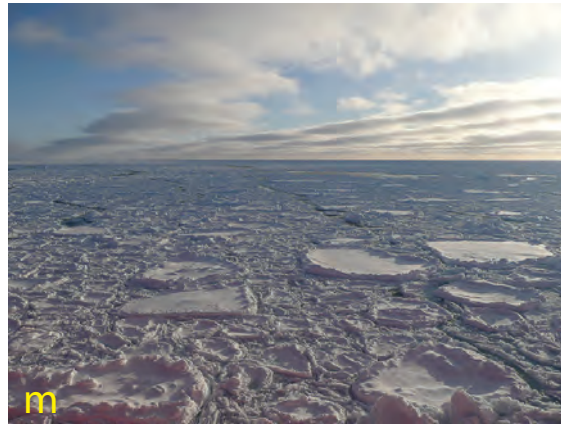


pancake ice

1 meter

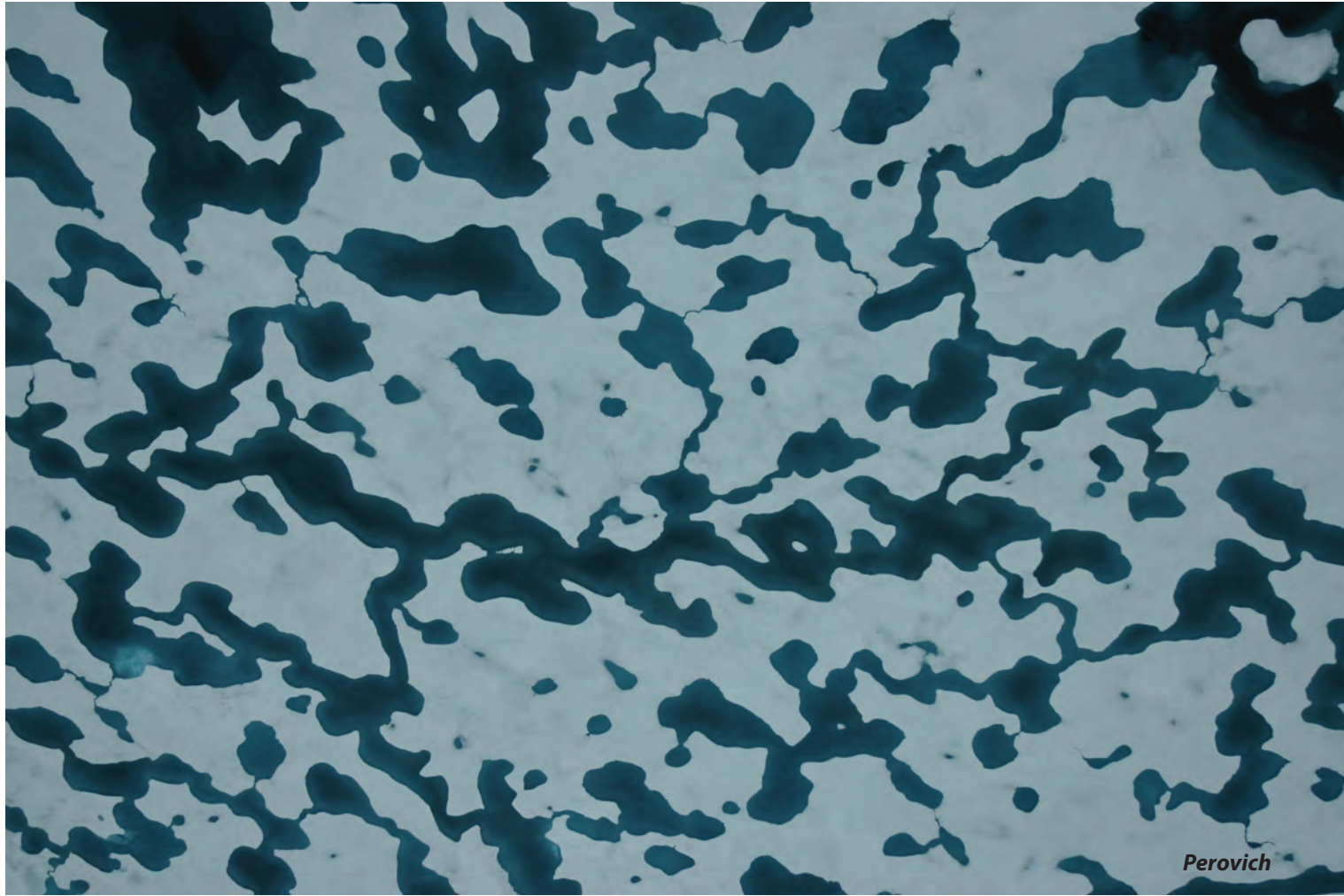


100 kilometers



melt pond formation and albedo evolution:

