

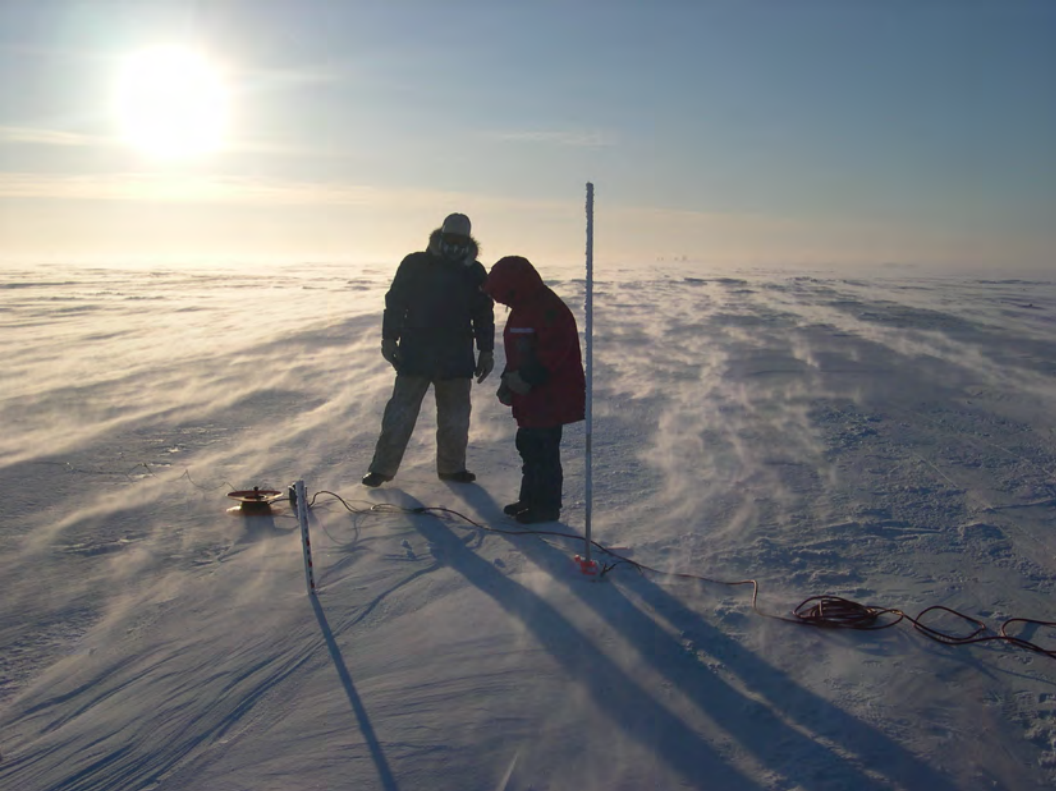
# On Thinning Ice: Modeling Sea Ice in a Warming Climate

**Kenneth M. Golden**  
**Department of Mathematics, University of Utah**



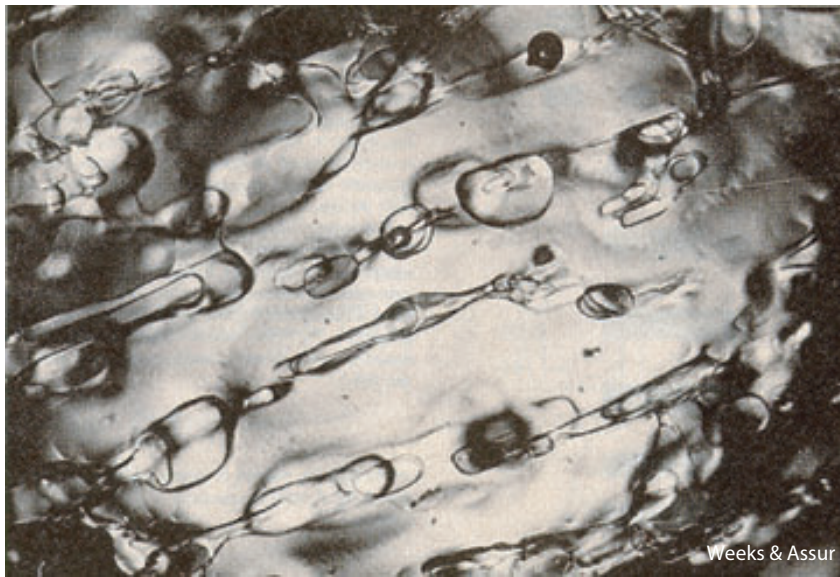
Antarctic Sea Ice and Southern Ocean Seminar  
16 February 2022



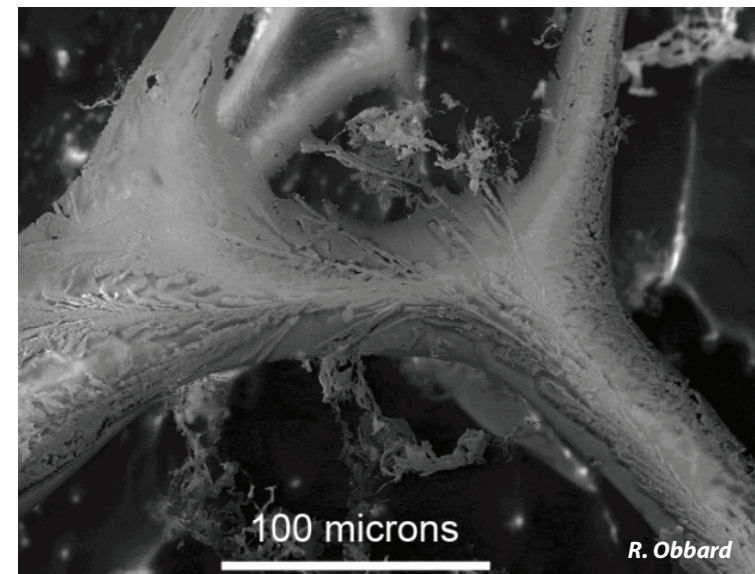


*sea ice may appear to be a  
barren, impermeable cap ...*





**brine inclusions in sea ice (mm)**



**micro - brine channel (SEM)**

***sea ice is a  
porous composite***

pure ice with brine, air, and salt inclusions

**brine channels (cm)**



horizontal section



vertical section



# fluid flow through the porous microstructure of sea ice governs key processes in polar climate and ecosystems

*evolution of Arctic melt ponds and sea ice albedo*



*nutrient flux for algal communities*



T. Maksym and T. Markus, 2008

*Antarctic surface flooding  
and snow-ice formation*

September  
snow-ice  
estimates

- *evolution of salinity profiles*
- *ocean-ice-air exchanges of heat, CO<sub>2</sub>*



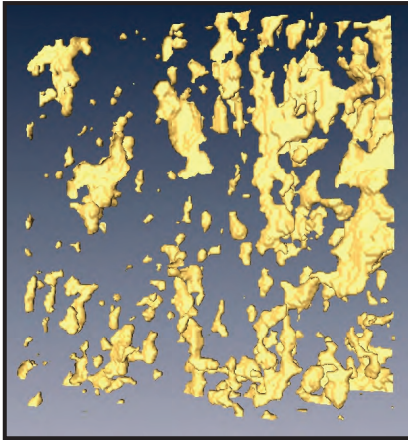
# Sea Ice is a Multiscale Composite Material

## *microscale*

brine inclusions

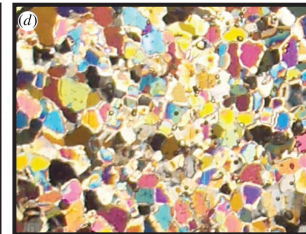
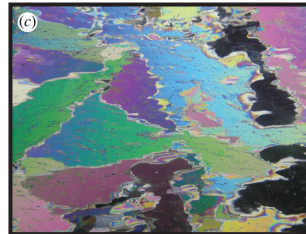
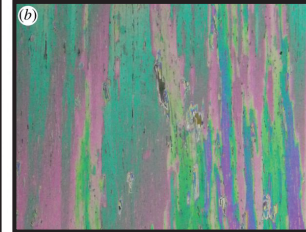


Weeks & Assur 1969



H. Eicken  
Golden et al. GRL 2007

polycrystals

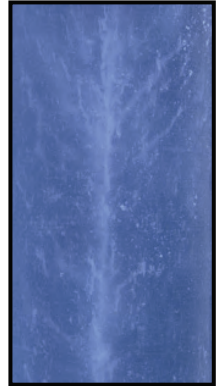


Gully et al. Proc. Roy. Soc. A 2015

brine channels



D. Cole



K. Golden

**millimeters**

**centimeters**

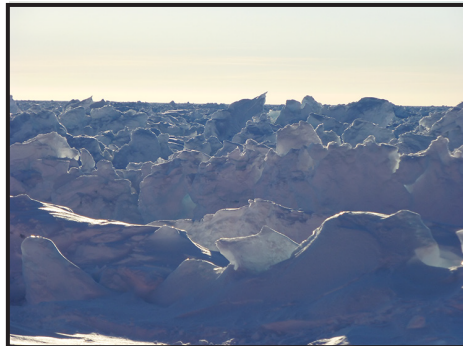
## *mesoscale*

Arctic melt ponds



K. Frey

Antarctic pressure ridges



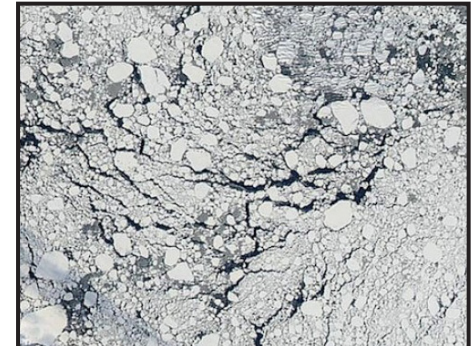
K. Golden

sea ice floes



J. Weller

sea ice pack



NASA

**meters**

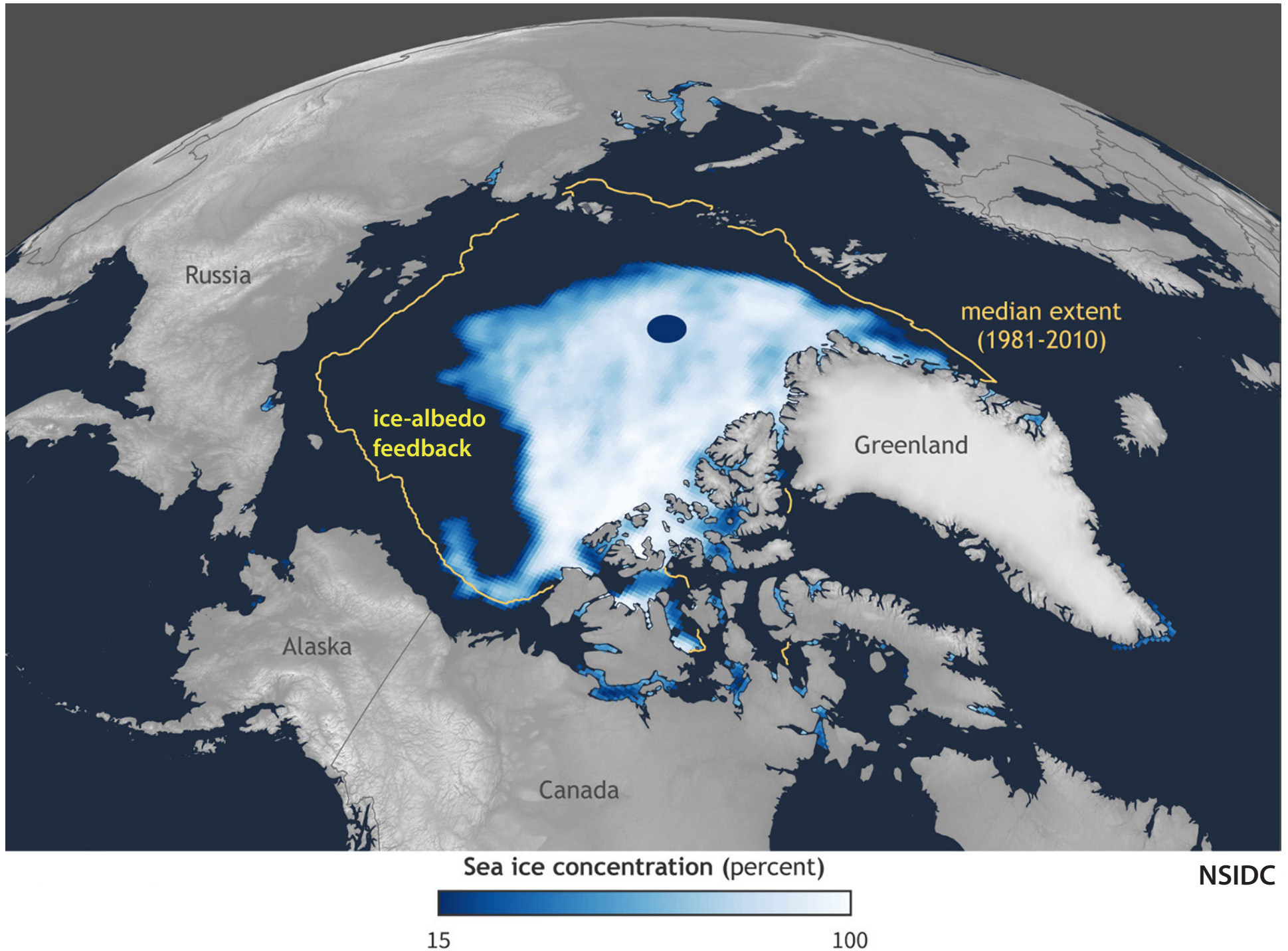
**kilometers**

## *macroscale*



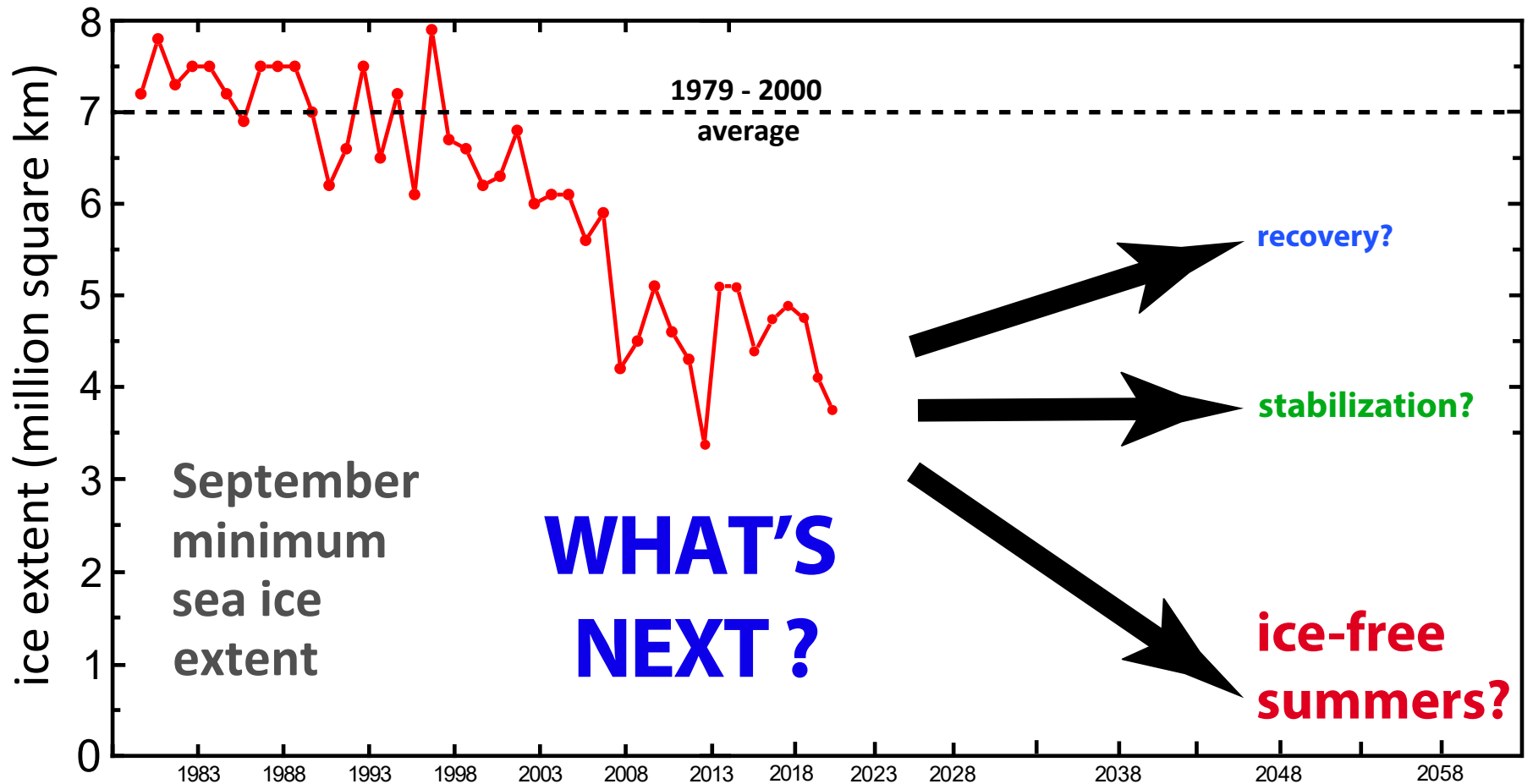
# Arctic sea ice extent

## September 15, 2020



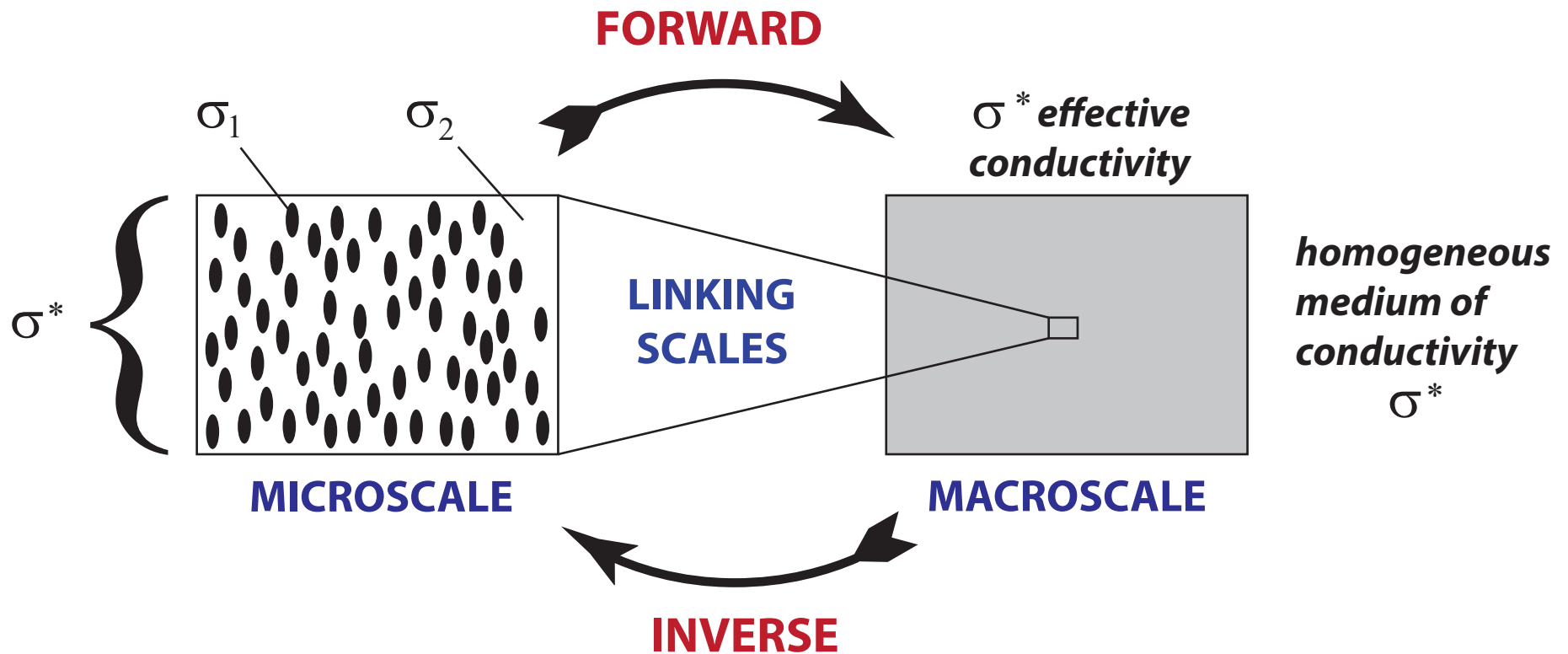


# *Predicting what may come next requires lots of math modeling.*





# ***HOMOGENIZATION for Composite Materials***



*Maxwell 1873 : effective conductivity of a dilute suspension of spheres*

*Einstein 1906 : effective viscosity of a dilute suspension of rigid spheres in a fluid*

*Wiener 1912 : arithmetic and harmonic mean **bounds** on effective conductivity*

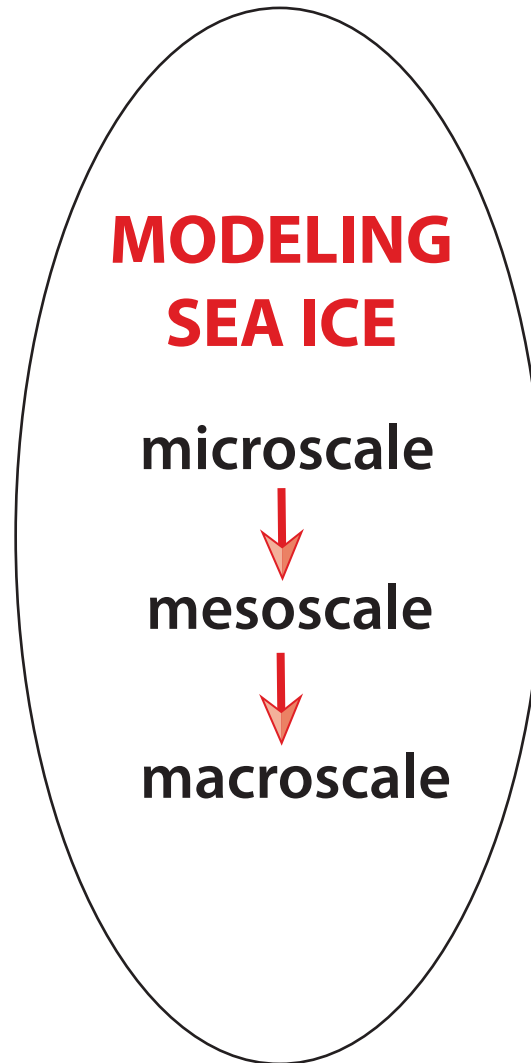
*Hashin and Shtrikman 1962 : variational **bounds** on effective conductivity*

widespread use of composites in late 20th century due in large part to advances in mathematically predicting their effective properties



# What is this talk about?

Using methods of **homogenization and statistical physics** to model sea ice effective behavior and advance representation of sea ice in climate models, process studies, ...



**A tour of key sea ice processes on micro, meso, and macro scales.**



# What is our research about?

Using methods of **homogenization and statistical physics** to model sea ice effective behavior and advance representation of sea ice in climate models, process studies, ...