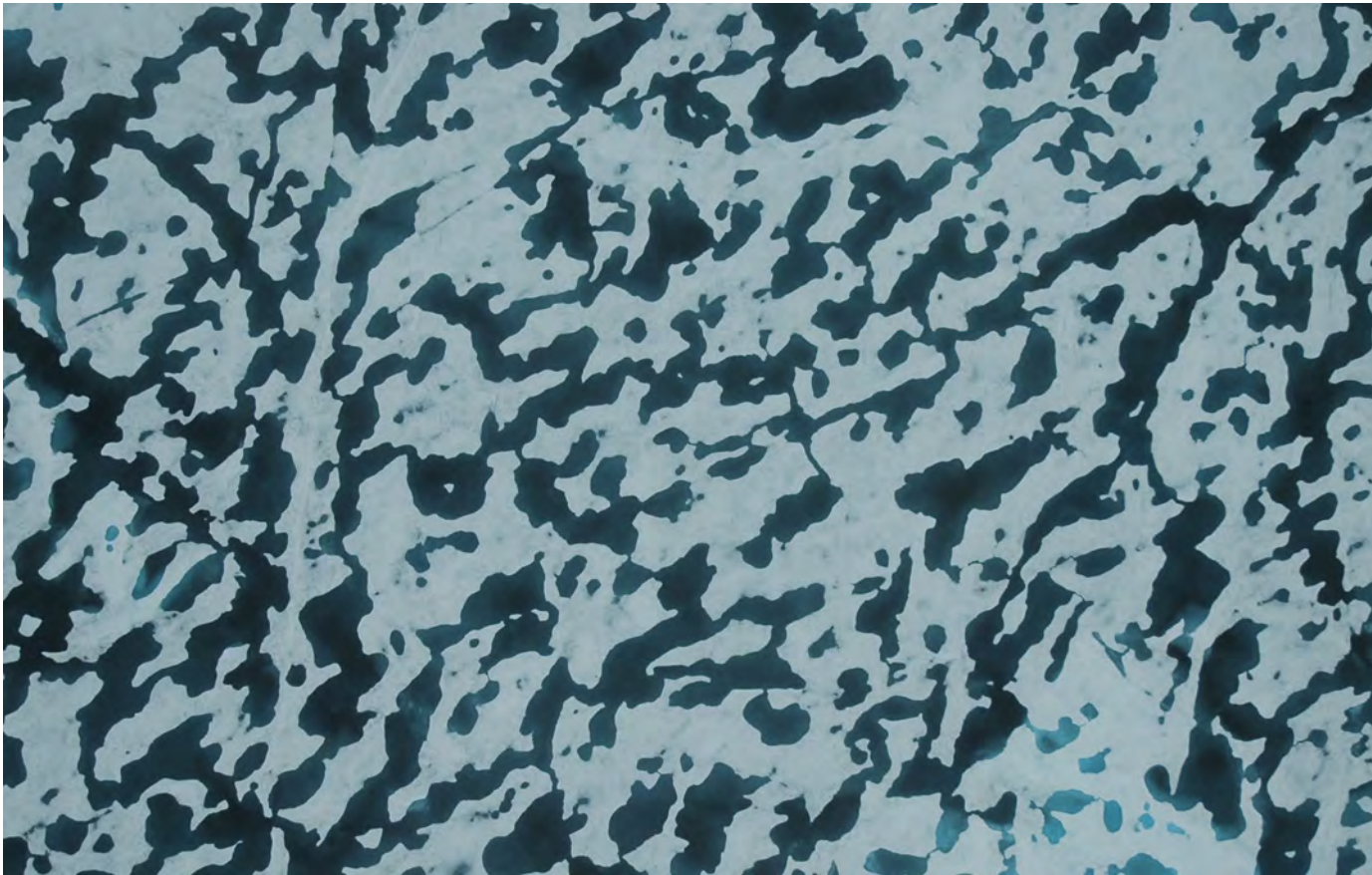
- *major drivers in polar climate*
- *key challenge for global climate models*



Perovich

Do melt ponds exhibit interesting multiscale structure?