## *Inputs, Ingredients*

### COMPOSITE MATERIALS

electrical engineering,  
stealth technology

porous media,  
oil extraction

statistical mechanics  
of ferromagnets

Anderson localization,  
semiconductor physics

random matrix theory

differential equations



### MODELING SEA ICE

microscale

mesoscale

macroscale

## *Outputs, Impacts*

### CLIMATE MODELING

sea ice physics  
& biology

composites,  
polycrystals

remote sensing

advection diffusion

biomedical imaging,  
biomaterials, EPS

polar microbial ecology



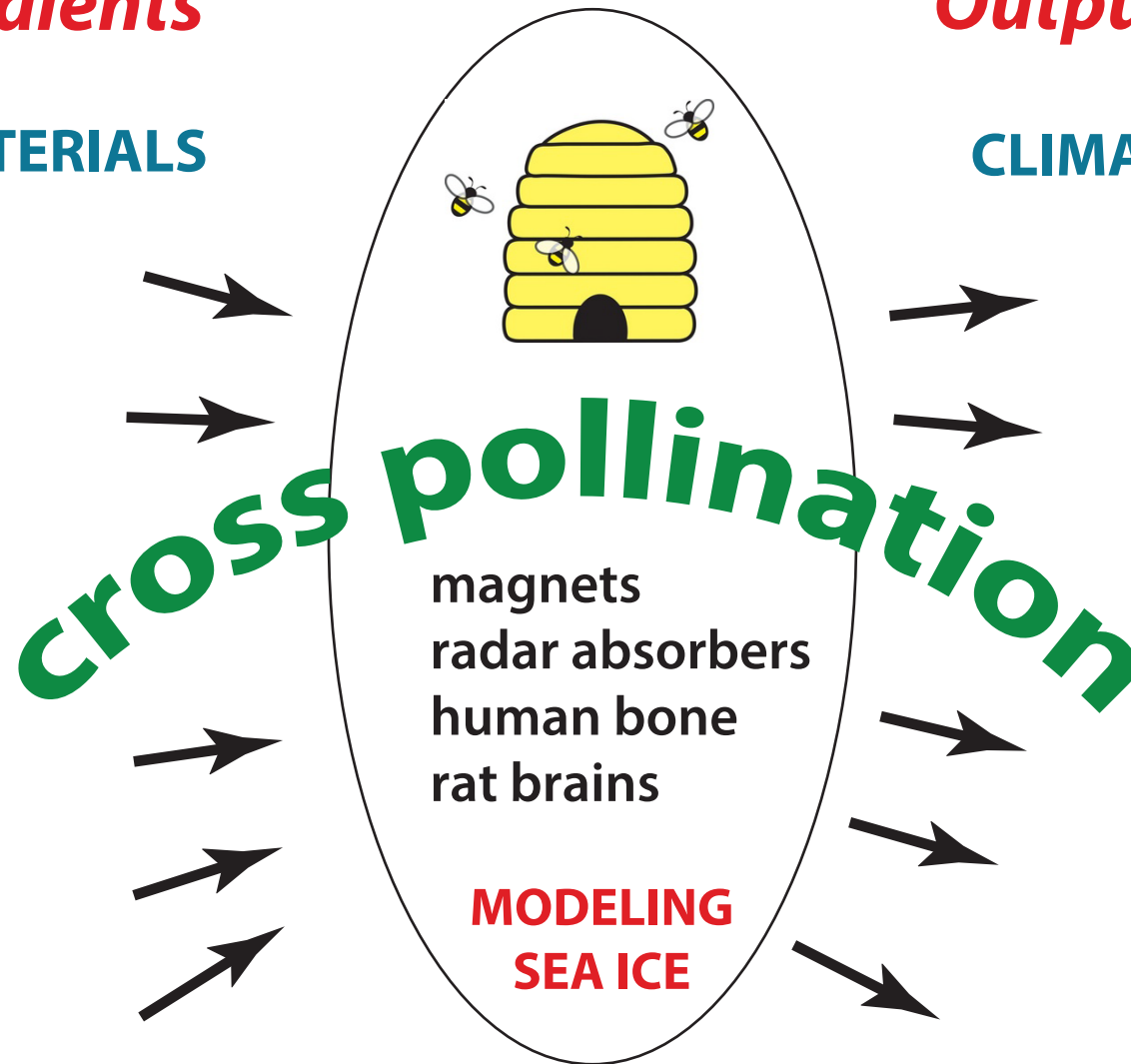
# What is our research about?

*Inputs, Ingredients*

COMPOSITE MATERIALS

*Outputs, Impacts*

CLIMATE MODELING

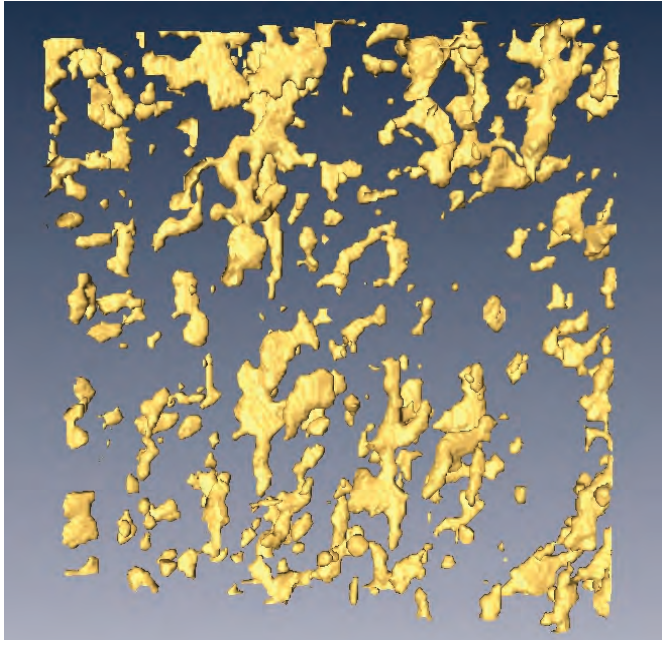


*Modeling sea ice drives advances in many areas of science and engineering.*

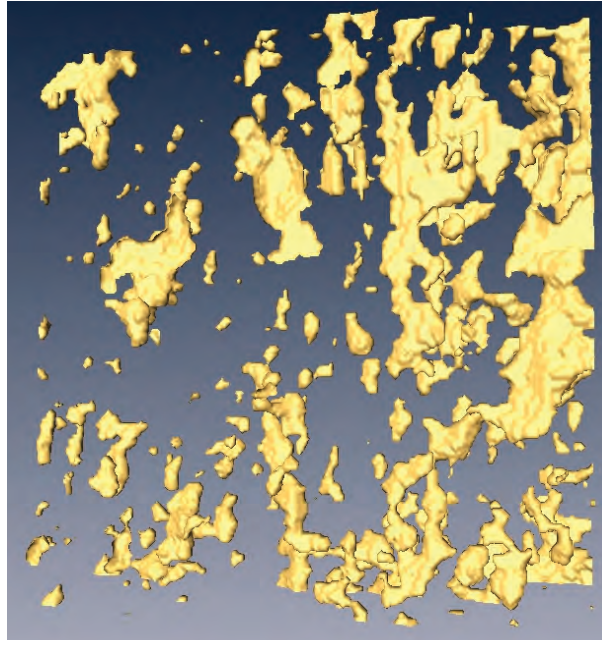


**microscale**

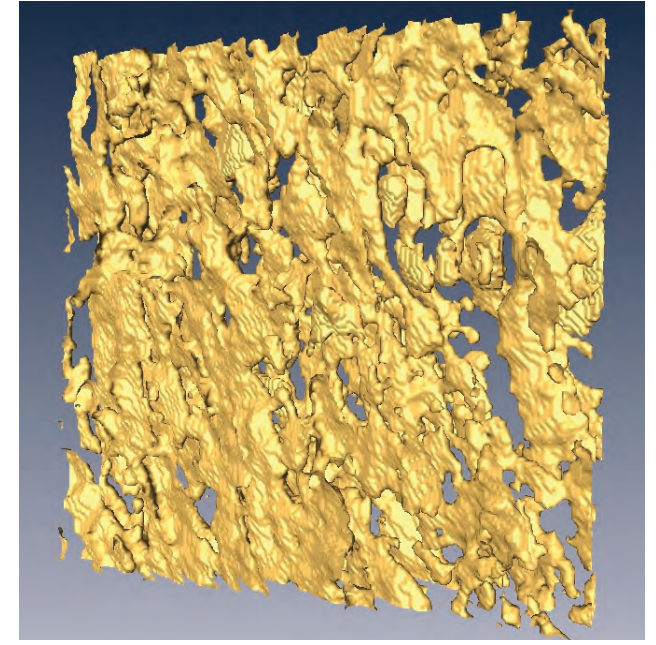
brine volume fraction and **connectivity** increase with temperature



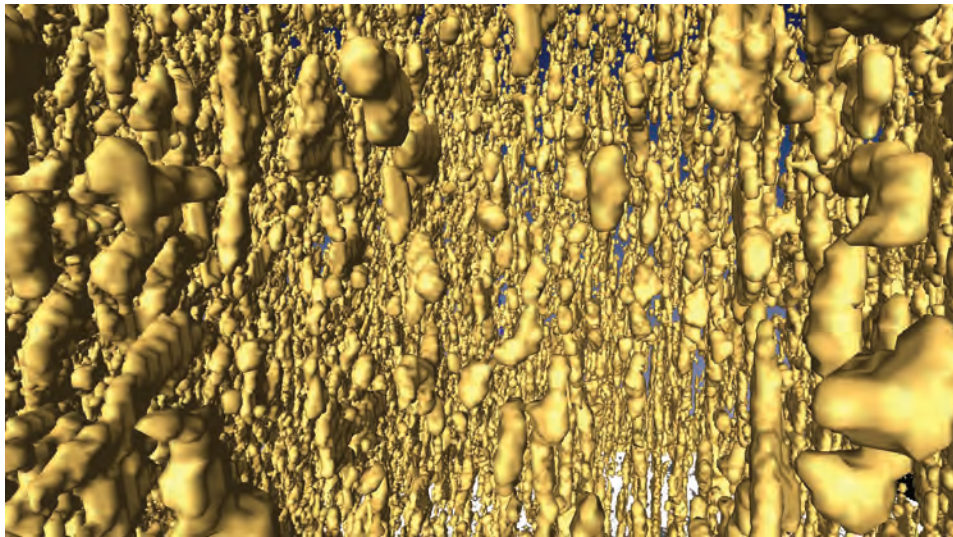
$T = -15\text{ }^{\circ}\text{C}$ ,  $\phi = 0.033$



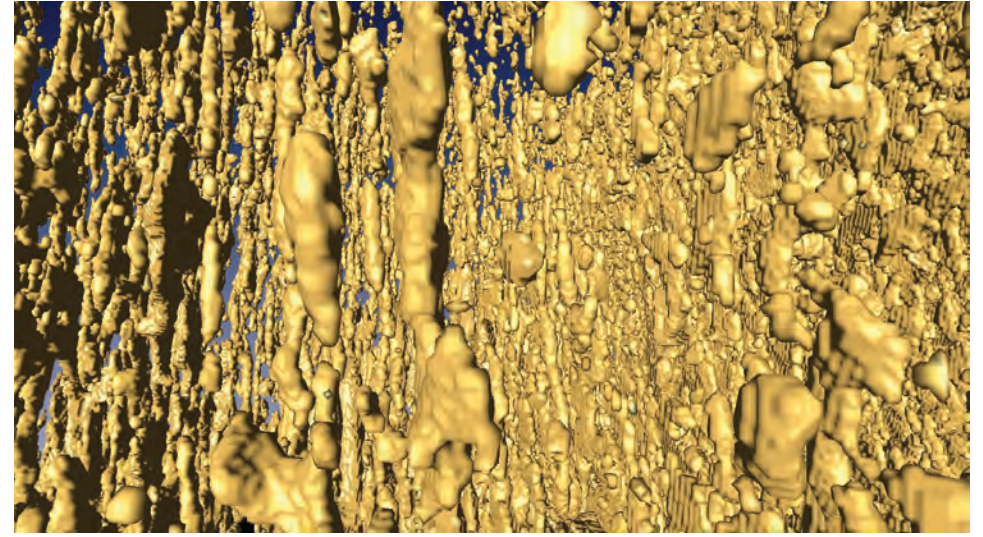
$T = -6\text{ }^{\circ}\text{C}$ ,  $\phi = 0.075$



$T = -3\text{ }^{\circ}\text{C}$ ,  $\phi = 0.143$



$T = -8\text{ }^{\circ}\text{C}$ ,  $\phi = 0.057$



$T = -4\text{ }^{\circ}\text{C}$ ,  $\phi = 0.113$

***X-ray tomography for brine in sea ice***

Golden et al., *Geophysical Research Letters*, 2007



# fluid permeability of a porous medium



how much water gets through the sample per unit time?

## *Darcy's Law*

for slow viscous flow in a porous medium

averaged fluid velocity

pressure gradient

$$\mathbf{v} = -\frac{\mathbf{k}}{\eta} \nabla p$$

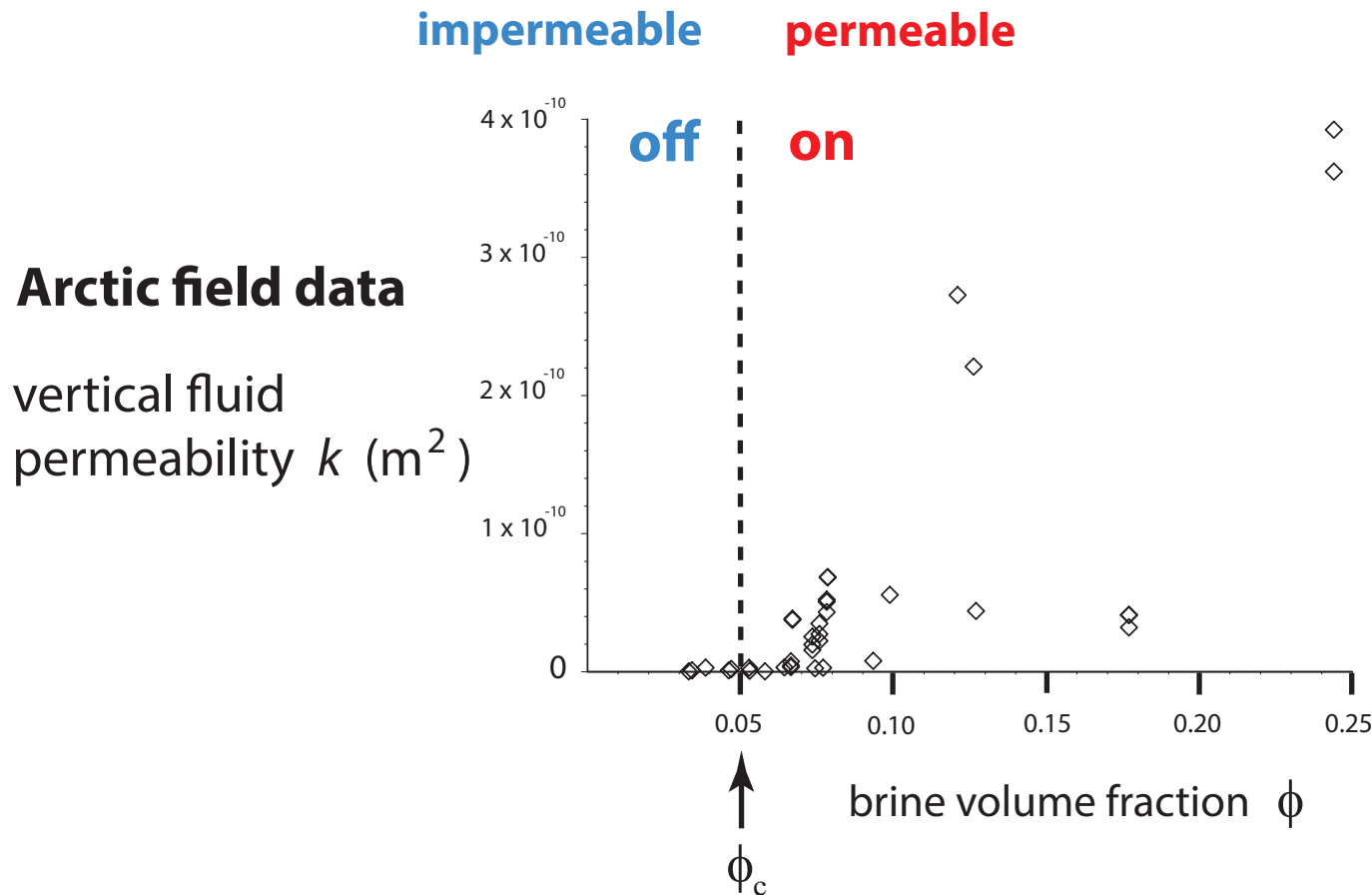
viscosity

$\mathbf{k}$  = fluid permeability tensor

## *HOMOGENIZATION*

*mathematics for analyzing effective behavior of heterogeneous systems*

# Critical behavior of fluid transport in sea ice



***“on - off” switch  
for fluid flow***

critical brine volume fraction  $\phi_c \approx 5\% \longleftrightarrow T_c \approx -5^\circ \text{C}, S \approx 5 \text{ ppt}$

**RULE OF FIVES**

**Golden, Ackley, Lytle Science 1998**

**Golden, Eicken, Heaton, Miner, Pringle, Zhu GRL 2007**

**Pringle, Miner, Eicken, Golden J. Geophys. Res. 2009**





# sea ice algal communities

D. Thomas 2004

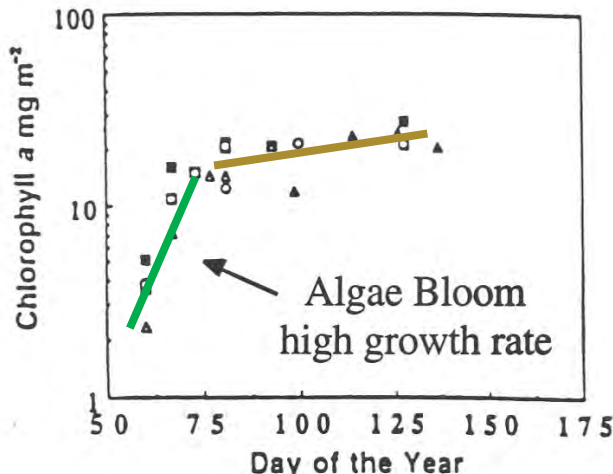
nutrient replenishment  
controlled by ice permeability

biological activity turns on  
or off according to  
**rule of fives**

**Golden, Ackley, Lytle**      **Science 1998**

**Fritsen, Lytle, Ackley, Sullivan**      **Science 1994**

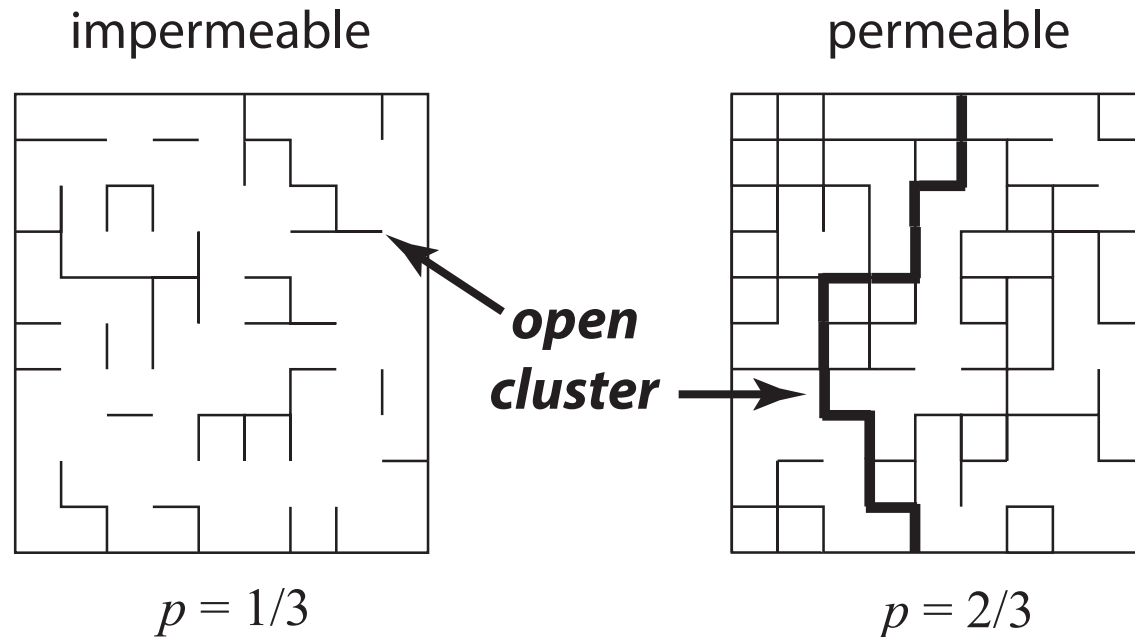
## critical behavior of microbial activity



Convection-fueled algae bloom  
Ice Station Weddell

# percolation theory

## *probabilistic theory of connectedness*



bond  $\longrightarrow$  *open* with probability  $p$   
*closed* with probability  $1-p$

## percolation threshold

$$p_c = 1/2 \quad \text{for } d = 2$$

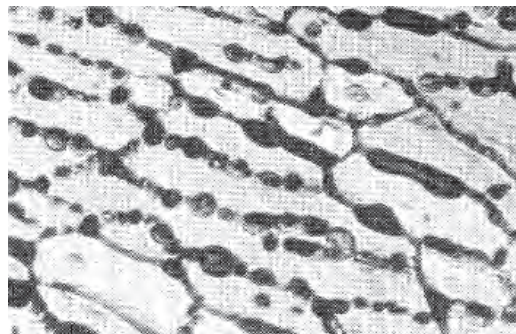
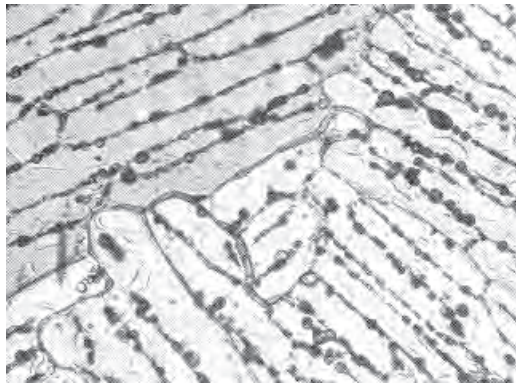
smallest  $p$  for which there is an infinite open cluster



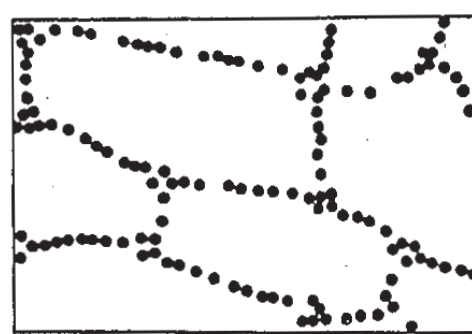
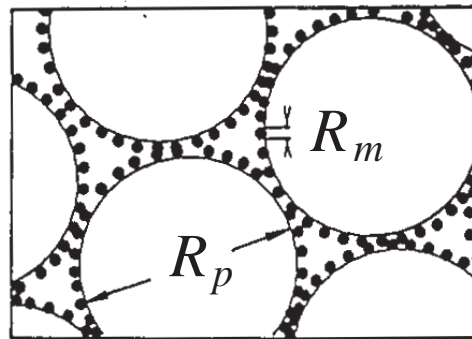
*Continuum* percolation model for **stealthy** materials applied to sea ice microstructure explains **Rule of Fives** and Antarctic data on **ice production** and **algal growth**

$$\phi_c \approx 5 \%$$

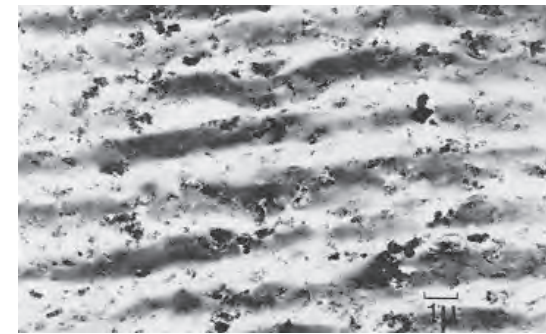
Golden, Ackley, Lytle, *Science*, 1998



sea ice



compressed  
powder

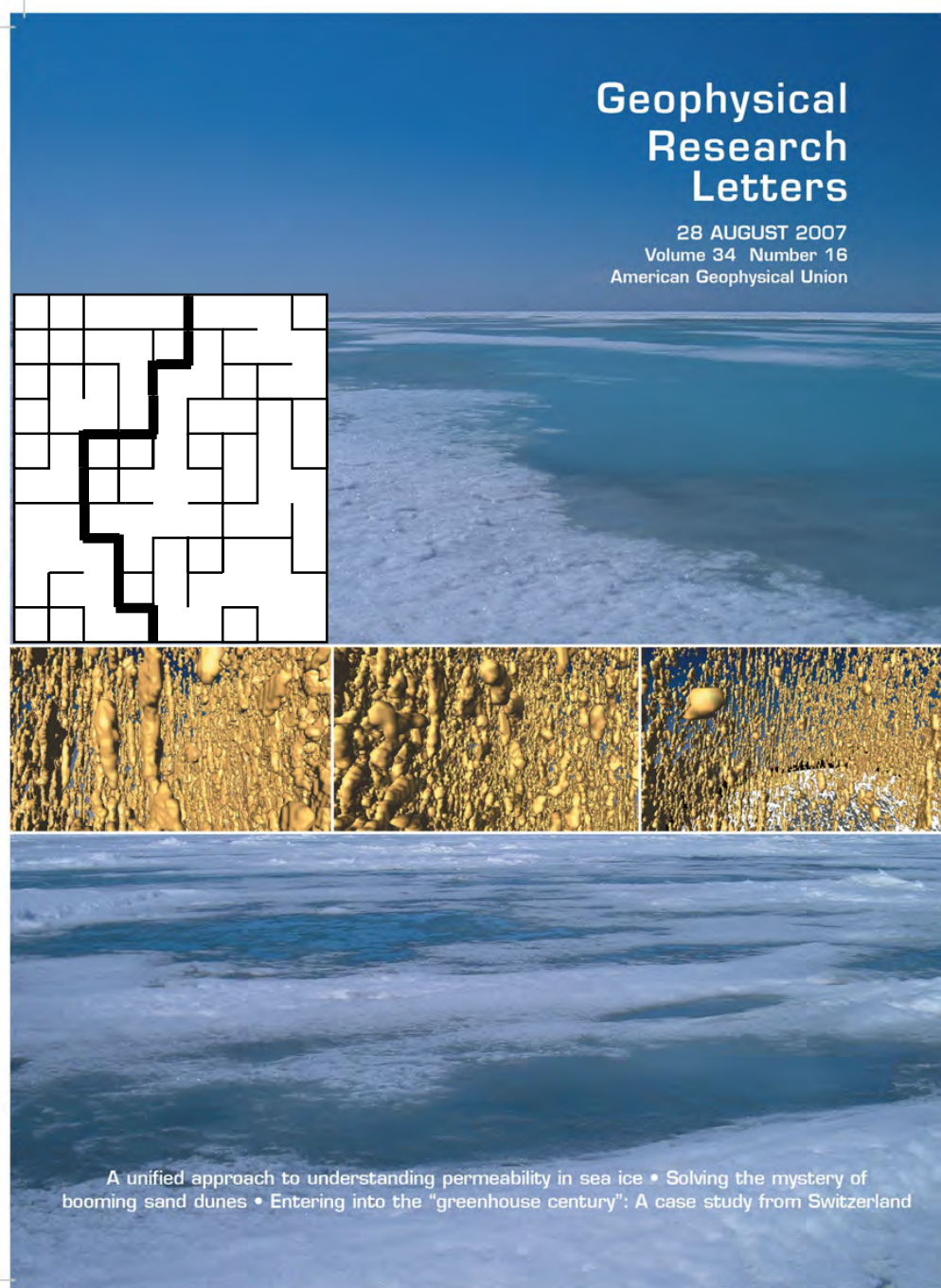


radar absorbing  
composite

**sea ice is radar absorbing**

# Thermal evolution of permeability and microstructure in sea ice

Golden, Eicken, Heaton\*, Miner, Pringle, Zhu, *Geophysical Research Letters* 2007