Are there universal features of the evolution
similar to phase transitions in statistical mechanics?

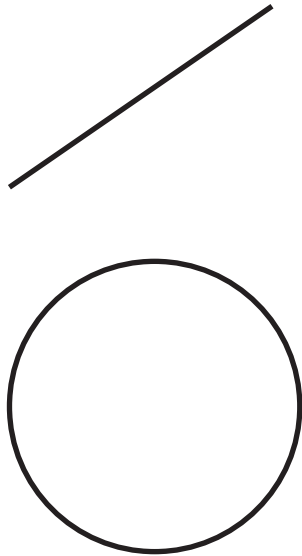


Transition in the fractal geometry of Arctic melt ponds

Christel Hohenegger, Bacim Alali, Kyle Steffen, Don Perovich, Ken Golden

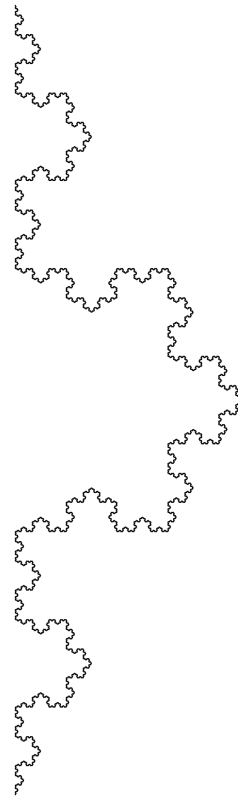
fractal curves in the plane

they wiggle so much that their dimension is >1



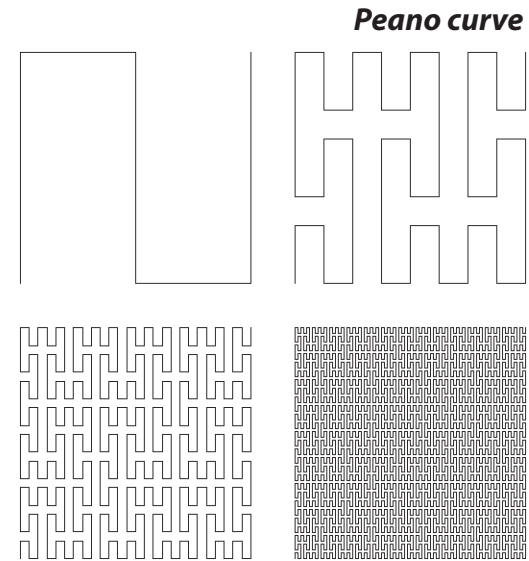
simple curves

$$D = 1$$

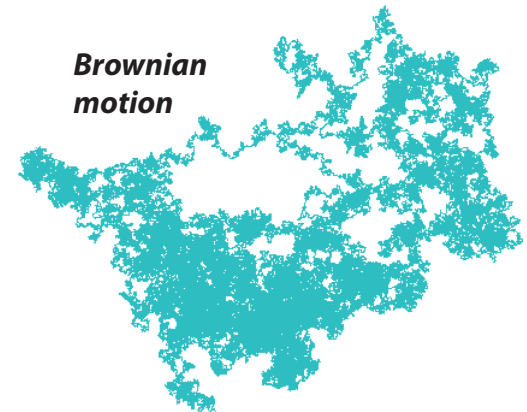


Koch snowflake

$$D = 1.26$$



Brownian motion



space filling curves

$$D = 2$$

clouds exhibit fractal behavior from 1 to 1000 km

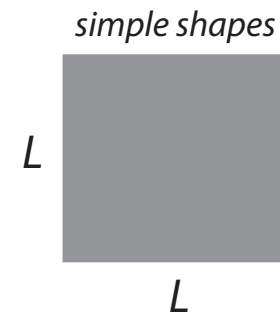
use **perimeter-area** data to find that
cloud and rain boundaries are fractals