**percolation theory  
for fluid permeability**

$$k(\phi) = k_0 (\phi - 0.05)^2$$

critical  
exponent  
 $t$

$$k_0 = 3 \times 10^{-8} \text{ m}^2$$

from critical path analysis  
in **hopping conduction**

**hierarchical model**

**rock physics**

**network model**

**rigorous bounds**

**X-ray tomography for  
brine inclusions**

**confirms rule of fives**

*Pringle, Miner, Eicken, Golden  
J. Geophys. Res. 2009*

**theories agree closely  
with field data**

microscale  
governs

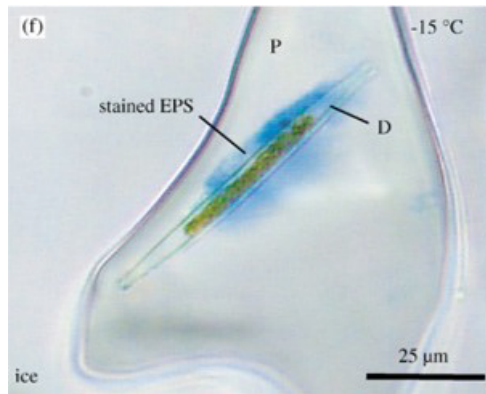
mesoscale  
processes

**melt pond  
evolution**

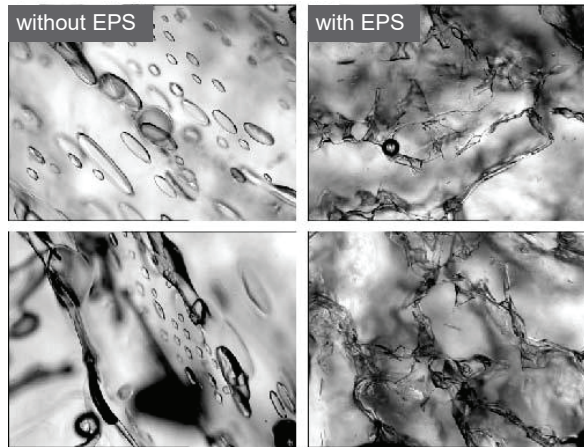


# Sea ice algae secrete extracellular polymeric substances (EPS) affecting evolution of brine microstructure.

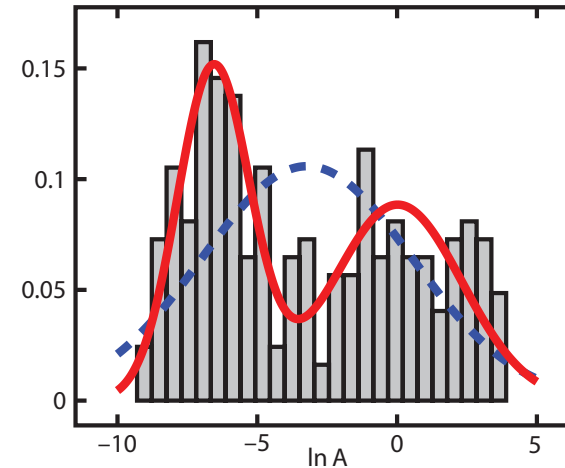
How does EPS affect fluid transport? How does the biology affect the physics?



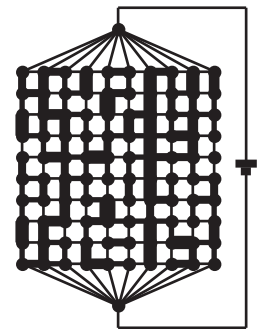
Krembs



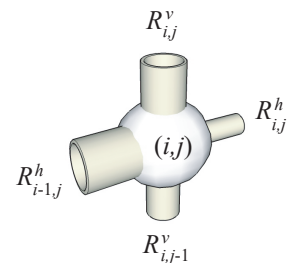
Krembs, Eicken, Deming, PNAS 2011



**RANDOM  
PIPE  
MODEL**



- 2D random pipe model with bimodal distribution of pipe radii
- Rigorous bound on permeability  $k$ ; results predict observed drop in  $k$



Steffen, Epshteyn, Zhu, Bowler, Deming, Golden  
*Multiscale Modeling and Simulation*, 2018

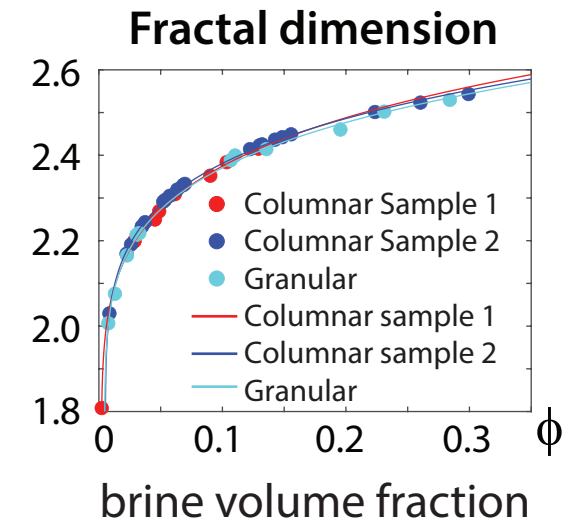
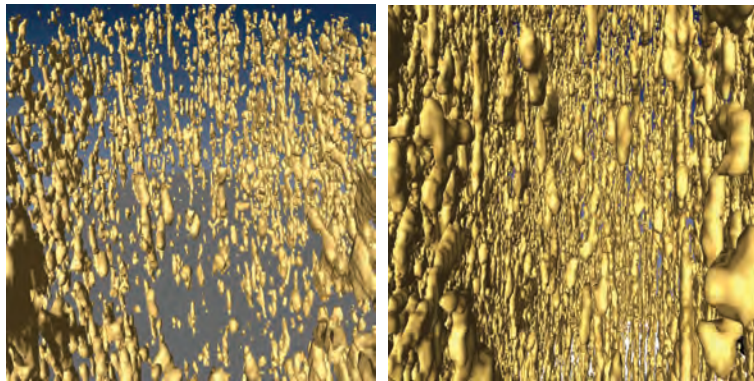
Zhu, Jabini, Golden,  
Eicken, Morris  
*Ann. Glac.* 2006



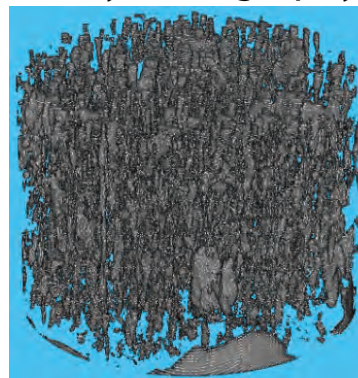
# Thermal evolution of the fractal geometry of the brine microstructure in sea ice

N. Ward, D. Hallman, H. Eicken, M. Oggier and K. M. Golden, 2022

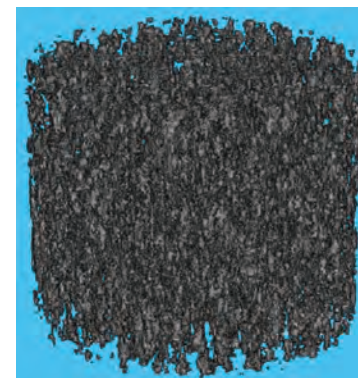
$T = -12\text{ }^{\circ}\text{C}$ ,  $\phi = 0.033$        $T = -8\text{ }^{\circ}\text{C}$ ,  $\phi = 0.057$



X-ray tomography



DLA model



# Arctic and Antarctic field experiments

*develop electromagnetic methods  
of monitoring fluid transport and  
microstructural transitions*

extensive measurements of fluid and  
electrical transport properties of sea ice:

**2007    Antarctic    SIPEX**

**2010    Antarctic    McMurdo Sound**

**2011    Arctic        Barrow AK**

**2012    Arctic        Barrow AK**

**2012    Antarctic    SIPEX II**

**2013    Arctic        Barrow AK**

**2014    Arctic        Chukchi Sea**





# Notices

of the American Mathematical Society

May 2009

Volume 56, Number 5

Climate Change and  
the Mathematics of  
Transport in Sea Ice

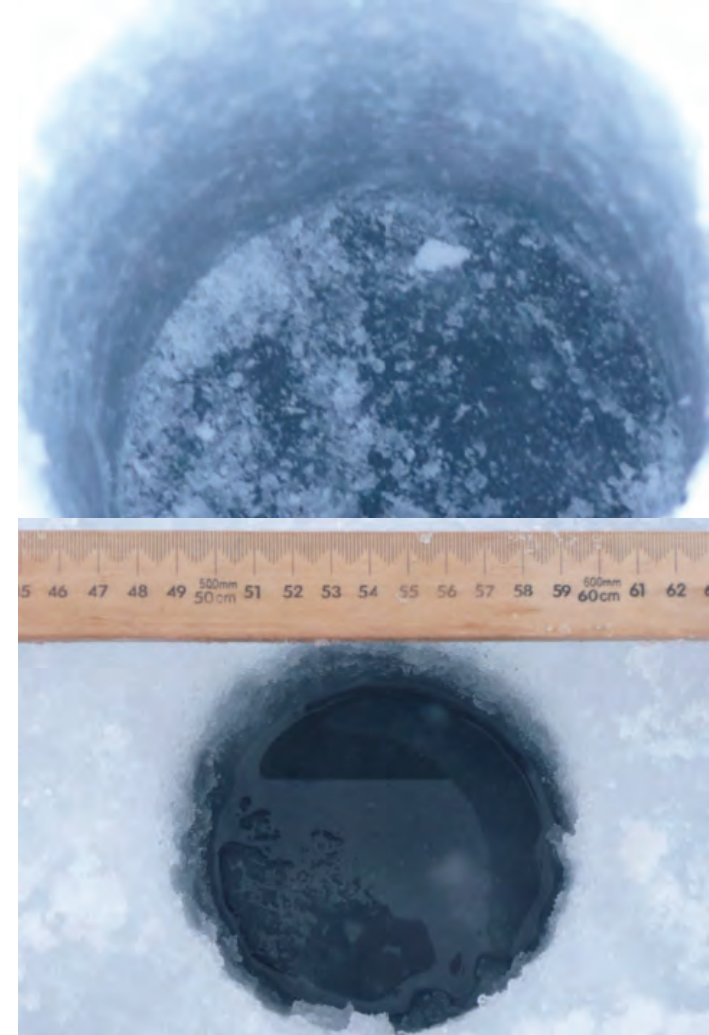
page 562

Mathematics and the  
Internet: A Source of  
Enormous Confusion  
and Great Potential

page 586

photo by Jan Lieser

Real analysis in polar coordinates (see page 613)



***measuring  
fluid permeability  
of Antarctic sea ice***

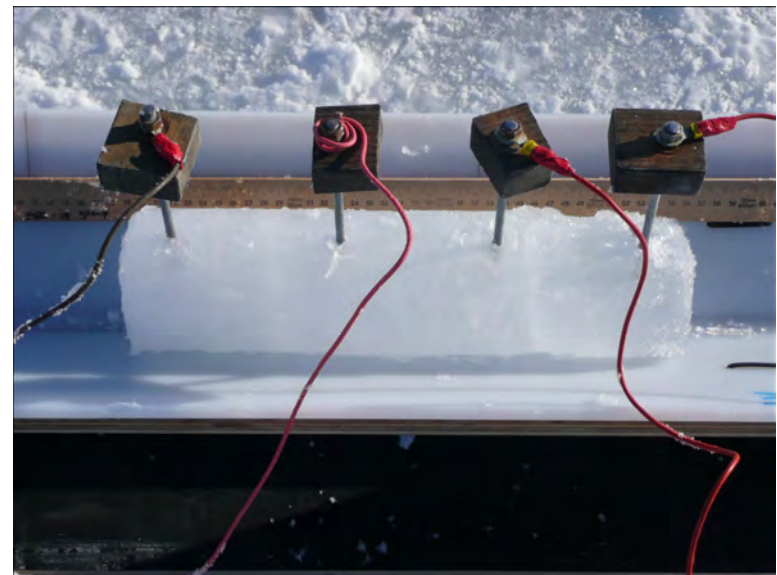
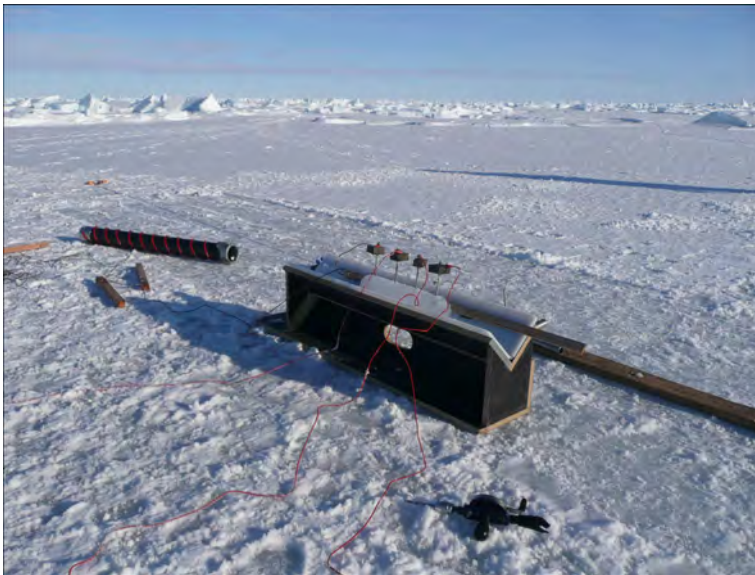
***SIPEX 2007***



## electrical measurements



## Wenner array



## vertical conductivity

Zhu, Golden, Gully, Sampson *Physica B* 2010

Sampson, Golden, Gully, Worby *Deep Sea Research* 2011

# ***cross borehole tomography***



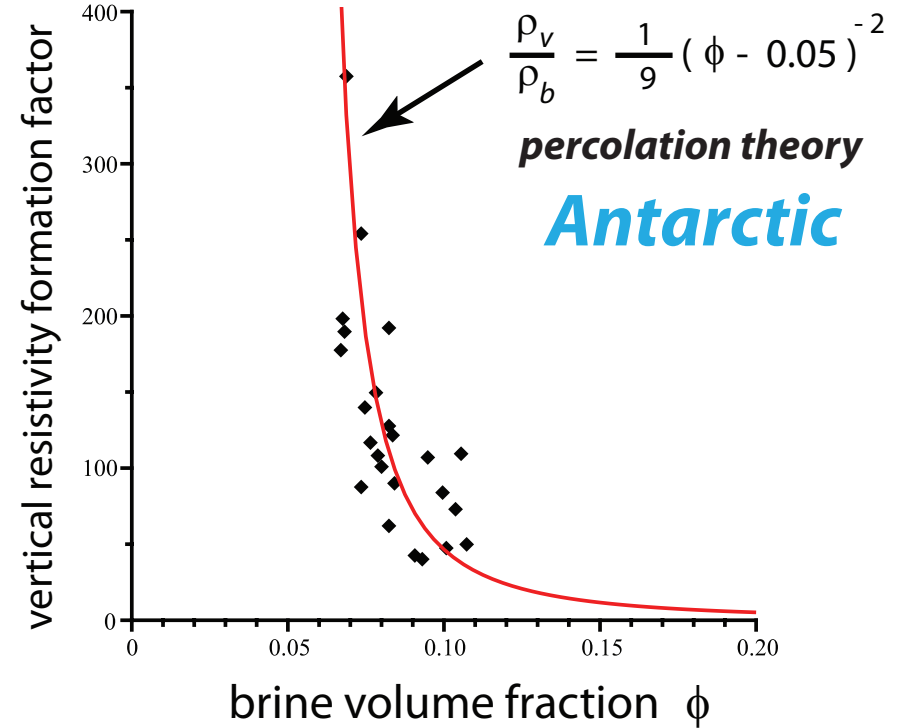
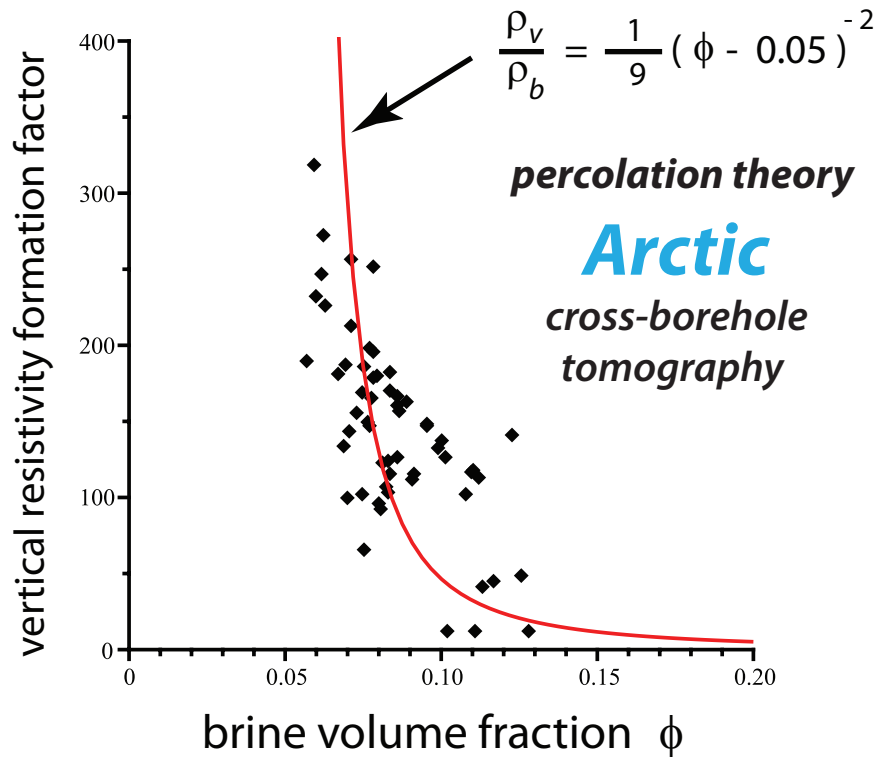
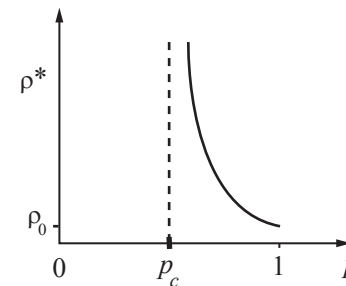
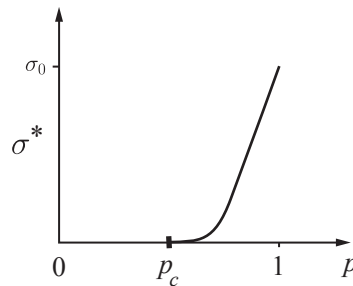
***Ingham, Jones, Buchanan  
Victoria University, Wellington, NZ***

# critical behavior of electrical transport in sea ice

## electrical signature of the on-off switch for fluid flow

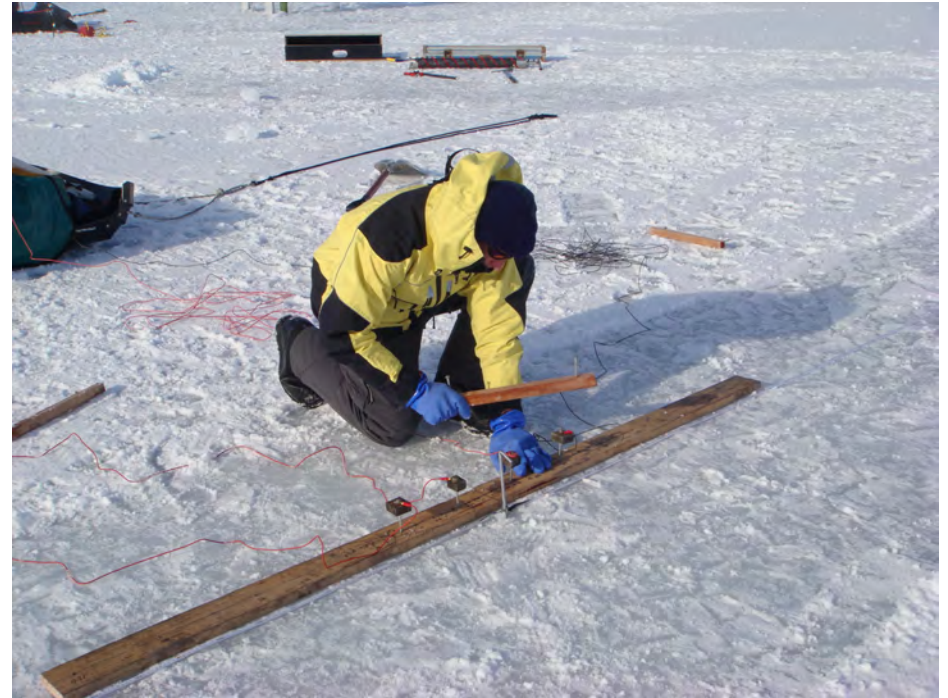
same universal critical exponent as for fluid permeability

studied for over 50 years but no previous observations or theory of critical behavior





# Measuring sea ice thickness







# Modeling of anisotropic electromagnetic reflection from sea ice

## Golden & Ackley, JGR 1981

measure sea ice thickness using 100 MHz waves

Model sea ice as a two phase composite of pure ice with anisotropic ellipsoidal brine inclusions; use mean field theory to estimate complex permittivity

Explained marked anisotropy in bottom return in terms of relation between E field polarization and current-induced anisotropy in brine microstructure (c-axis direction).

Weeks & Gow, 1979  
Kovacs & Morey, 1978

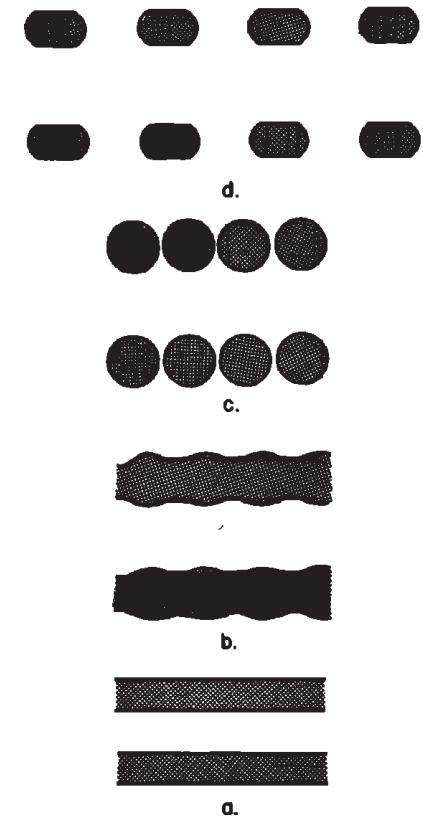


Fig. 4. Brine layers (top view) near the bottom of sea ice (a) begin to 'neck' with decreasing temperature further up in the ice sheet(s) and freeze out into cylinders (c) and ellipitical cylinders (d) [from *Anderson and Weeks*, 1958].