$$D \approx 1.35$$

S. Lovejoy, Science, 1982



$$P \sim \sqrt{A}$$



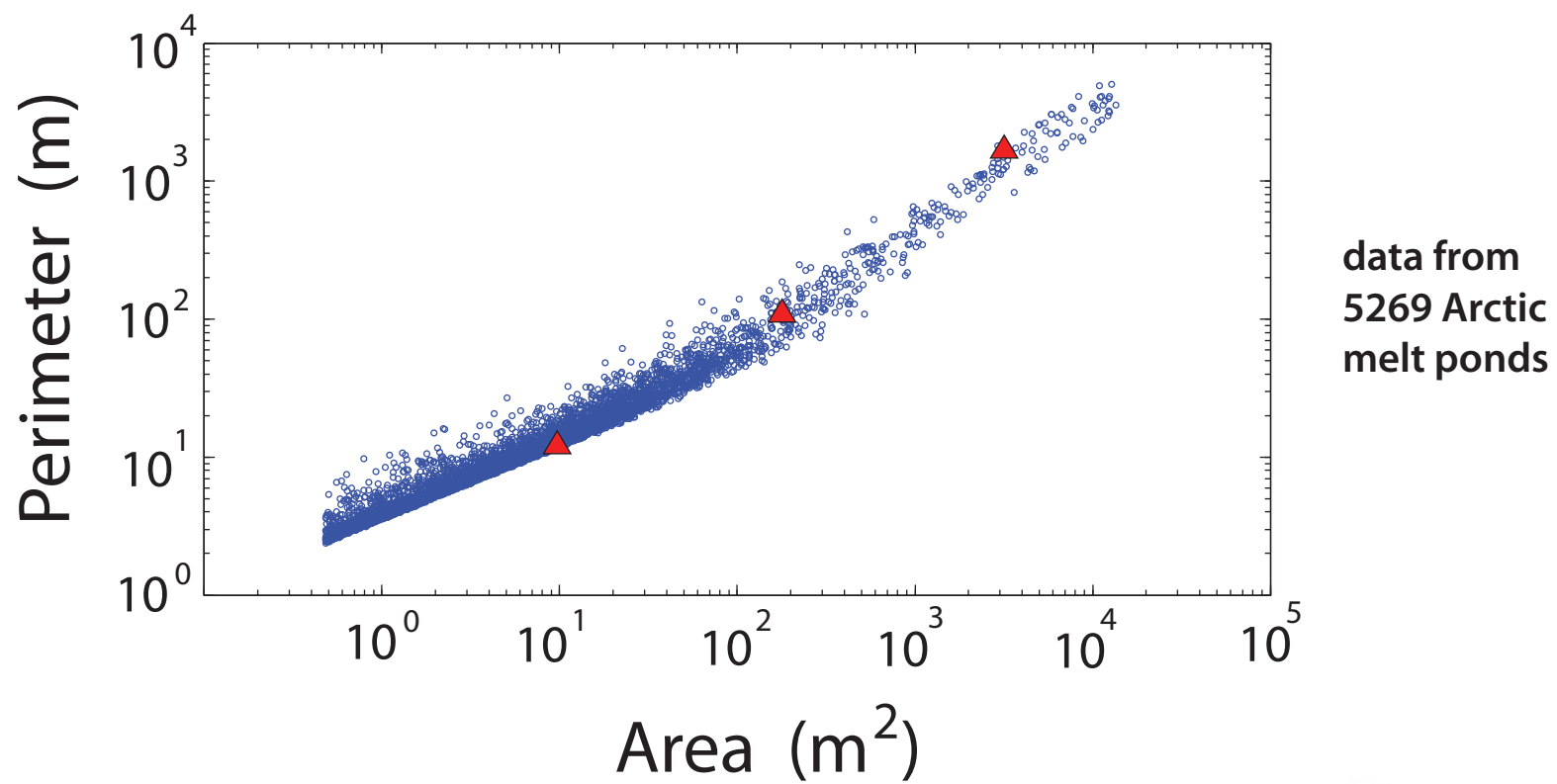
$$A = L^2$$
$$P = 4L = 4\sqrt{A}$$

$$P \sim \sqrt{A}^D$$



for fractals with
dimension D

$D = 1.52...$

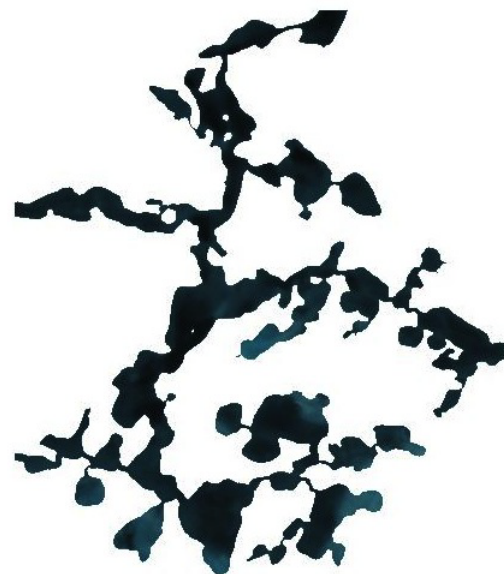


simple pond



~ 30 m

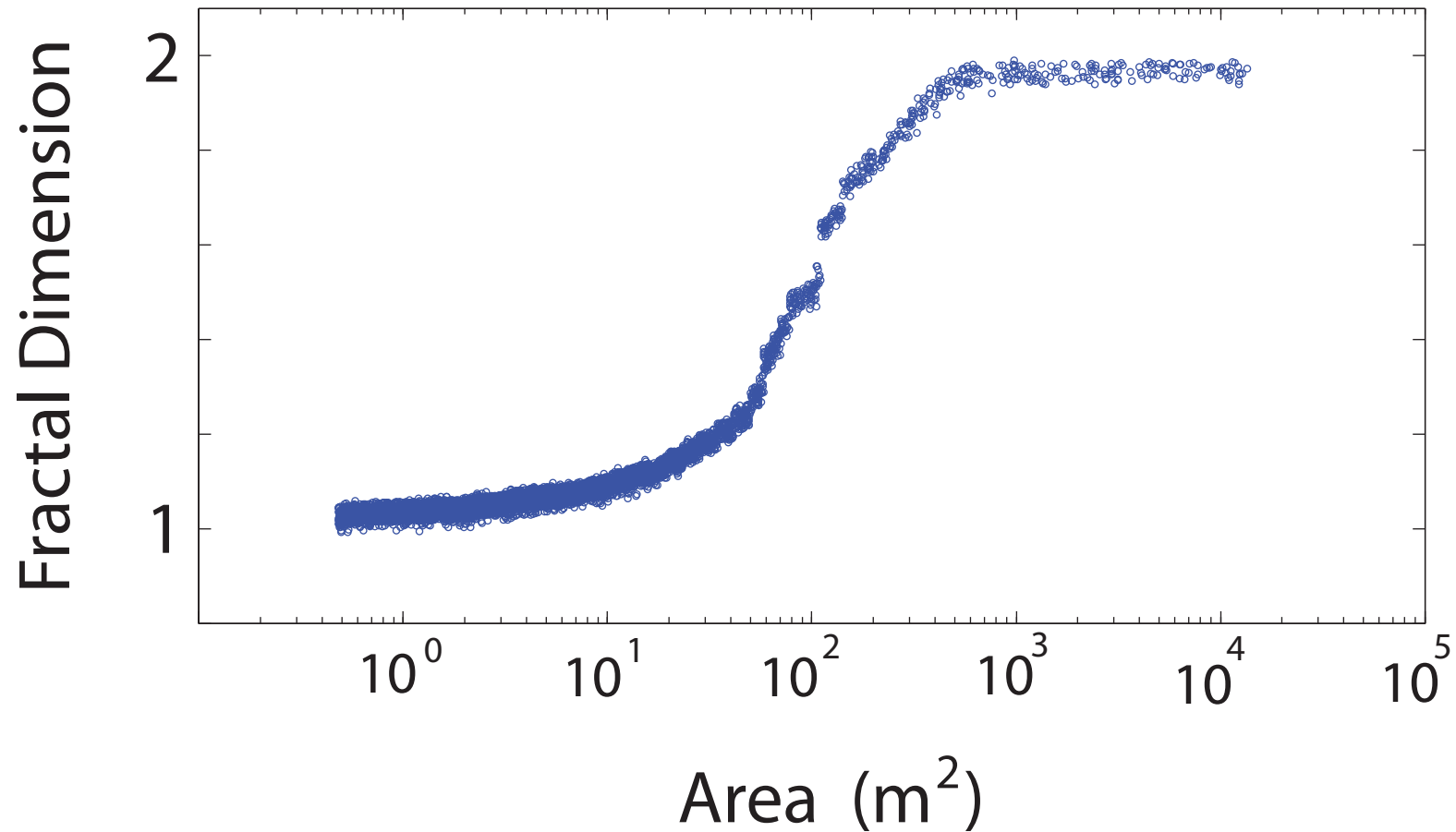
transitional pond



complex pond