Bounds on the complex permittivity of polycrystalline sea ice  
with anisotropy in the horizontal plane  
K. McLean, E. Cherkaev, K. M. Golden, 2022



# Remote sensing of sea ice



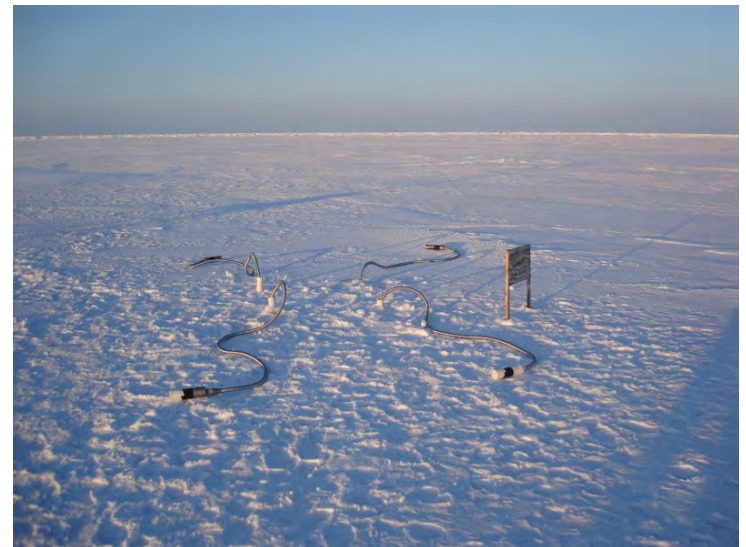
*sea ice thickness*  
*ice concentration*

## **INVERSE PROBLEM**

Recover sea ice  
properties from  
electromagnetic  
(EM) data

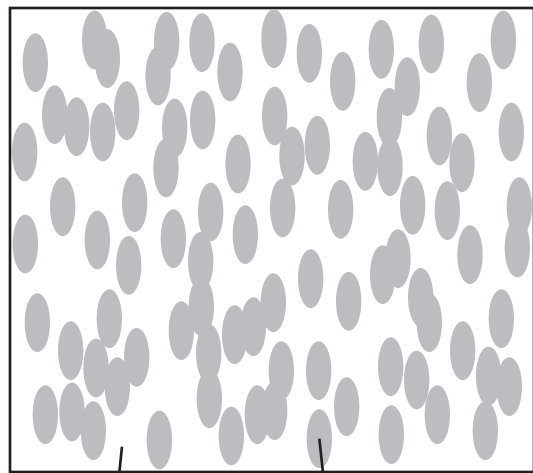
$$\epsilon^*$$

effective complex permittivity  
(dielectric constant, conductivity)



*brine volume fraction*  
*brine inclusion connectivity*

# Effective complex permittivity of a two phase composite in the quasistatic (long wavelength) limit



$\epsilon_1$

$\epsilon_2$



$\epsilon^*$

$$D = \epsilon E$$

$$\nabla \cdot D = 0$$

$$\nabla \times E = 0$$

$$\langle D \rangle = \epsilon^* \langle E \rangle$$

$p_1, p_2$  = volume fractions of  
the components

$$\epsilon^* = \epsilon^* \left( \frac{\epsilon_1}{\epsilon_2}, \text{ composite geometry} \right)$$

**What are the effective propagation characteristics  
of an EM wave (radar, microwaves) in the medium?**

# Analytic Continuation Method for Homogenization

Bergman (1978), Milton (1979), Golden and Papanicolaou (1983), Theory of Composites, Milton (2002)

## Stieltjes integral representation for homogenized parameter

*separates geometry from parameters*

$$F(s) = 1 - \frac{\epsilon^*}{\epsilon_2} = \int_0^1 \frac{d\mu(z)}{s - z}$$

← geometry

← material parameters

$$s = \frac{1}{1 - \epsilon_1 / \epsilon_2}$$

$\mu$

- spectral measure of self adjoint operator  $\Gamma\chi$
- mass =  $p_1$
- higher moments depend on  $n$ -point correlations

$$\Gamma = \nabla(-\Delta)^{-1}\nabla.$$

$\chi$  = characteristic function of the brine phase

$$E = s (s + \Gamma\chi)^{-1} e_k$$

$\Gamma\chi$  : microscale  $\rightarrow$  macroscale

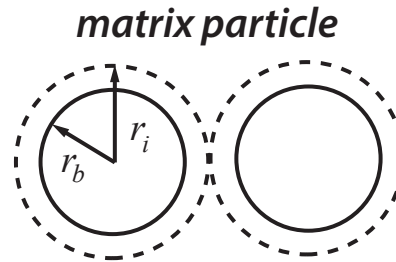
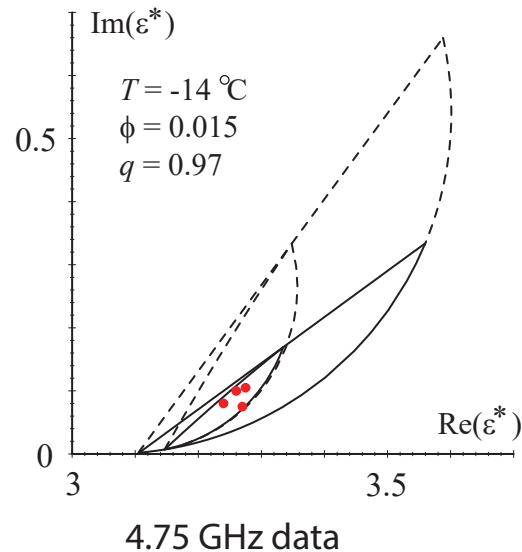
$\Gamma\chi$  *links scales*



**This representation distills the complexities of mixture geometry into the spectral properties of an operator like the Hamiltonian in physics.**

# forward and inverse bounds on the complex permittivity of sea ice

## forward bounds

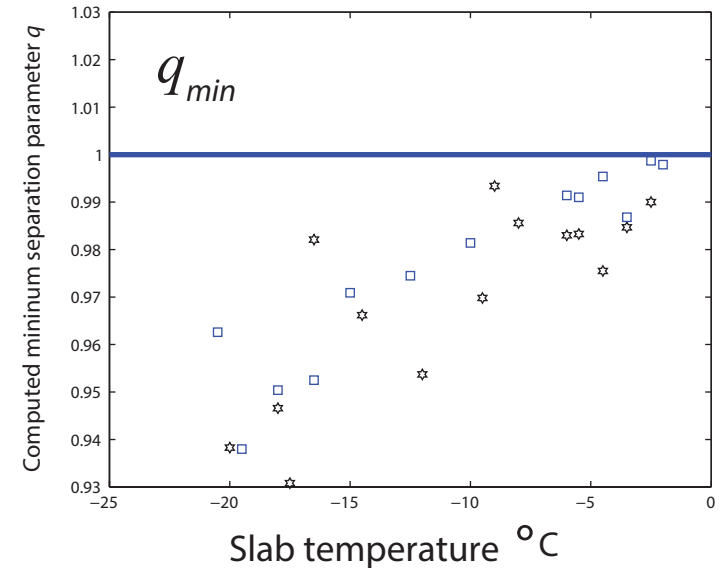


$$q = r_b / r_i$$

$$0 < q < 1$$

**Golden 1995, 1997**

## inverse bounds



## Inverse Homogenization

Cherkaev and Golden (1998), Day and Thorpe (1999), Cherkaev (2001), McPhedran, McKenzie, Milton (1982), *Theory of Composites*, Milton (2002)

$\epsilon^*$   $\longrightarrow$  composite geometry  
(spectral measure  $\mu$ )

## inverse bounds and recovery of brine porosity

**Gully, Backstrom, Eicken, Golden**  
*Physica B*, 2007

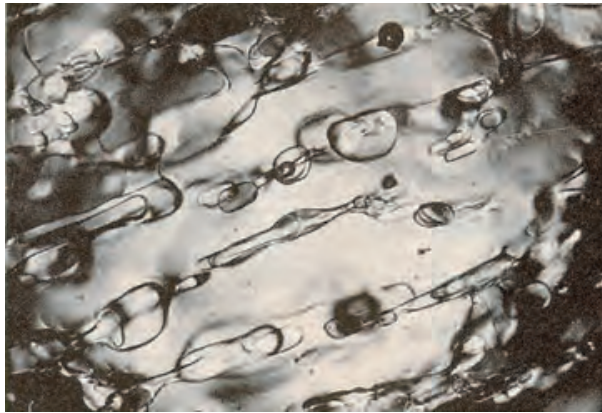
## inversion for brine inclusion separations in sea ice from measurements of effective complex permittivity $\epsilon^*$

### rigorous inverse bound on spectral gap

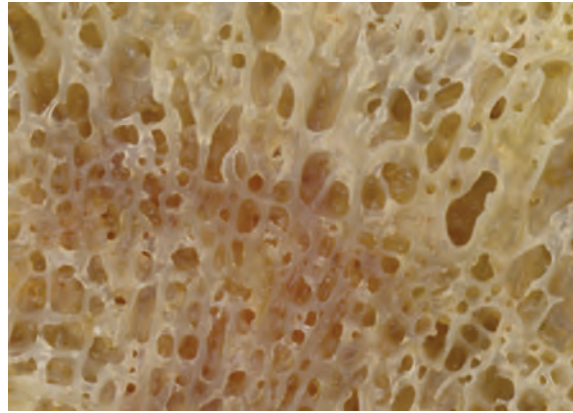
construct algebraic curves which bound admissible region in  $(p, q)$ -space

**Orum, Cherkaev, Golden**  
*Proc. Roy. Soc. A*, 2012

## SEA ICE

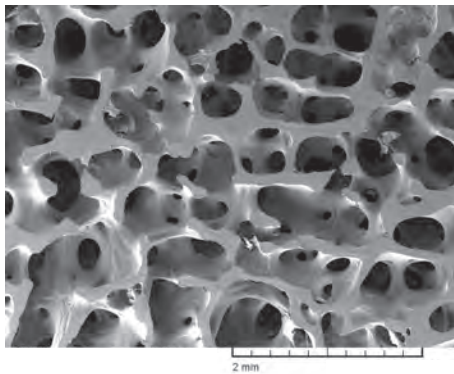


## HUMAN BONE

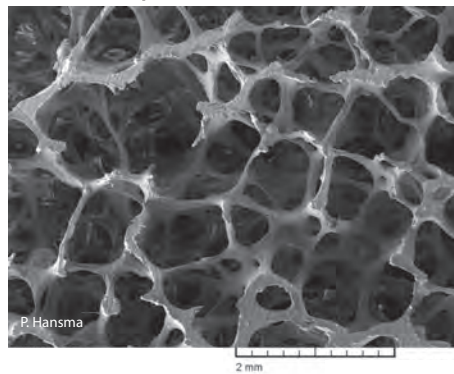


*spectral characterization  
of porous microstructures  
in human bone*

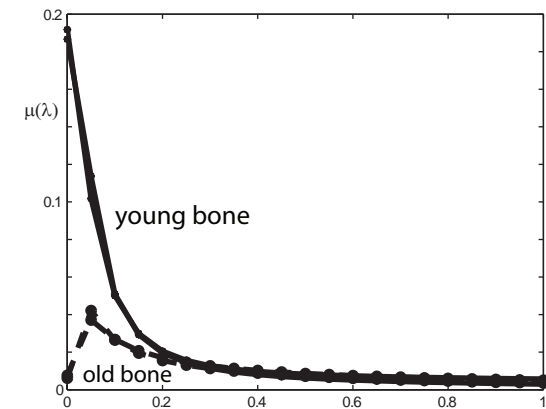
young healthy trabecular bone



old osteoporotic trabecular bone



reconstruct spectral measures  
from complex permittivity data



use regularized inversion scheme

*apply spectral measure analysis of brine connectivity and  
spectral inversion to electromagnetic monitoring of osteoporosis*

Golden, Murphy, Cherkaev, J. Biomechanics 2011

***the math doesn't care if it's sea ice or bone!***



# direct calculation of spectral measures

Murphy, Hohenegger, Cherkaev, Golden, *Comm. Math. Sci.* 2015

- depends only on the composite geometry
- discretization of microstructural image gives binary network
- fundamental operator becomes a random matrix
- spectral measure computed from eigenvalues and eigenvectors

**once we have the spectral measure  $\mu$  it can be used in  
Stieltjes integrals for other transport coefficients:**

***electrical and thermal conductivity, complex permittivity,  
magnetic permeability, diffusion, fluid flow properties***

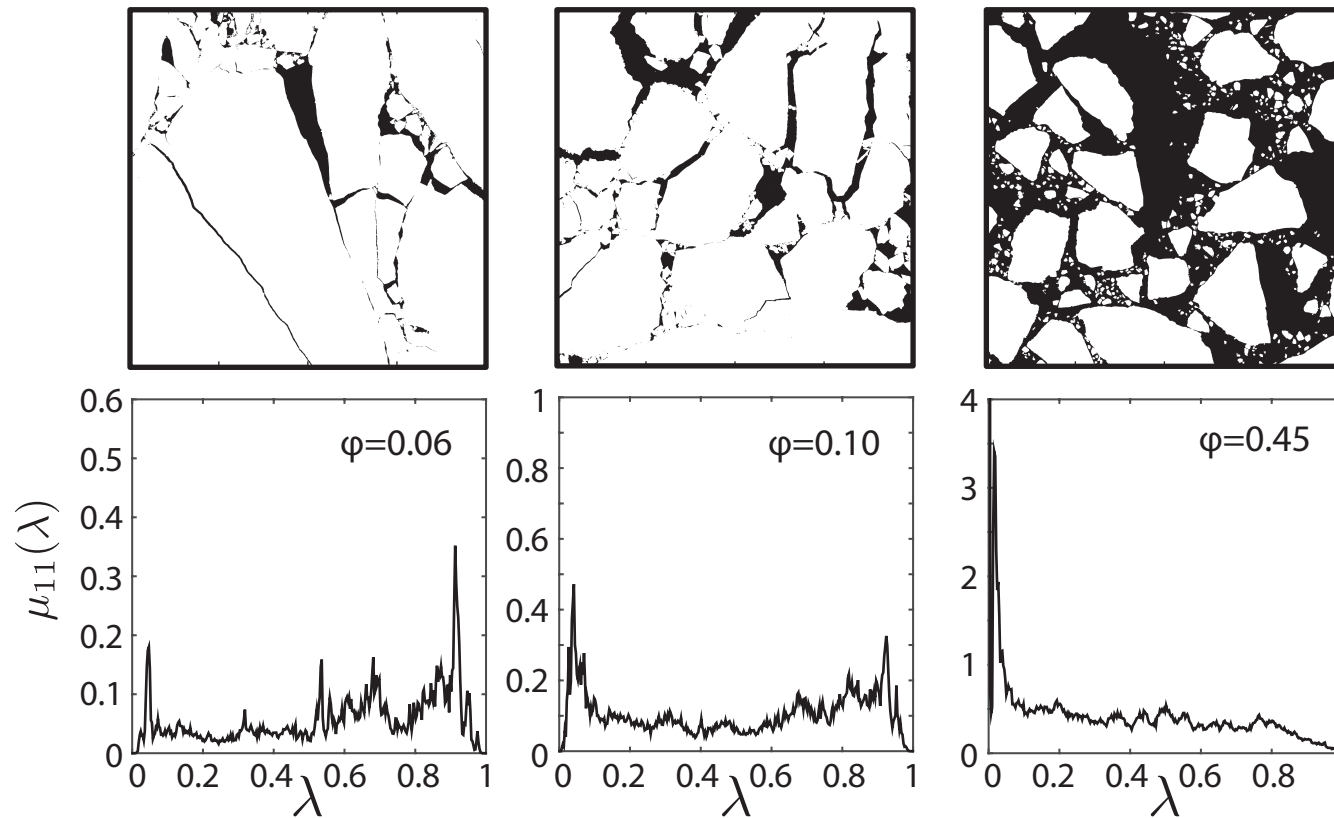
earlier studies of spectral measures

Day and Thorpe 1996

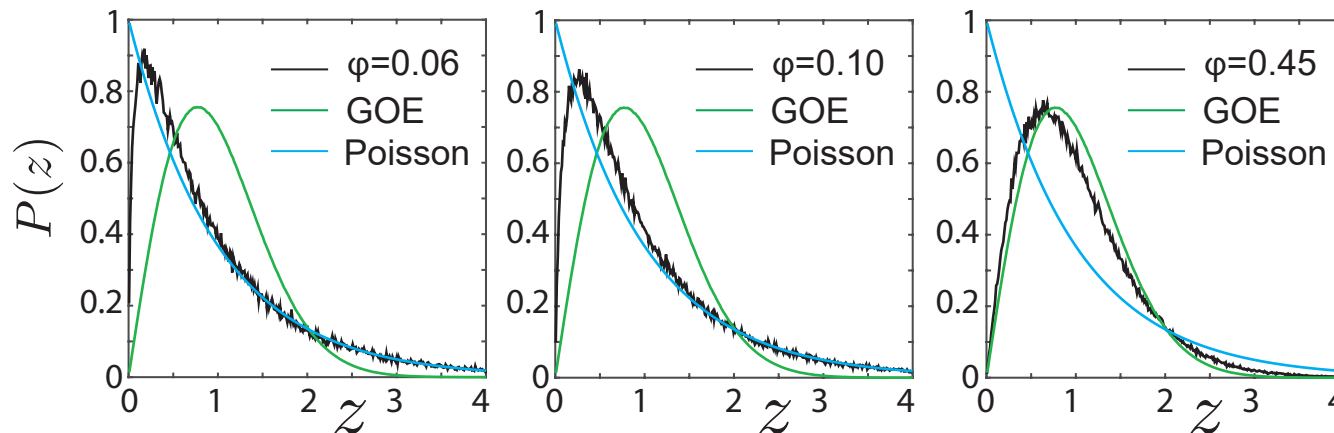
Helsing, McPhedran, Milton 2011

# Spectral computations for sea ice floe configurations

spectral  
measures



eigenvalue  
spacing  
distributions



uncorrelated



level repulsion

**UNIVERSAL**  
**Wigner-Dyson**  
**distribution**

# Eigenvalue Statistics of Random Matrix Theory

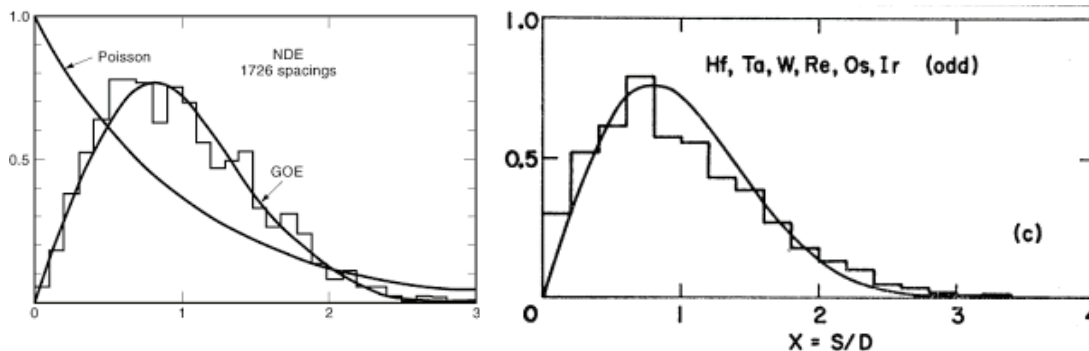
*Wigner (1951) and Dyson (1953) first used random matrix theory (RMT) to describe quantized energy levels of heavy atomic nuclei.*

$[N]_{ij} \sim N(0,1), \quad A = (N + N^T)/2 \quad \text{Gaussian orthogonal ensemble (GOE)}$

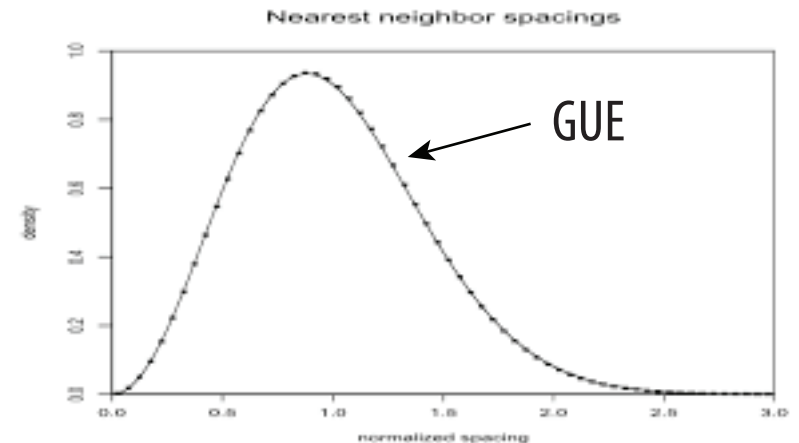
$[N]_{ij} \sim N(0,1) + iN(0,1), \quad A = (N + N^\dagger)/2 \quad \text{Gaussian unitary ensemble (GUE)}$

*Short range and long range correlations of eigenvalues are measured by various eigenvalue statistics.*

Spacing distributions of energy levels for heavy atomic nuclei

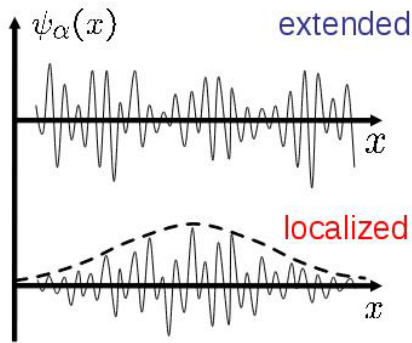


Spacing distributions of the first billion zeros of the Riemann zeta function



**Universal eigenvalue statistics arise in a broad range of “unrelated” problems!**





electronic transport in semiconductors

metal / insulator transition

**localization**

*Anderson 1958*  
*Mott 1949*  
*Shklovshii et al 1993*  
*Evangelou 1992*

**Anderson transition in wave physics:  
 quantum, optics, acoustics, water waves, ...**

**from analysis of spectral measures for brine, melt ponds, ice floes**

we find percolation-driven

***Anderson transition for classical transport in composites***

*Murphy, Cherkaev, Golden Phys. Rev. Lett. 2017*

**PERCOLATION  
 TRANSITION**



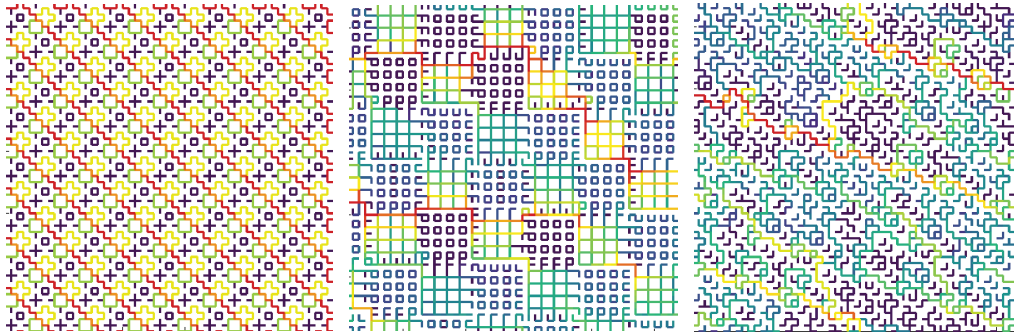
**universal eigenvalue statistics (GOE)  
 extended states, mobility edges**

**-- but with NO wave interference or scattering effects ! --**

# Order to disorder in quasiperiodic composites

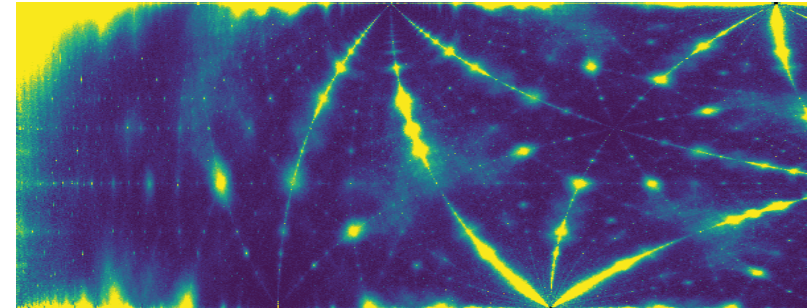
Morison, Murphy, Cherkaev, Golden, *Commun. Phys.* 2022

Parameterized Moiré Pattern Creates Tunable Microgeometry



Anderson transition as QP is tuned

Poisson  
Wigner-Dyson



parameter space

periodic

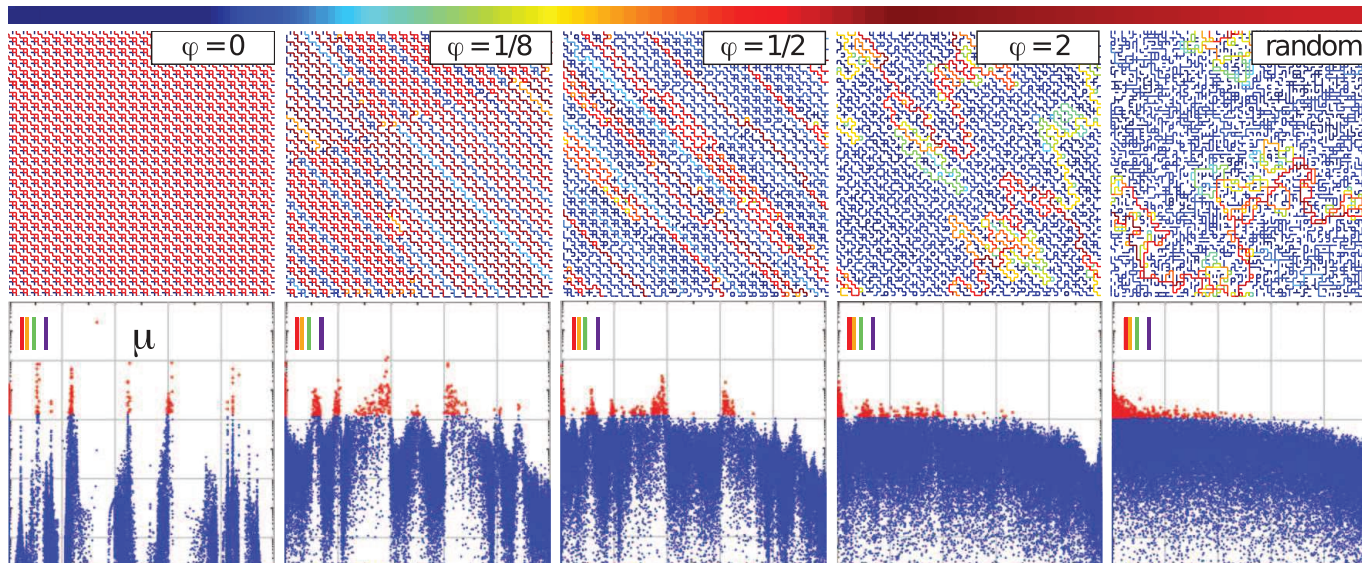


quasiperiodic

electric field  
strength

spectral  
measure

$10^{-4}$   
 $10^{-6}$   
 $10^{-8}$



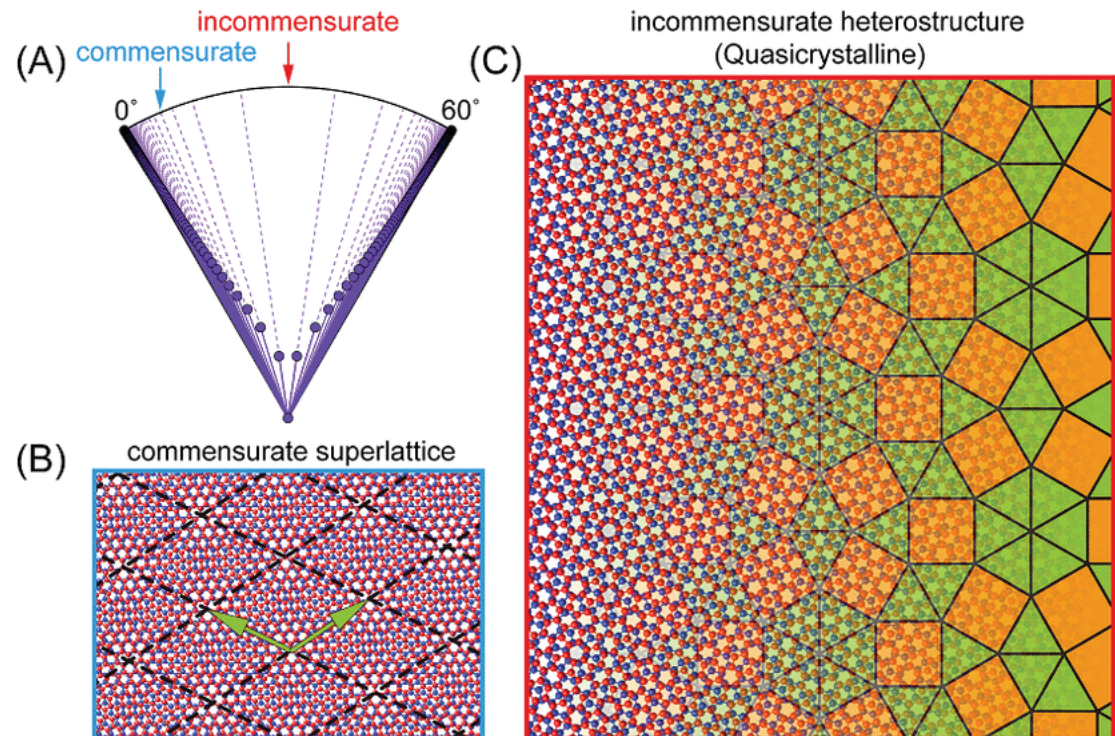
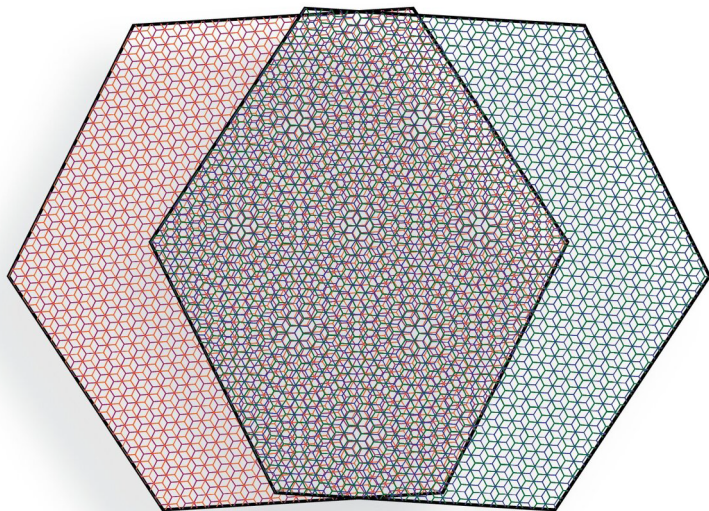
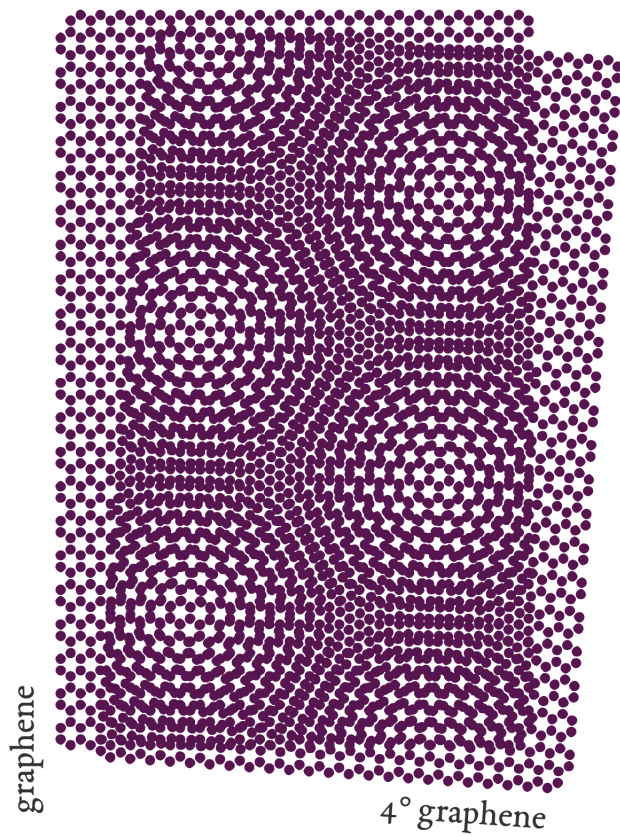
RRN at  
percolation  
threshold

we bring the framework of solid state physics of electronic transport and band gaps in semiconductors to classical transport in periodic and quasiperiodic composites

photonic crystals and quasicrystals

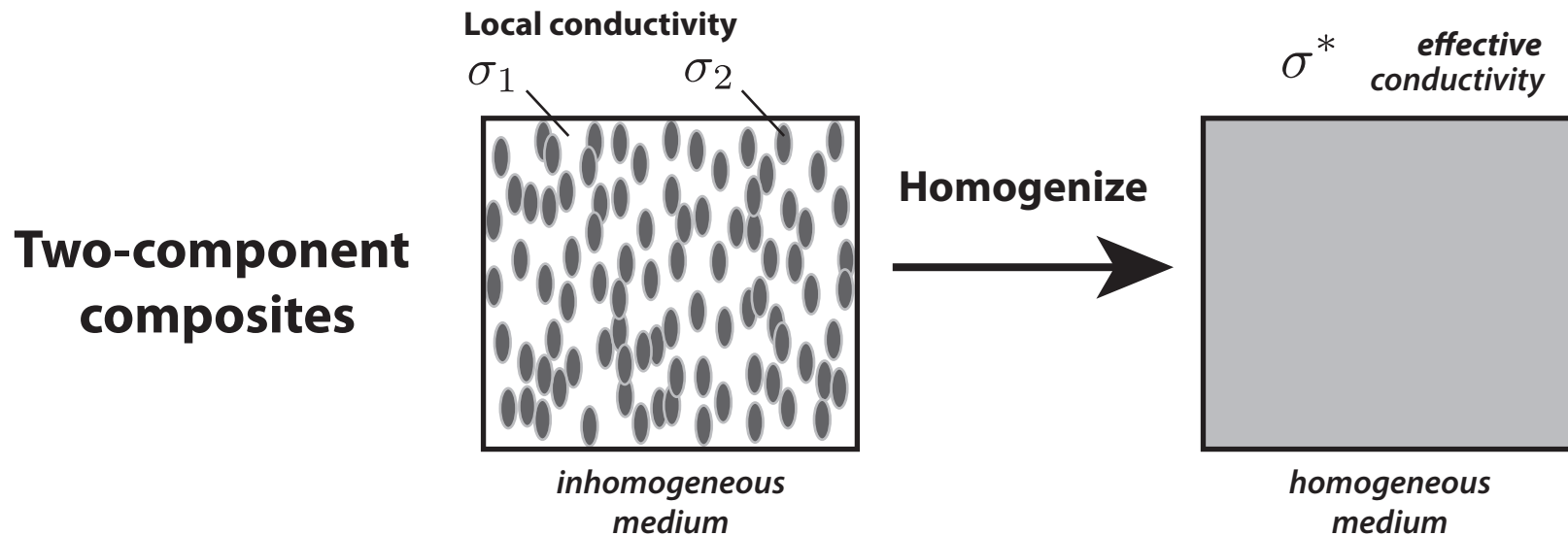


# twisted bilayer graphene

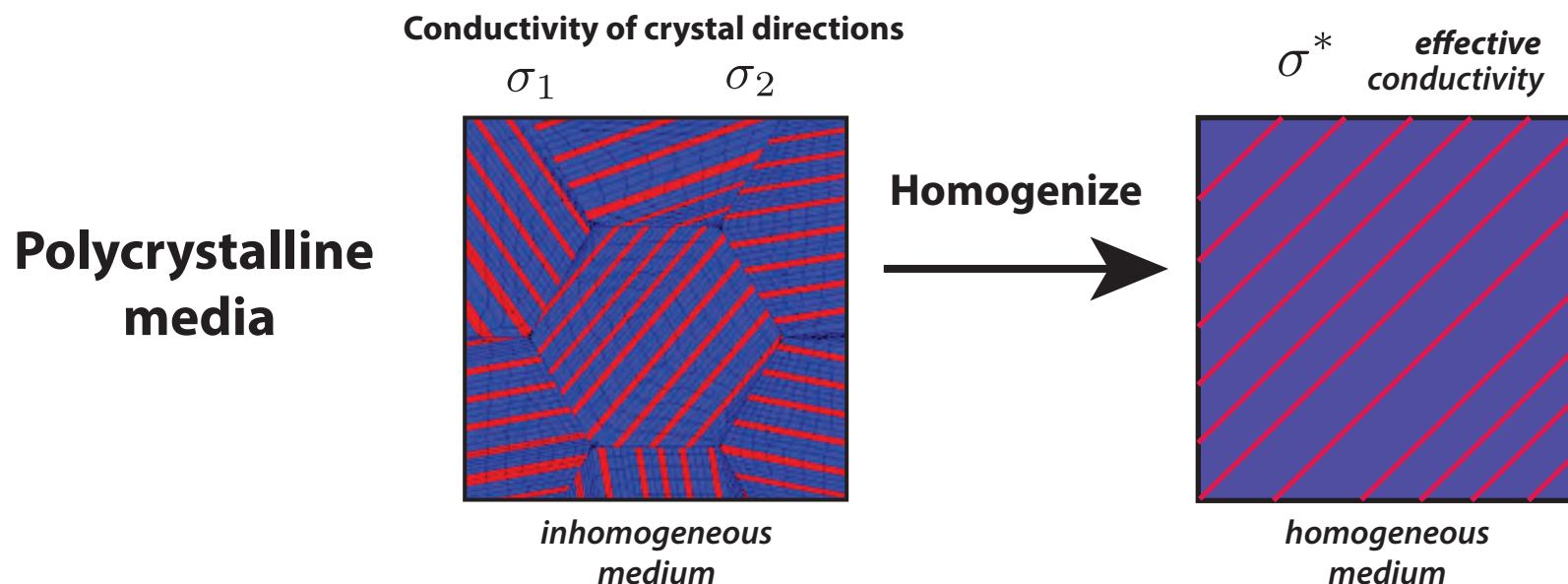




# Homogenization for polycrystalline materials



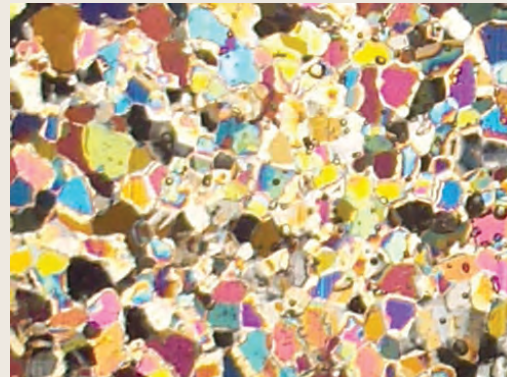
**Find the homogeneous medium which behaves macroscopically the same as the inhomogeneous medium**



# Bounds on the complex permittivity of polycrystalline materials by analytic continuation

Adam Gully, Joyce Lin,  
Elena Cherkaev, Ken Golden

- **Stieltjes integral representation for effective complex permittivity**  
Milton (1981, 2002), Barabash and Stroud (1999), ...
- **Forward and inverse bounds**  
*orientation statistics*
- **Applied to sea ice using two-scale homogenization**
- **Inverse bounds give method for distinguishing ice types using remote sensing techniques**



## PROCEEDINGS A

350 YEARS  
OF SCIENTIFIC  
PUBLISHING

An invited review  
commemorating 350 years  
of scientific publishing at the  
Royal Society

A method to distinguish  
between different types  
of sea ice using remote  
sensing techniques

A computer model to  
determine how a human  
should walk so as to expend  
the least energy



THE  
ROYAL  
SOCIETY  
PUBLISHING



# higher threshold for fluid flow in granular sea ice

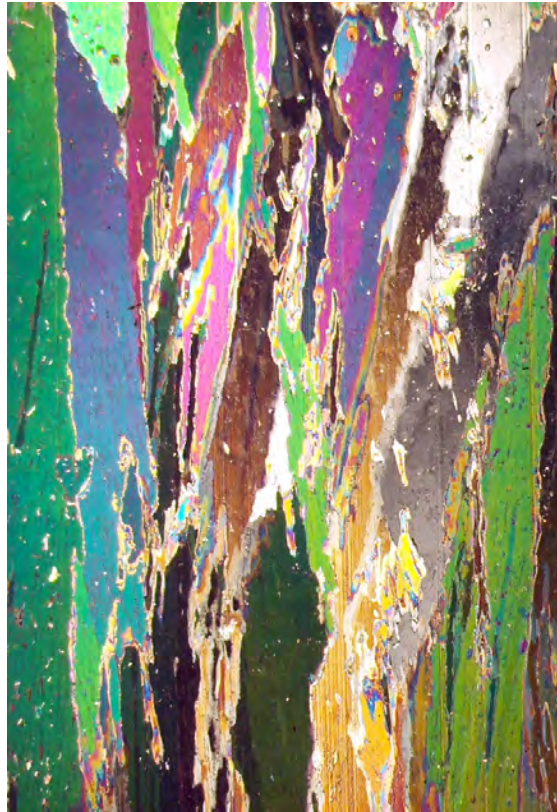
*microscale details impact “mesoscale” processes*

nutrient fluxes for microbes  
melt pond drainage  
snow-ice formation

columnar

granular

**5%**



**10%**



Tison

Golden, Sampson, Gully, Lubbers, Tison 2022

**electromagnetically distinguishing ice types**  
Kitsel Lusted, Elena Cherkaev, Ken Golden



**mesoscale**

# wave propagation in the marginal ice zone (MIZ)

Stieltjes integral representation and bounds for the complex viscoelasticity of the ice - ocean layer

Sampson, Murphy, Cherkaev, Golden 2022

first theory of key parameter in wave-ice interactions only fitted to wave data before

Keller, 1998

Mosig, Montiel, Squire, 2015

Wang, Shen, 2012

**Analytic Continuation Method**

Bergman (78) - Milton (79)  
integral representation for  $\epsilon^*$

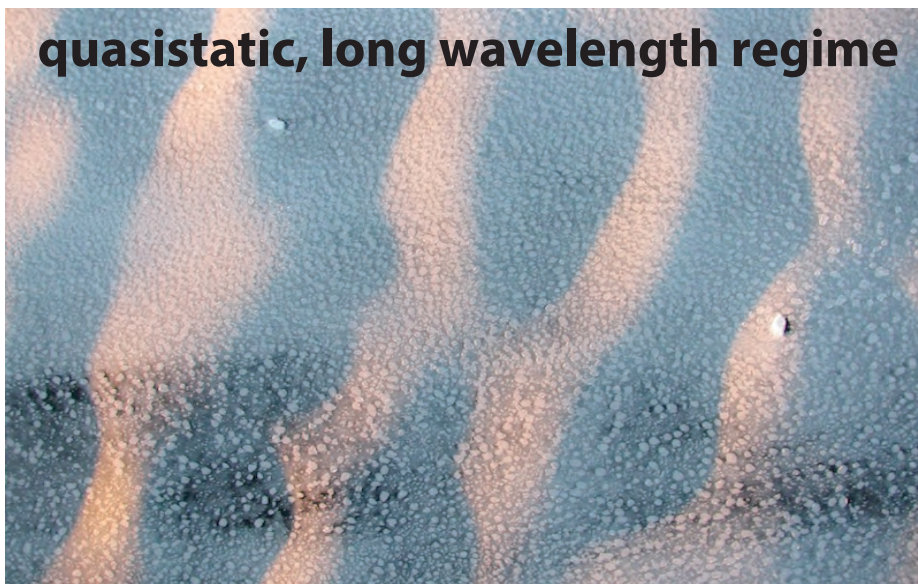
Golden and Papanicolaou (83)

Milton, *Theory of Composites* (02)

**quasistatic, long wavelength regime**

homogenized parameter depends on sea ice concentration and ice floe geometry

like EM waves





# advection enhanced diffusion

## effective diffusivity

nutrient and salt transport in sea ice  
heat transport in sea ice with convection  
sea ice floes in winds and ocean currents  
tracers, buoys diffusing in ocean eddies  
diffusion of pollutants in atmosphere

advection diffusion equation with a velocity field  $\vec{u}$

$$\frac{\partial T}{\partial t} + \vec{u} \cdot \vec{\nabla} T = \kappa_0 \Delta T$$

$$\vec{\nabla} \cdot \vec{u} = 0$$



homogenize

$$\frac{\partial \bar{T}}{\partial t} = \kappa^* \Delta \bar{T}$$

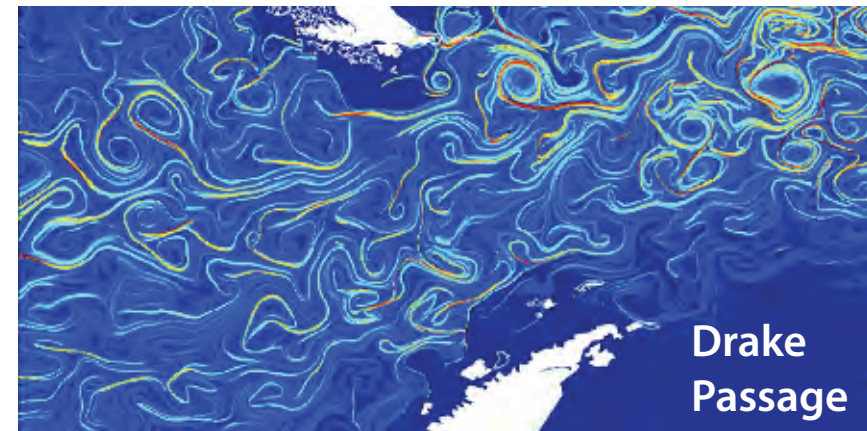
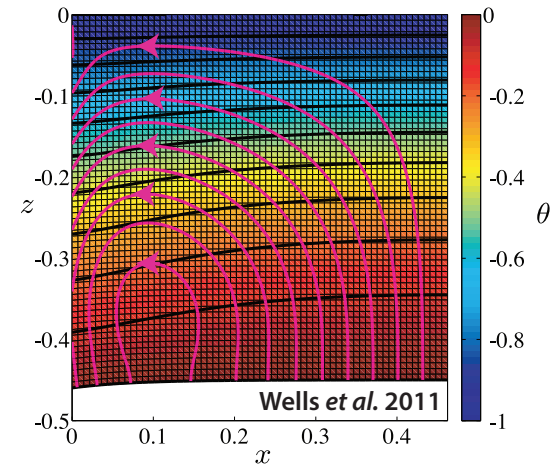
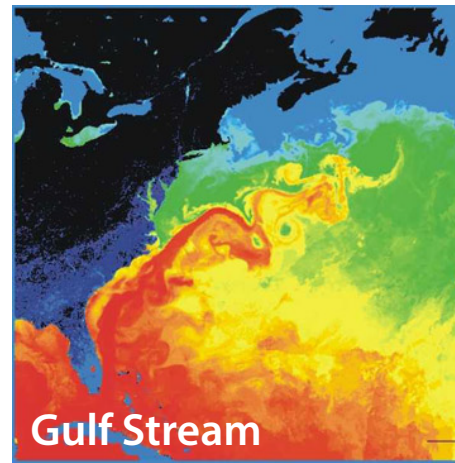
$\kappa^*$  effective diffusivity