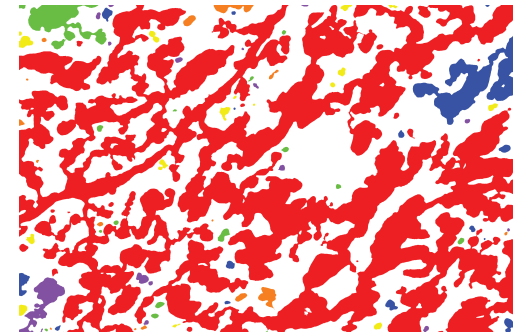
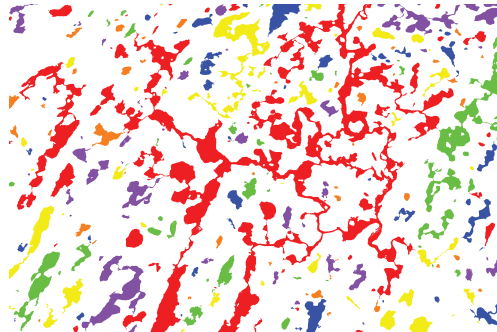
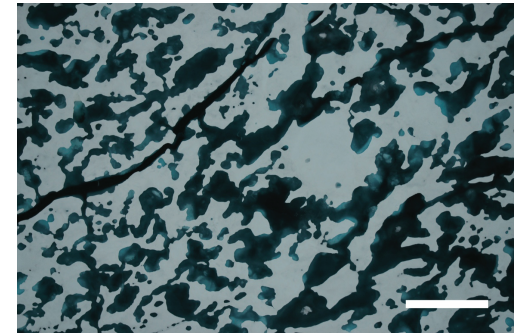
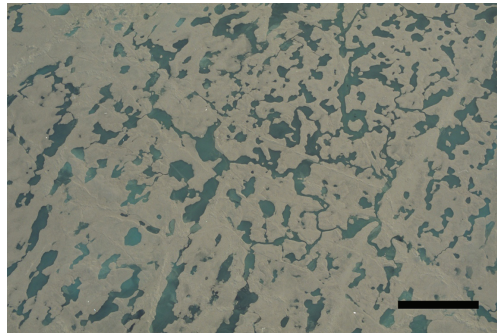
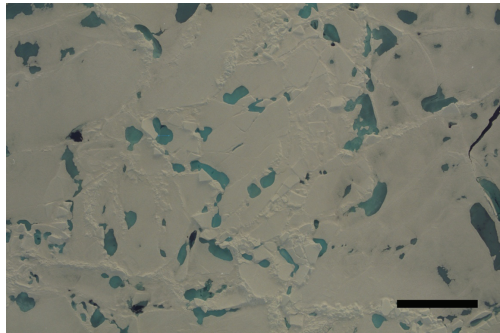
transition in the fractal dimension

complexity grows with length scale



compute “derivative” of area - perimeter data

***small simple ponds coalesce to form
large connected structures with complex boundaries***



melt pond percolation

THANK YOU

National Science Foundation

Division of Mathematical Sciences

Arctic Natural Sciences

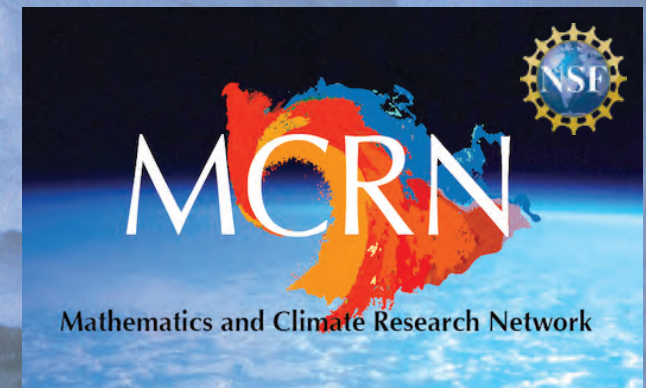
Office of Polar Programs

CMG Program

(Collaboration in Mathematical Geosciences)

VIGRE Program

REU Program



Buchanan Bay, Antarctica Mertz Glacier Polynya Experiment July 1999