**Stieltjes integral for  $\kappa^*$  with spectral measure**

*Avellaneda and Majda, PRL 89, CMP 91*

Murphy, Cherkaev, Xin, Zhu, Golden, *Ann. Math. Sci. Appl.* 2017

Murphy, Cherkaev, Zhu, Xin, Golden, *J. Math. Phys.* 2020



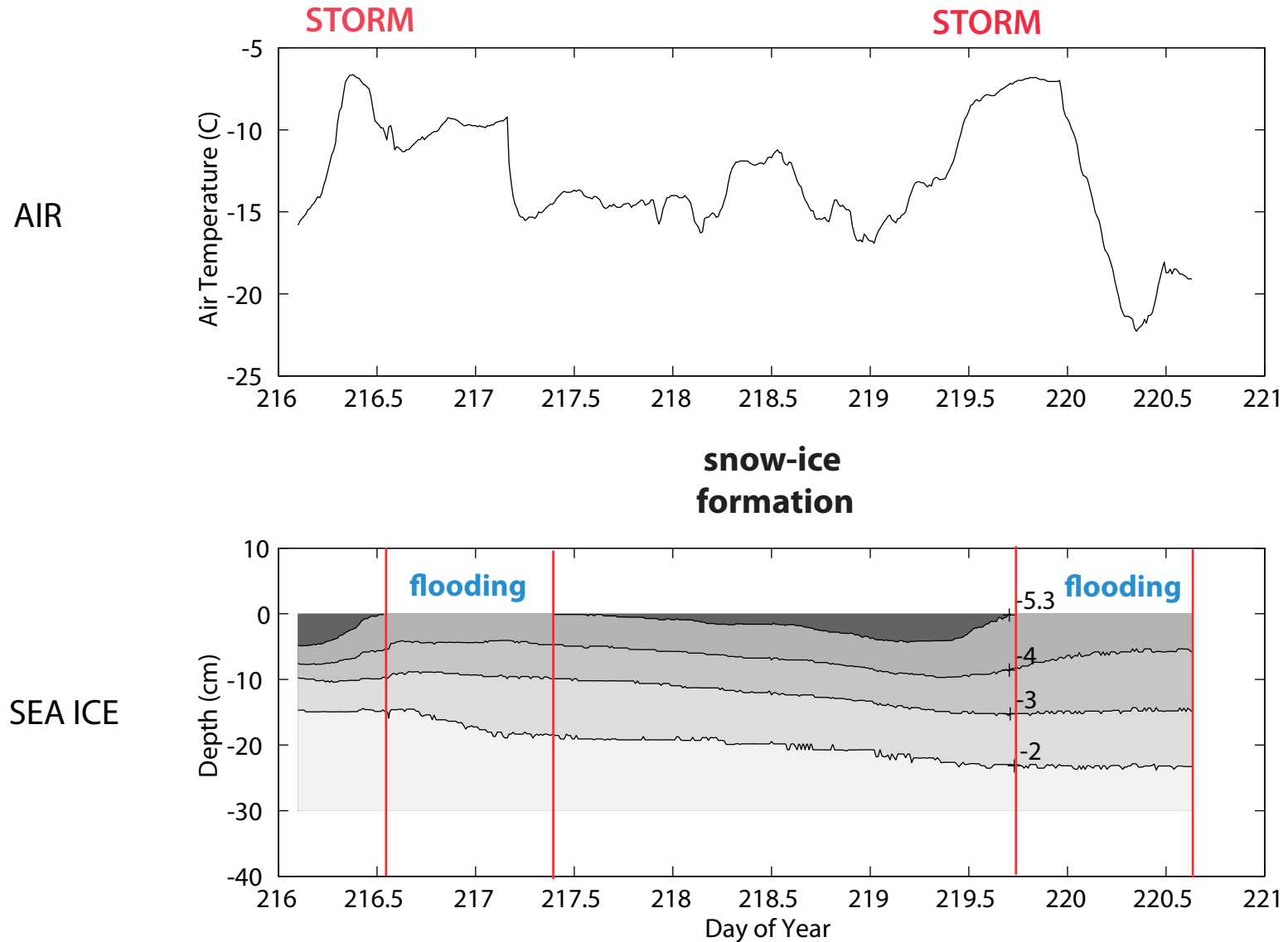


# tracers flowing through inverted sea ice blocks



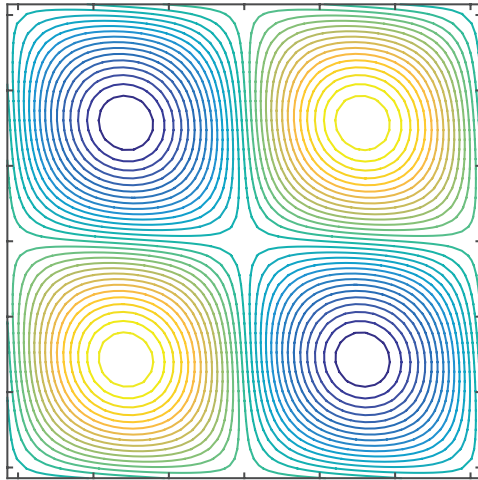
# *ANZFLUX drift camp*

**snow loading, surface flooding and subsequent snow - ice formation**



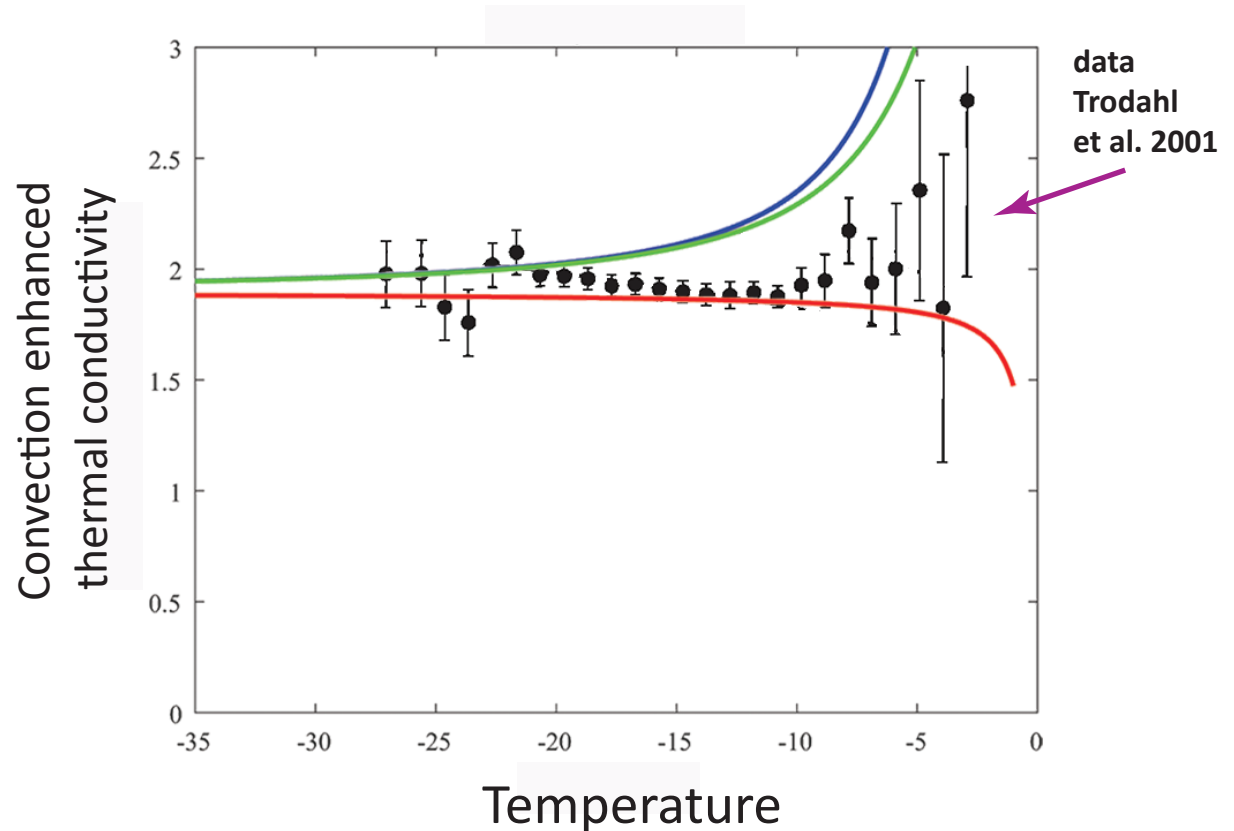
# Rigorous bounds on convection enhanced thermal conductivity of sea ice

Kraitzman, Hardenbrook, Dinh, Murphy, Zhu, Cherkaev, Golden 2022



cat's eye flow model for  
brine convection cells

similar bounds  
for shear flows



rigorous Padé bounds from Stieltjes integral +  
analytical calculations of moments of measure

Lytle & Ackley 1996



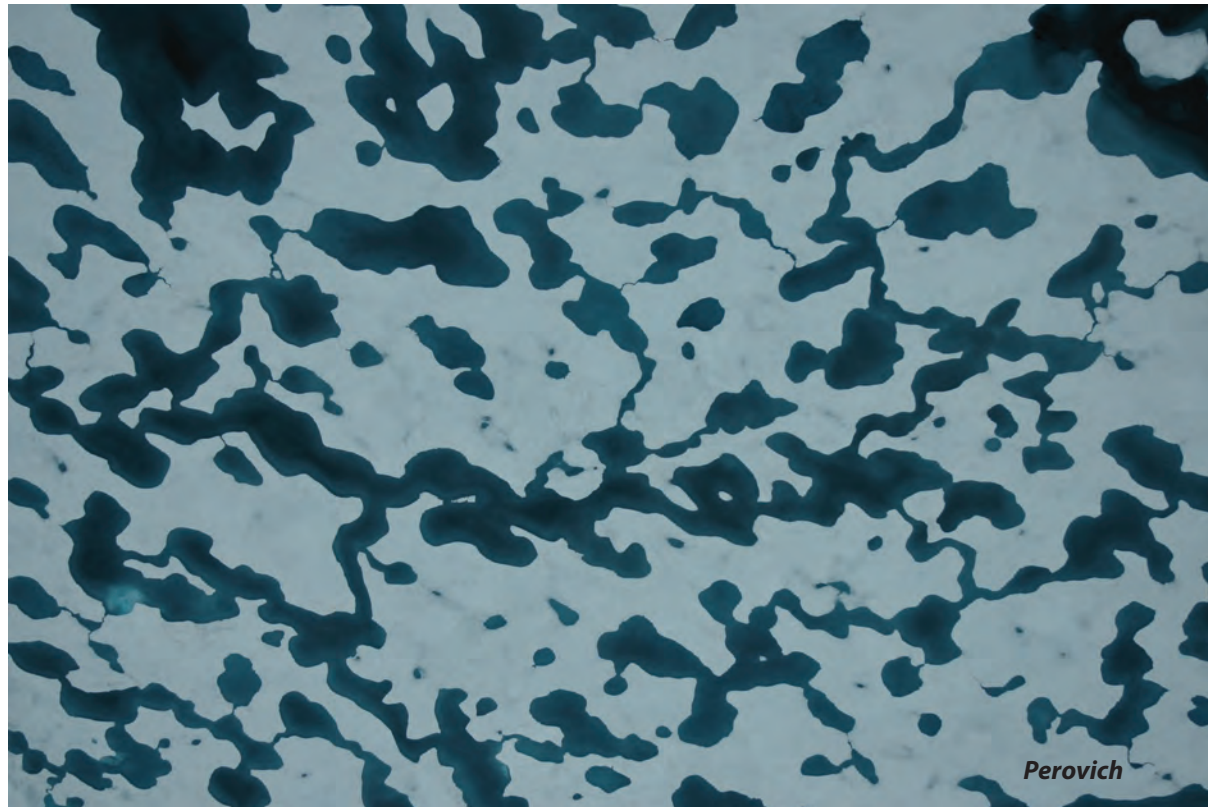
# *melt pond formation and albedo evolution:*

- *major drivers in polar climate*
- *key challenge for global climate models*

**numerical models of melt pond evolution, including topography, drainage (permeability), etc.**

Lüthje, Feltham,  
Taylor, Worster 2006  
Flocco, Feltham 2007

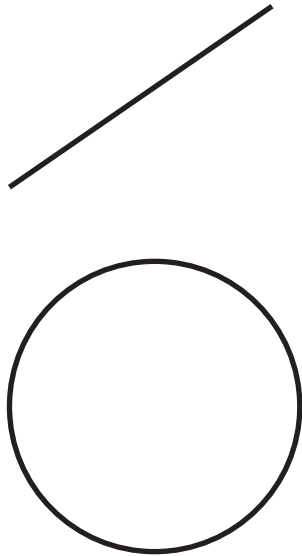
Skyllingstad, Paulson,  
Perovich 2009  
Flocco, Feltham,  
Hunke 2012



**Are there universal features of the evolution similar to phase transitions in statistical physics?**

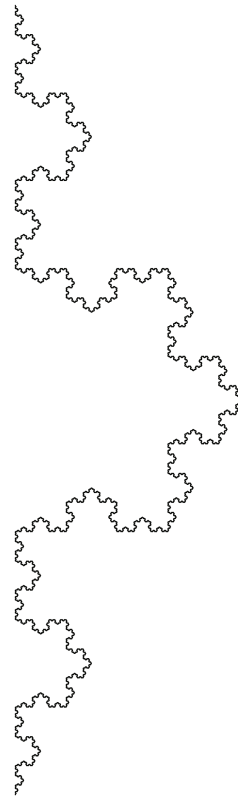
# *fractal curves in the plane*

*they wiggle so much that their dimension is  $>1$*



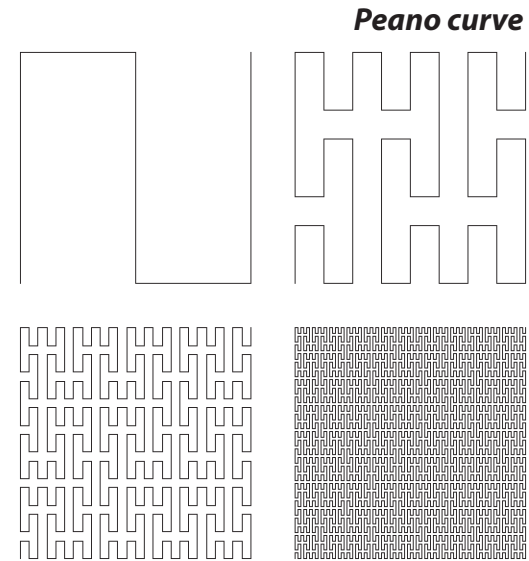
*simple curves*

$$D = 1$$

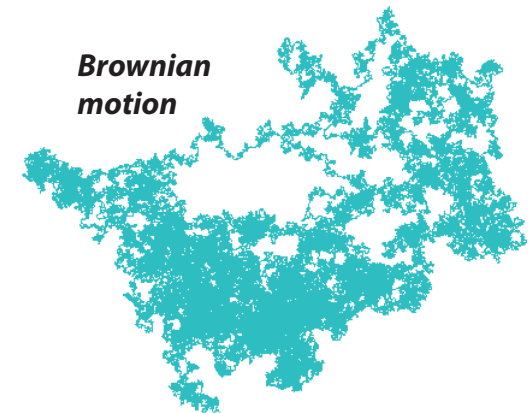


*Koch snowflake*

$$D = 1.26$$



*Peano curve*



*Brownian motion*

*space filling curves*

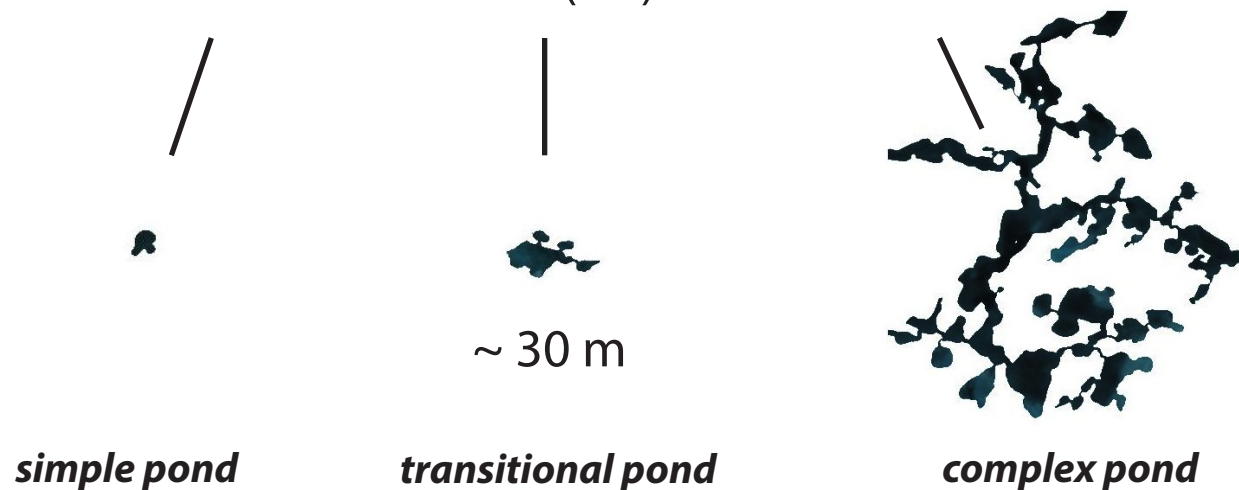
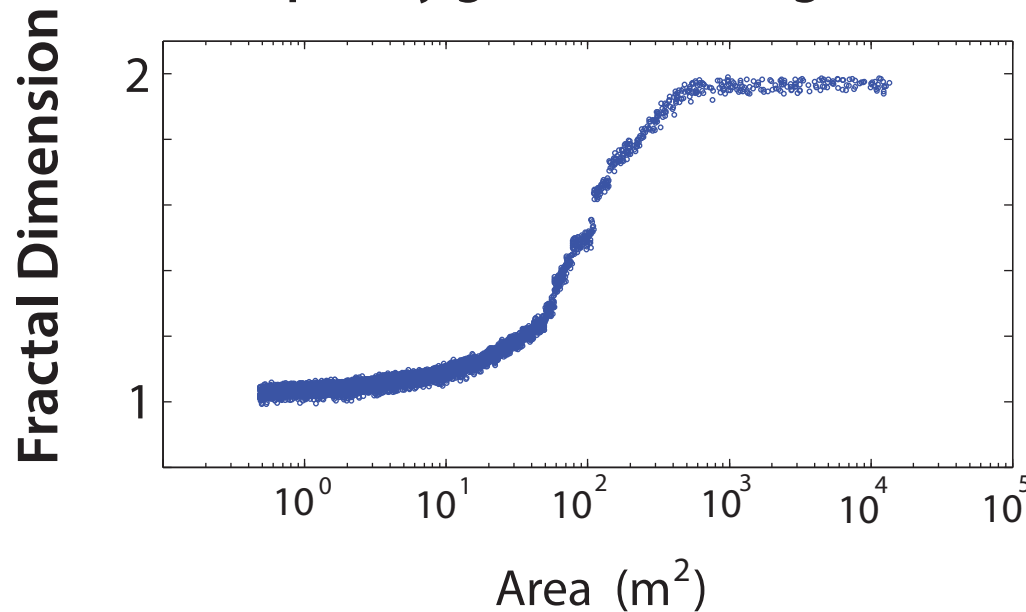
$$D = 2$$

# *Transition in the fractal geometry of Arctic melt ponds*

Christel Hohenegger, Bacim Alali, Kyle Steffen, Don Perovich, Ken Golden

*The Cryosphere, 2012*

complexity grows with length scale

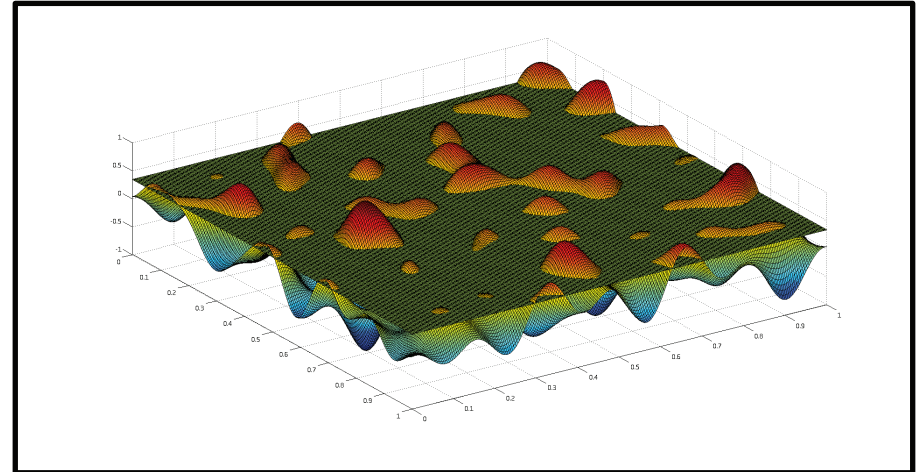
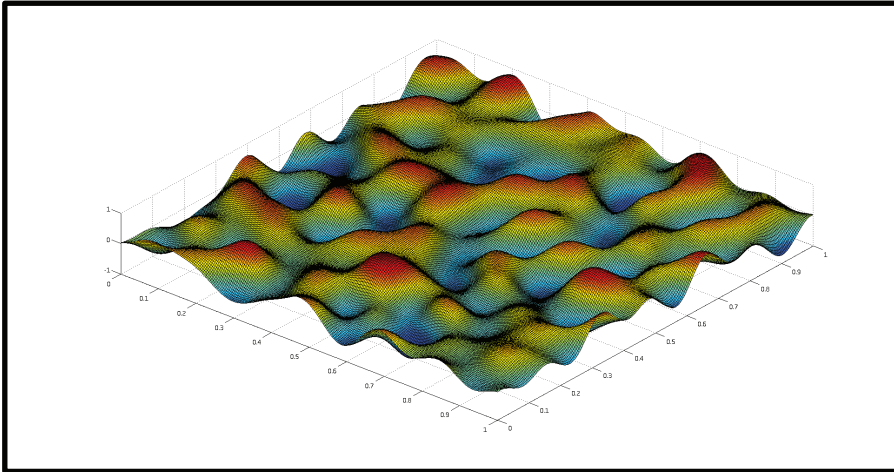




# Continuum percolation model for melt pond evolution

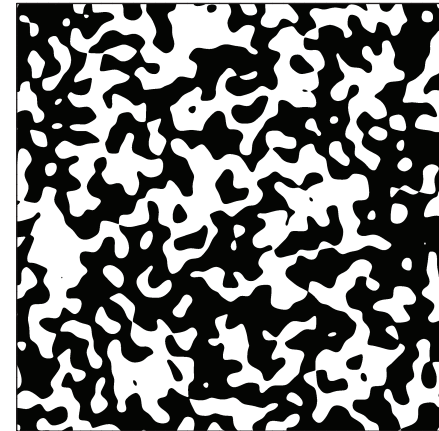
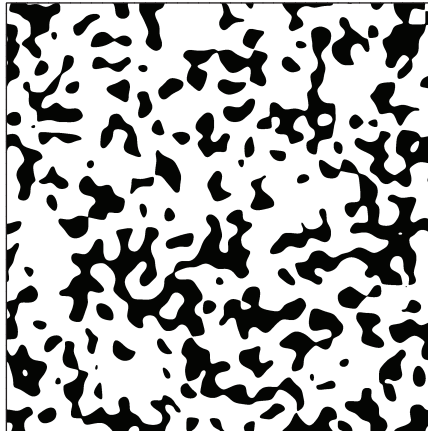
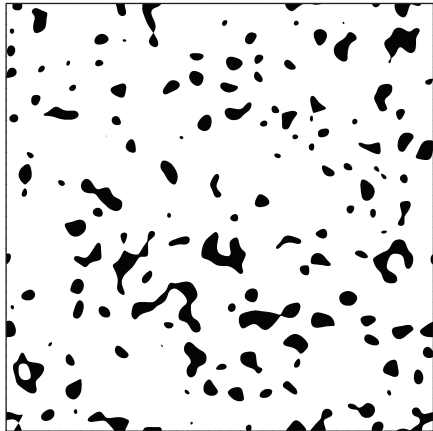
## *level sets of random surfaces*

*Brady Bowen, Court Strong, Ken Golden, J. Fractal Geometry 2018*



random Fourier series representation of surface topography

intersections of a plane with the surface define melt ponds

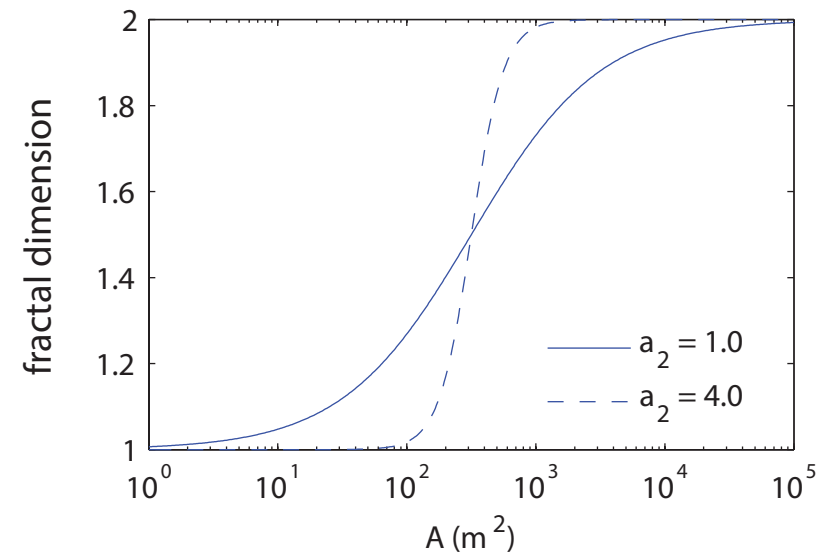
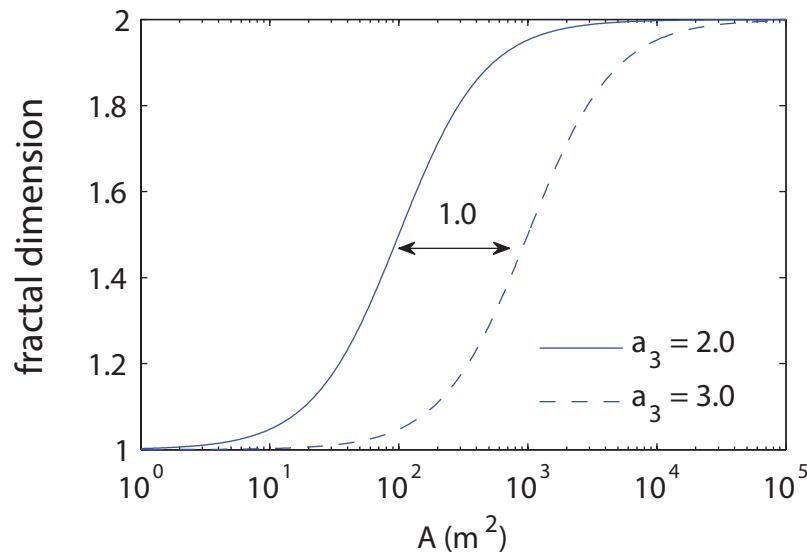


*electronic transport in disordered media*

*diffusion in turbulent plasmas*

*Isichenko, Rev. Mod. Phys., 1992*

# fractal dimension curves depend on statistical parameters defining random surface



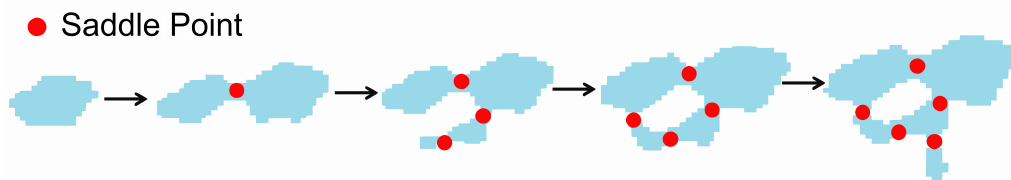
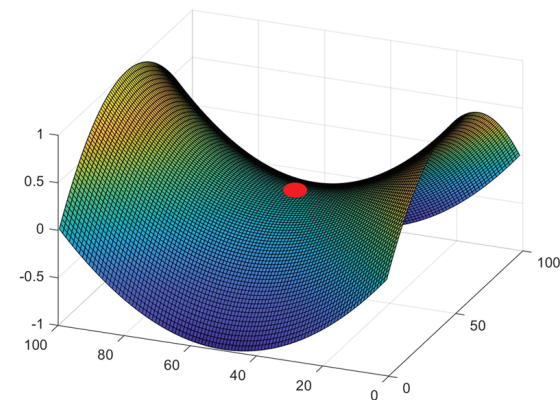
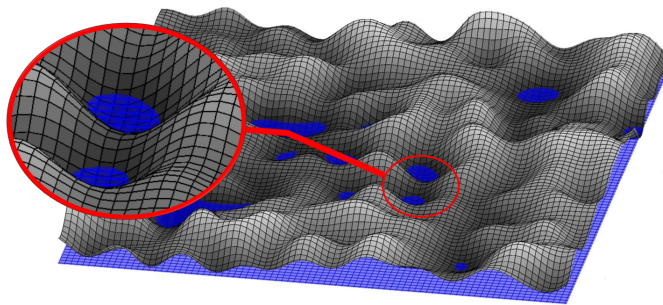
# Saddle points of the sea ice surface and the fractal geometry of Arctic melt ponds

*Physical Review Research* (invited, under review)

Ryleigh Moore, Jacob Jones, Dane Gollero,  
Court Strong, Ken Golden

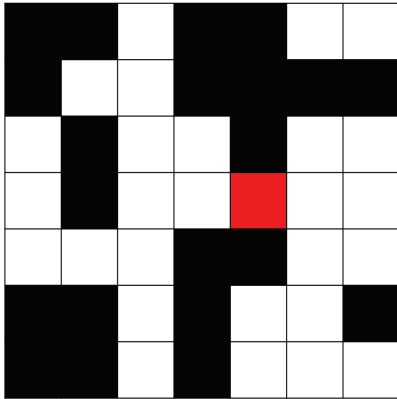
Several models replicate the transition in fractal dimension, but none explain how it arises.

We use Morse theory applied to the random surface model to show that **saddle points** play the critical role in the fractal transition.

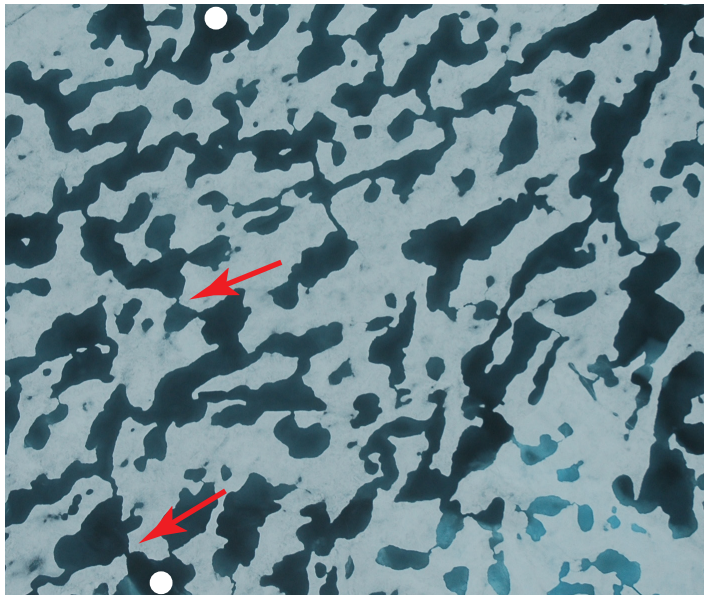


ponds coalesce  
(change topology) and  
complexify at saddle points



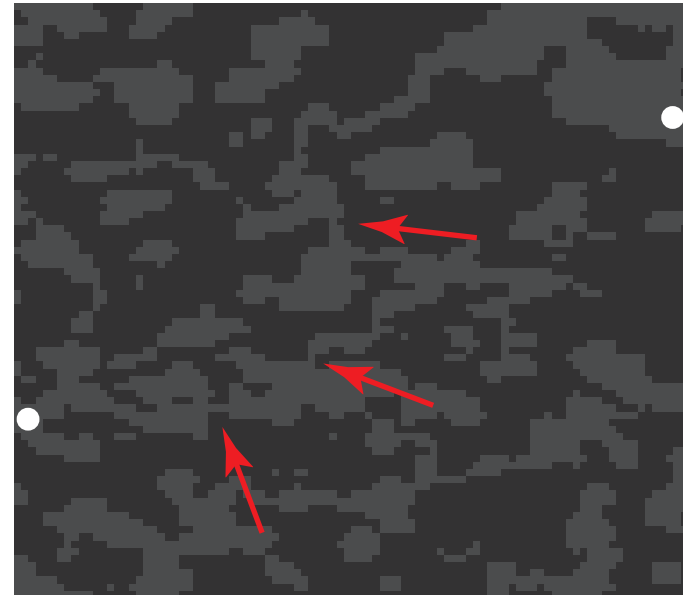


- Ponds connect through saddle points (Morse Theory).
- Red bonds in lattice percolation theory ~ saddle points.



Perovich

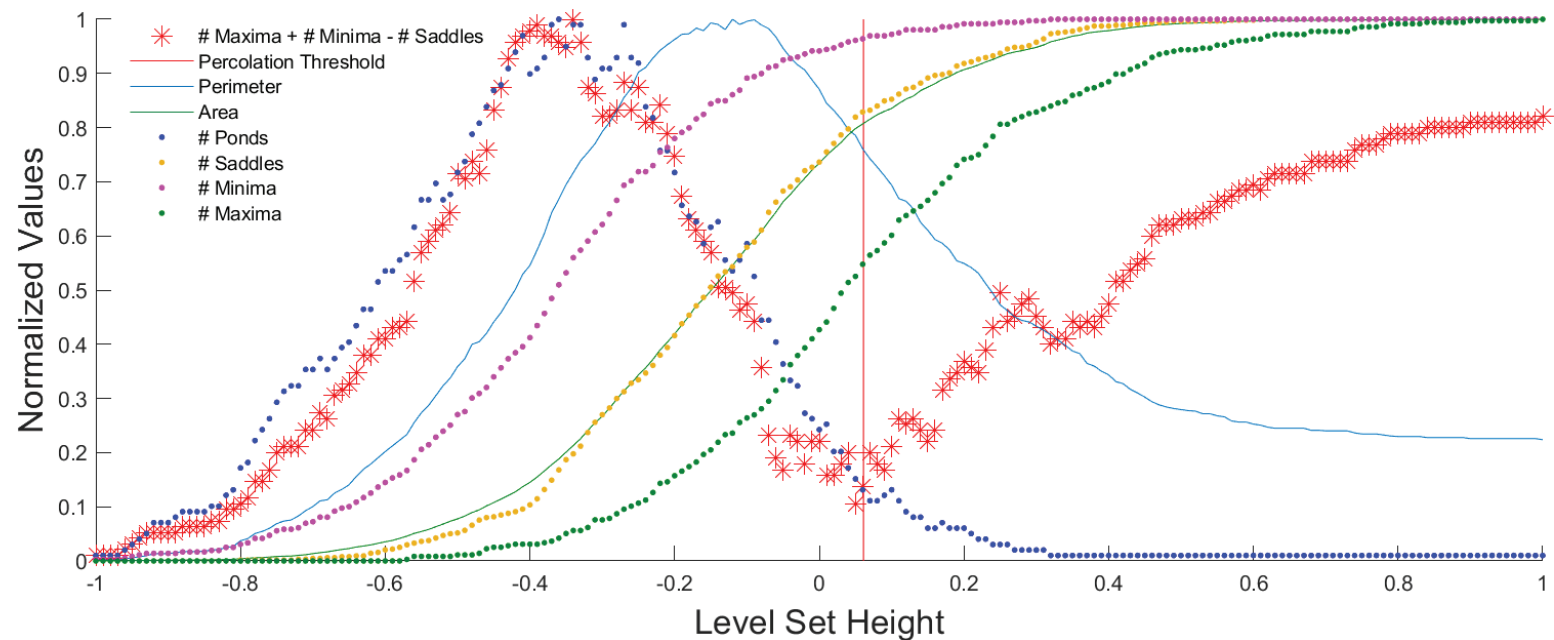
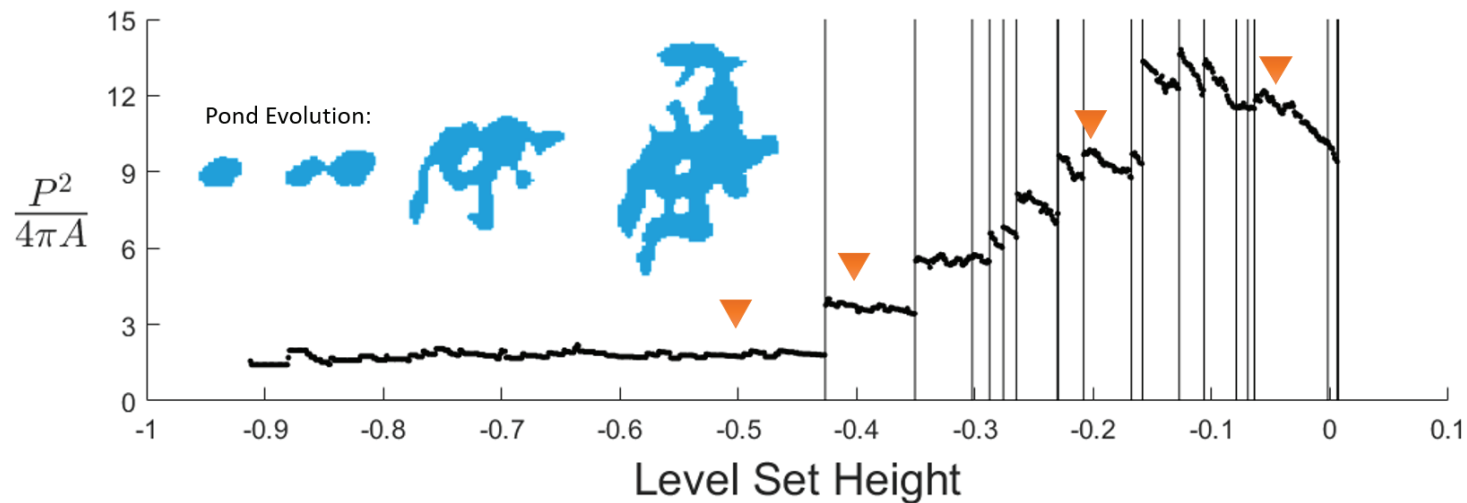
saddles



"red squares"

# Main results

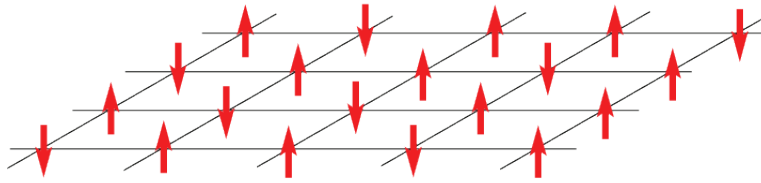
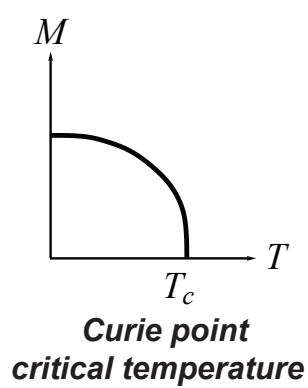
Isoperimetric quotient - as a proxy for fractal dimension - takes discrete jumps up when ponds coalesce at saddle points.



Euler characteristic reaches minimum at percolation threshold.

Horizontal fluid permeability “controlled” by saddles ~ electronic transport in 2D random potential.

# Ising Model for a Ferromagnet



applied  
magnetic  
field



$$s_i = \begin{cases} +1 & \text{spin up} & \text{blue} \\ -1 & \text{spin down} & \text{white} \end{cases}$$

$$\mathcal{H} = -H \sum_i s_i - J \sum_{\langle i,j \rangle} s_i s_j$$

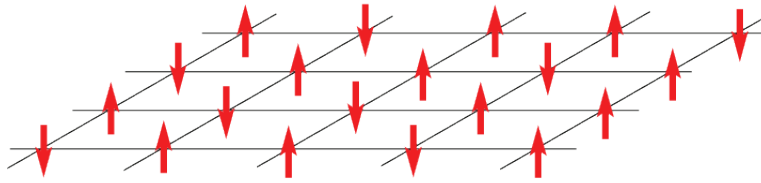
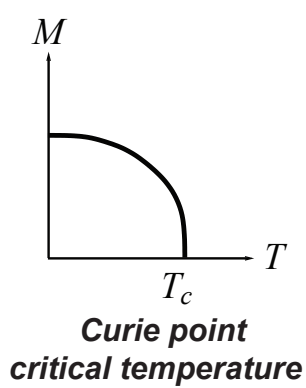
nearest neighbor Ising Hamiltonian

$$M(T, H) = \lim_{N \rightarrow \infty} \frac{1}{N} \left\langle \sum_j s_j \right\rangle$$

effective magnetization



# Ising Model for a Ferromagnet



$$s_i = \begin{cases} +1 & \text{spin up} \\ -1 & \text{spin down} \end{cases} \quad \begin{matrix} \text{blue} \\ \text{white} \end{matrix}$$

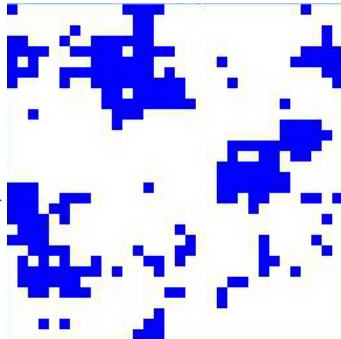
$$\mathcal{H} = -H \sum_i s_i - J \sum_{\langle i,j \rangle} s_i s_j$$

nearest neighbor Ising Hamiltonian

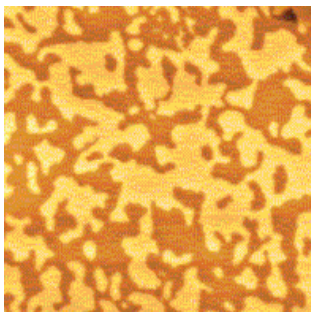
$$M(T, H) = \lim_{N \rightarrow \infty} \frac{1}{N} \left\langle \sum_j s_j \right\rangle$$

effective magnetization

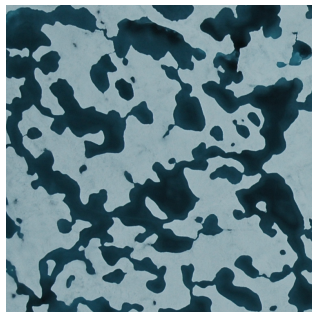
islands of like spins



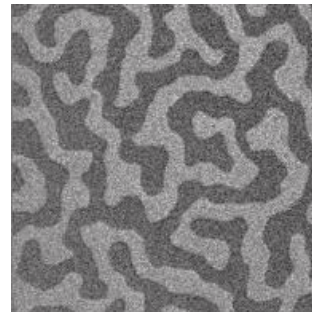
energy is lowered when nearby spins align with each other, forming **magnetic domains**



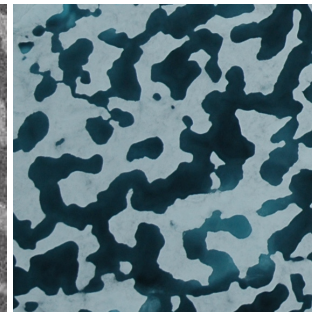
magnetic domains in cobalt



melt ponds (Perovich)



magnetic domains in cobalt-iron-boron



melt ponds (Perovich)

# Ising model for ferromagnets $\longrightarrow$ Ising model for melt ponds

Ma, Sudakov, Strong, Golden, *New J. Phys.*, 2019

$$\mathcal{H} = - \sum_i^N H_i s_i - J \sum_{\langle i,j \rangle}^N s_i s_j \quad s_i = \begin{cases} \uparrow & +1 \text{ water (spin up)} \\ \downarrow & -1 \text{ ice (spin down)} \end{cases}$$

random magnetic field  
represents snow topography

magnetization  $M$       pond area fraction  $F = \frac{(M+1)}{2}$       only nearest neighbor patches interact

$\sim$  *albedo*

Starting with random initial configurations, as Hamiltonian energy is minimized by Glauber spin flip dynamics, system “flows” toward metastable equilibria.

*Order from Disorder*

# Ising model for ferromagnets $\longrightarrow$ Ising model for melt ponds

Ma, Sudakov, Strong, Golden, *New J. Phys.*, 2019

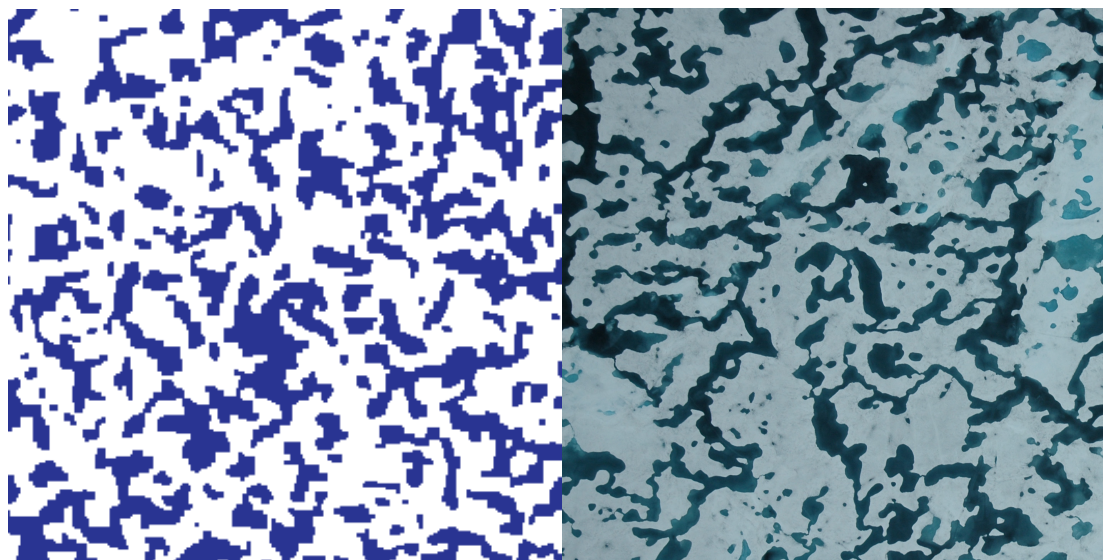
$$\mathcal{H} = - \sum_i^N H_i s_i - J \sum_{\langle i,j \rangle}^N s_i s_j \quad s_i = \begin{cases} \uparrow & +1 \text{ water (spin up)} \\ \downarrow & -1 \text{ ice (spin down)} \end{cases}$$

random magnetic field  
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magnetization  $M$       pond area fraction  $F = \frac{(M+1)}{2}$       only nearest neighbor patches interact  
 *$\sim$  albedo*

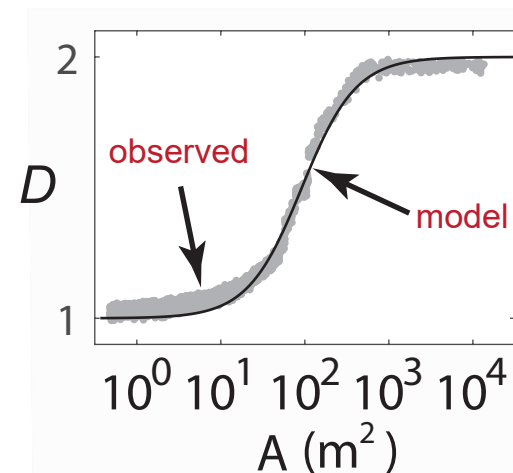
Starting with random initial configurations, as Hamiltonian energy is minimized by Glauber spin flip dynamics, system “flows” toward metastable equilibria.

## *Order from Disorder*



Ising  
model

melt pond  
photo (Perovich)



pond size  
distribution exponent

observed -1.5

(Perovich, et al. 2002)

model -1.58

*Scientific American  
EOS, PhysicsWorld, ...*

**ONLY MEASURED INPUT = LENGTH SCALE (GRID SIZE) from snow topography data**

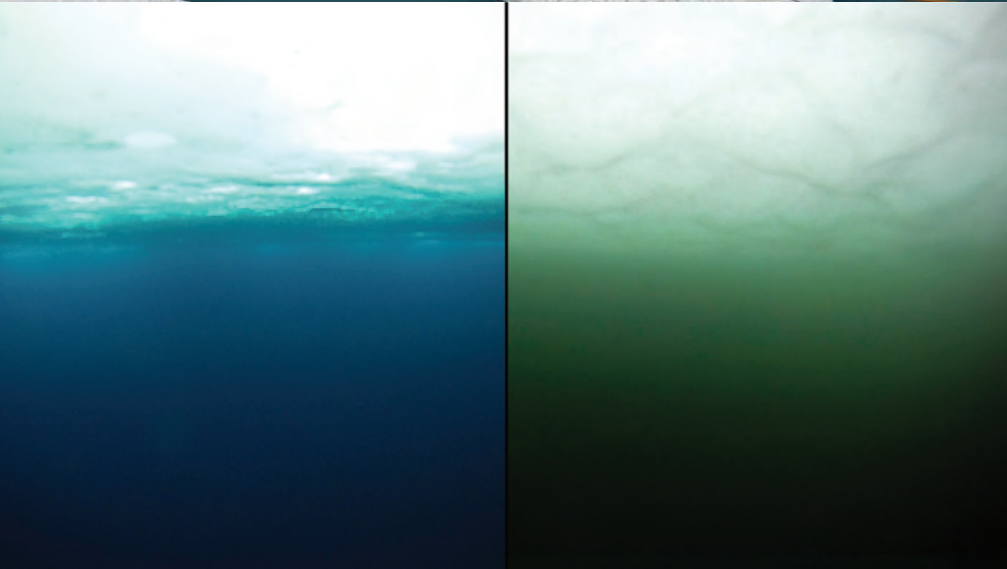




Perovich

Melt ponds control transmittance of solar energy through sea ice, impacting upper ocean ecology.

## WINDOWS



no bloom

bloom

massive under-ice **algal bloom**

Arrigo et al., *Science* 2012

***Have we crossed into a new ecological regime?***

The frequency and extent of sub-ice phytoplankton blooms in the Arctic Ocean

Horvat, Rees Jones, Iams, Schroeder, Flocco, Feltham, *Science Advances* 2017

The effect of melt pond geometry on the distribution of solar energy under first year sea ice

Horvat, Flocco, Rees Jones, Roach, Golden  
*Geophys. Res. Lett.* 2019

(2015 AMS MRC)

# **Uncertainty quantification and ecological dynamics in a model of a sea ice algae bloom, in prep. 2022**

**Jody Reimer, Fred Adler, Ken Golden, and Akil Narayan**

**Next week!**

**macroscale**

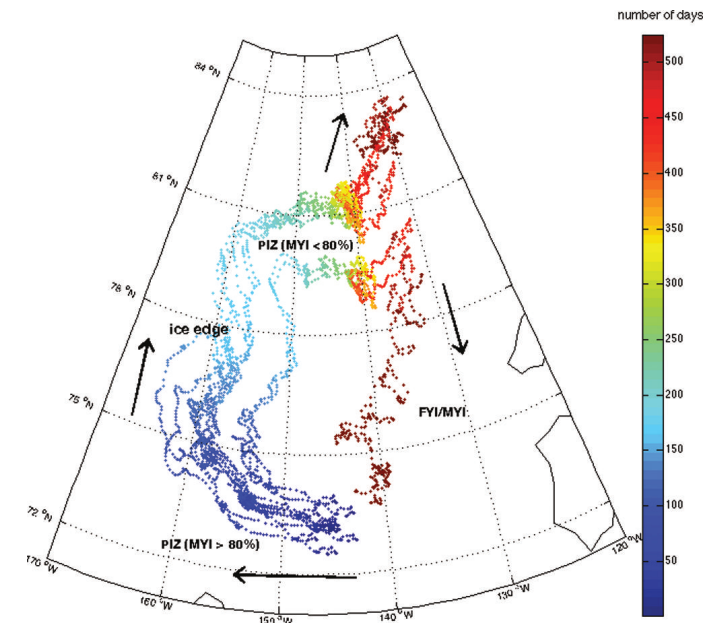


# Anomalous diffusion in sea ice dynamics

## *Ice floe diffusion in winds and currents*

### observations from GPS data:

Jennifer Lukovich, Jennifer Hutchings,  
David Barber, *Ann. Glac.* 2015



- On short time scales floes observed (buoy data) to exhibit Brownian-like behavior, but they are also being advected by winds and currents.
- Effective behavior is purely diffusive, sub-diffusive or super-diffusive depending on ice pack and advective conditions - **Hurst exponent**.

### modeling:

Huy Dinh, Ben Murphy, Elena Cherkaev,  
Court Strong, Ken Golden 2022

floe scale model to analyze transport regimes in  
terms of ice pack crowding, advective conditions

Delaney Mosier, Jennifer Hutchings, Jennifer Lukovich,  
Marta D'Elia, George Karniadakis, Ken Golden 2022

learning fractional PDE  
governing diffusion from data

# Floe Scale Model of Anomalous Diffusion in Sea Ice Dynamics

Huy Dinh, Ben Murphy, Elena Cherkaev, Court Strong, Ken Golden 2022

$$\langle |\mathbf{x}(t) - \mathbf{x}(0) - \langle \mathbf{x}(t) - \mathbf{x}(0) \rangle|^2 \rangle \sim t^\alpha$$

$\alpha$  = Hurst exponent

**diffusive**  $\alpha = 1$   
**sub-diffusive**  $\alpha < 1$   
**super-diffusive**  $\alpha > 1$

## Model Approximations

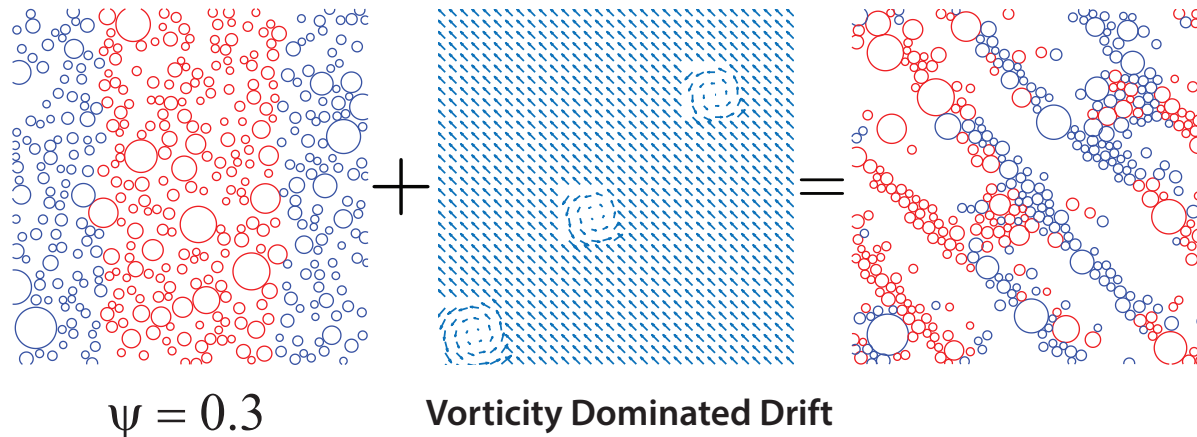
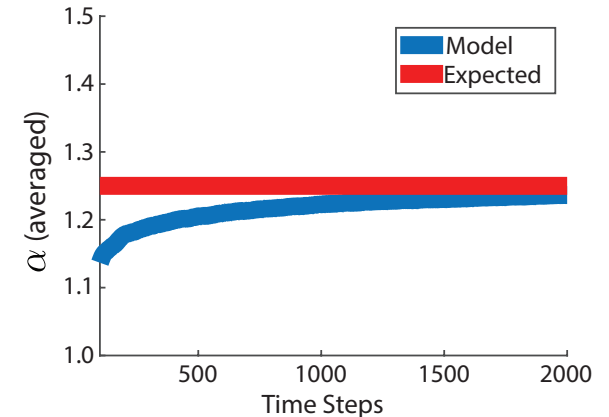
Power Law Size Distribution:  $N(D) \sim D^{-k}$

D. A. Rothrock and A. S. Thorndike Journal of Geophysical Research 1984

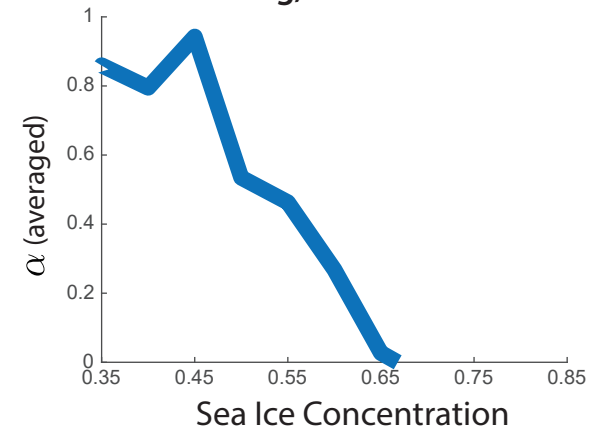
Floe-Floe Interactions: Linear Elastic Collisions

Advective Forcing: Passive, Linear Drag Law

Sparse Packing, Shear Dominated Drift



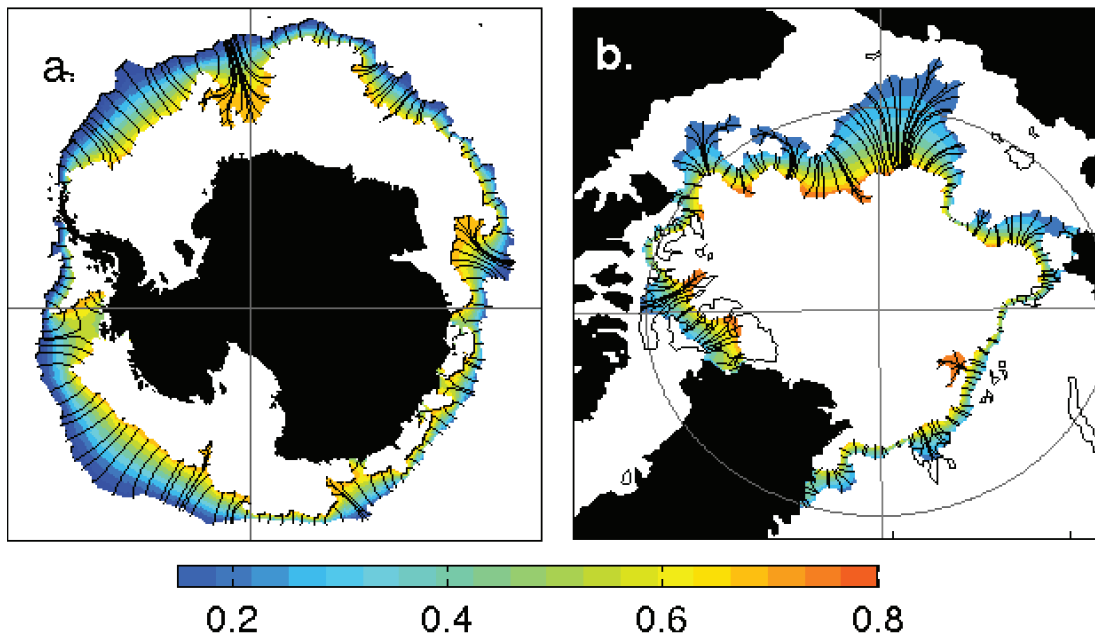
Crowding, Diffusive Drift



# Marginal Ice Zone

## MIZ

- biologically active region
- intense ocean-sea ice-atmosphere interactions
- region of significant wave-ice interactions



### MIZ WIDTH

fundamental length scale of  
ecological and climate dynamics

Strong, *Climate Dynamics* 2012

Strong and Rigor, *GRL* 2013

transitional region between  
dense interior pack ( $c > 80\%$ )  
sparse outer fringes ( $c < 15\%$ )

**How to objectively  
measure the “width”  
of this complex,  
non-convex region?**



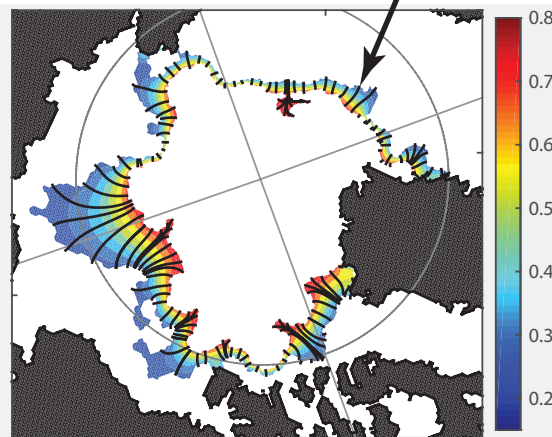
# Objective method for measuring MIZ width motivated by medical imaging and diagnostics

Strong, *Climate Dynamics* 2012  
Strong and Rigor, *GRL* 2013

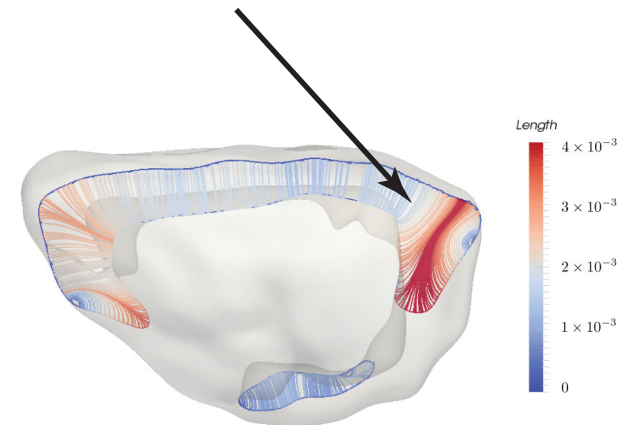
**39% widening**  
**1979 - 2012**

**“average” lengths of streamlines**

streamlines of a solution  
to Laplace’s equation



**Arctic Marginal Ice Zone**



**cross-section of the  
cerebral cortex of a rodent brain**

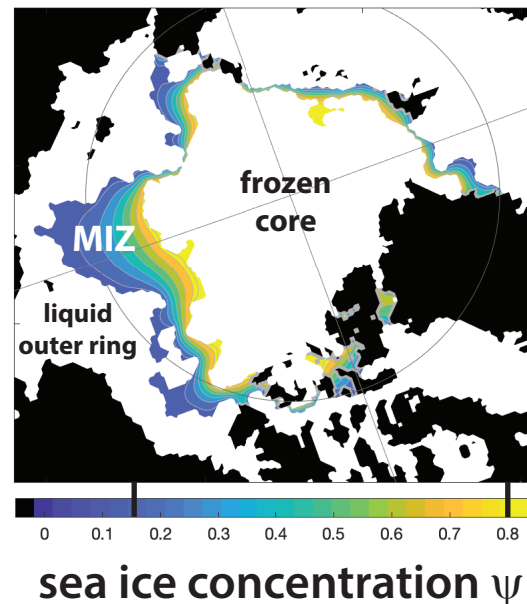
## ***analysis of different MIZ WIDTH definitions***

Strong, Foster, Cherkaev, Eisenman, Golden  
*J. Atmos. Oceanic Tech.* 2017

Strong and Golden  
*Society for Industrial and Applied Mathematics News*, April 2017

Model larger scale effective behavior  
with partial differential equations that  
*homogenize* complex local structure and dynamics.

## Arctic MIZ



Predict MIZ width and location with basin-scale phase change model.  
**dynamic transitional region - mushy layer - separating two “pure” phases**

seasonal and long term trends

C. Strong, E. Cherkaev, and K. M. Golden,  
Annual cycle of Arctic marginal ice zone location  
and width explained by phase change front model, 2022

# MIZ as a moving phase transition region

$$\rho c \frac{\partial T}{\partial t} = \nabla \cdot (k \nabla T) + S$$

$$S = [\rho(c_l - c_s)T + \rho L] \frac{\partial \psi}{\partial t}$$

$$\psi = 1 - \left( \frac{T - T_s}{T_l - T_s} \right)^\alpha$$

$$k_x = \left( \frac{\psi}{k_s} + \frac{1 - \psi}{k_l} \right)^{-1}$$

$$k_z = \psi k_s + (1 - \psi) k_l$$

**homogenization**

$\rho$  effective density

$T$  temperature

$c$  specific heat

$L$  latent heat of fusion

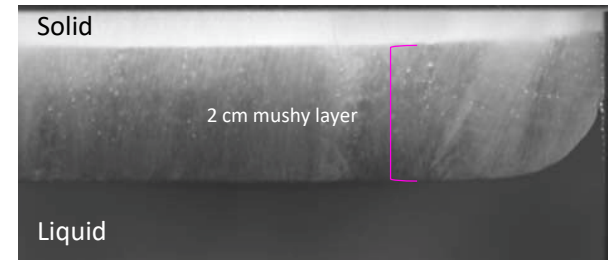
$S$  models nonlinear phase change

$\psi$  sea ice concentration

$k$  effective diffusivity

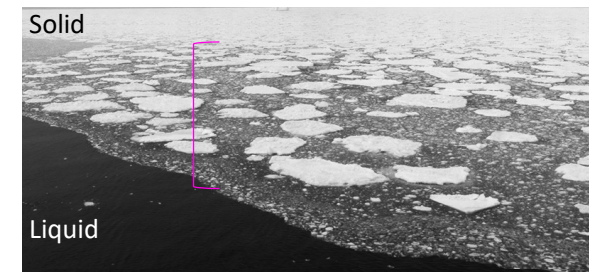
$l$  liquid,  $s$  solid

Classical small-scale application



NaCl-H<sub>2</sub>O in lab  
(Peppin et al., 2007; J. Fluid Mech.)

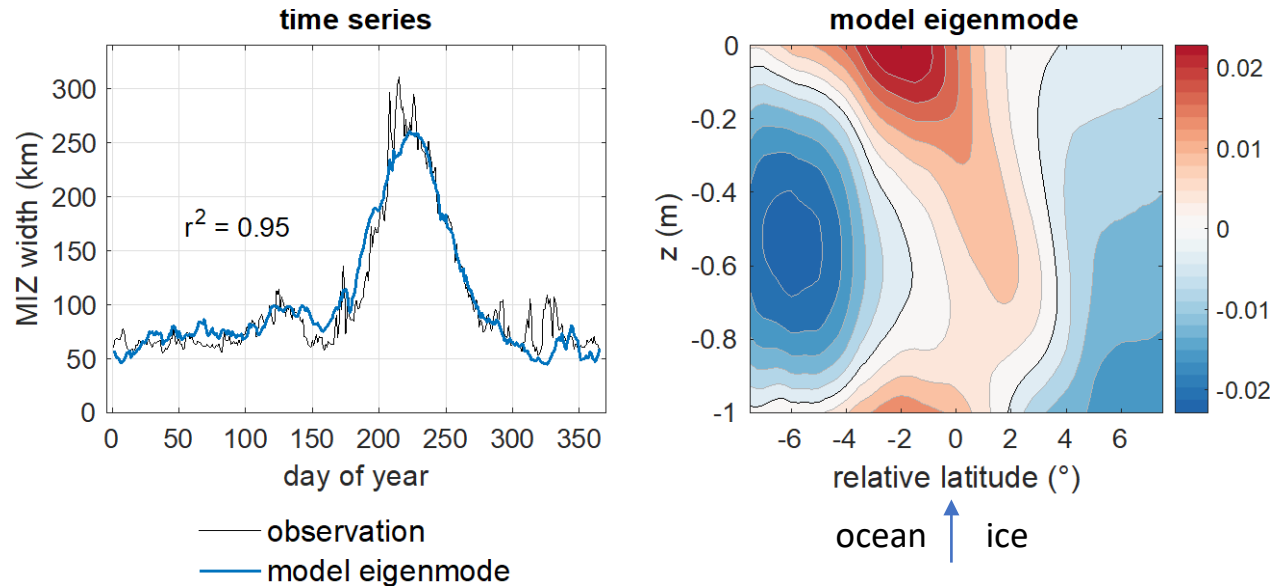
Macroscale application



- Develop multiscale PDE model for simulating phase transition fronts to predict MIZ seasonal cycles and decadal trends
- Model simulates MIZ as a large-scale mushy layer with effective thermal conductivity derived from physics of composite materials



# Model captures basic physics of MIZ dynamics



- Eigenmodes of temperature solution skillfully capture seasonal cycle of MIZ location and width.
- Eigenmode explaining MIZ width captures heat flux convergence into the MIZ layer from atmosphere above and oceanic mixed layer below.
- Model could ultimately be used to explain long term trends toward a wider and more poleward MIZ. Develop more sophisticated homogenization calculations; explore forcing scenarios and how to “drive” MIZ dynamics

# Learning the velocity field in an advection diffusion model for sea ice concentration

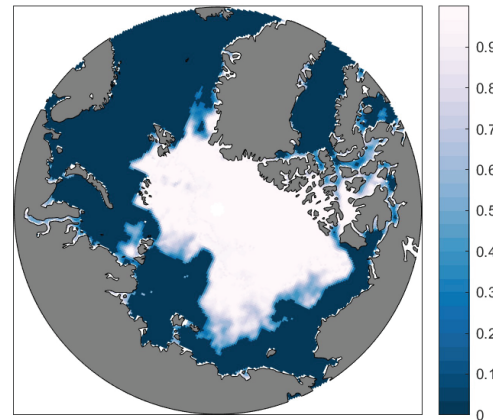
Eric Brown, Delaney Mosier, Bao Wang, Ken Golden, 2022

Goal: Develop PDE model to describe evolution of sea ice concentration field.

**advection diffusion model for sea ice concentration:**

$$\frac{\partial \psi}{\partial t} = -\mathbf{v} \cdot \nabla \psi + k \Delta \psi$$

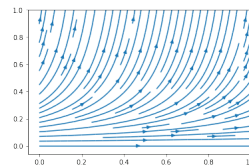
Use two-layer neural network to **infer advective fields** based on satellite imagery



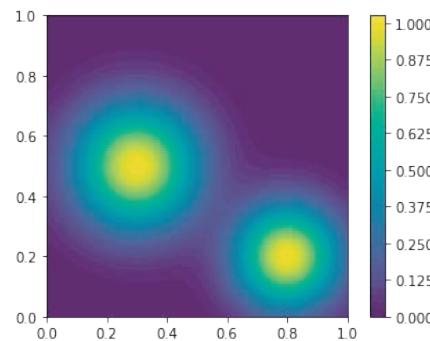
National Snow and Ice Data Center

**discretized satellite concentration data**

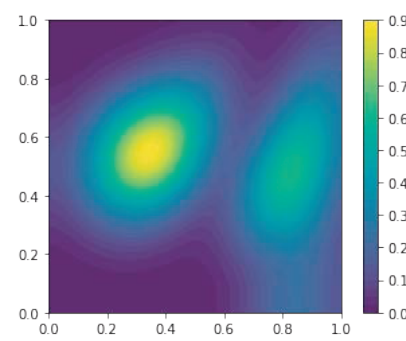
Figure 1. Arctic sea ice concentration in early August 2012.



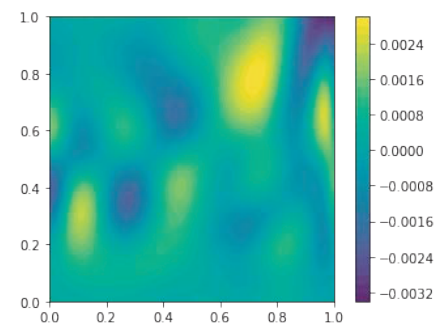
**learned velocity**



**initital test concentration**



**predicted concentration**



**error**

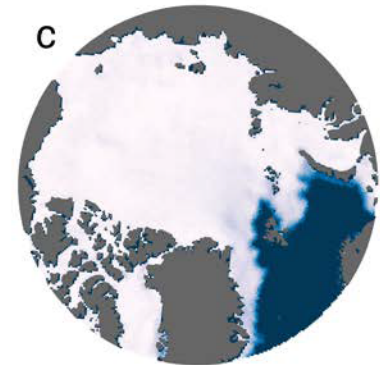
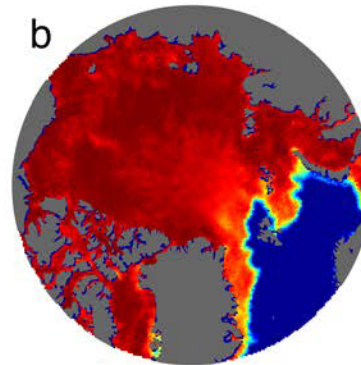
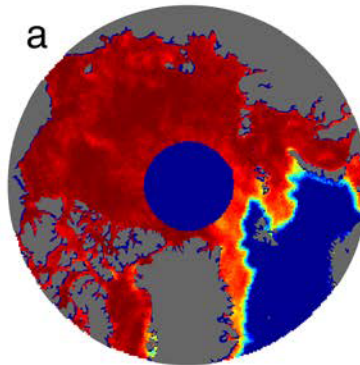
**2.5% absolute error** in preliminary study

# Filling the polar data gap with partial differential equations

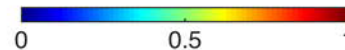
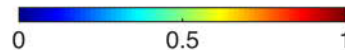
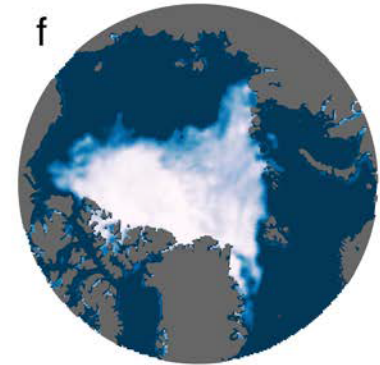
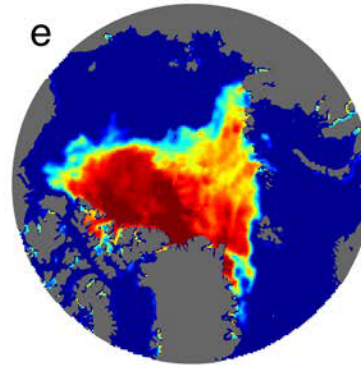
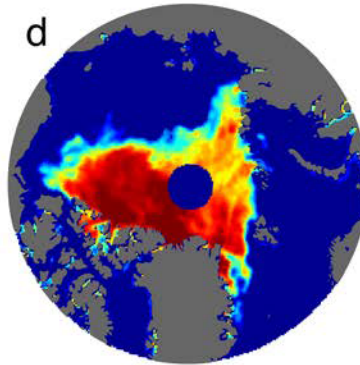
hole in satellite coverage  
of sea ice concentration field

previously assumed  
ice covered

Gap radius: 611 km  
06 January 1985



Gap radius: 311 km  
30 August 2007



$$\Delta\psi=0$$

**fill = harmonic function with  
learned stochastic term**

Strong and Golden, *Remote Sensing* 2016  
Strong and Golden, *SIAM News* 2017

**NOAA/NSIDC Sea Ice Concentration CDR  
product update will use our PDE method.**



# Conclusions

1. Sea ice is a fascinating multiscale composite with structure similar to many other natural and man-made materials.
2. Mathematical methods developed for sea ice advance theories of composites and inverse problems in science and engineering.
3. **Homogenization and statistical physics help *link scales in sea ice and composites***; provide rigorous methods for finding effective behavior; advance sea ice representations in climate models.
4. **Inverse problems of many types** arise naturally in studying sea ice and the impact of climate change in Earth's polar regions.
5. Field experiments are essential to developing relevant mathematics.
6. Our research is helping to **improve projections of climate change**, the fate of Earth's sea ice packs, and the ecosystems they support.

# University of Utah Sea Ice Modeling Group (2017-2021)

**Senior Personnel:** Ken Golden, Distinguished Professor of Mathematics  
Elena Cherkaev, Professor of Mathematics  
Court Strong, Associate Professor of Atmospheric Sciences  
Ben Murphy, Adjunct Assistant Professor of Mathematics

**Postdoctoral Researchers:** Noa Kraitzman (now at ANU), Jody Reimer

**Graduate Students:** Kyle Steffen (now at UT Austin with Clint Dawson)  
Christian Sampson (now at UNC Chapel Hill with Chris Jones)  
Huy Dinh (now a sea ice MURI Postdoc at NYU/Courant)  
Rebecca Hardenbrook  
David Morison (Physics Department)  
Ryleigh Moore  
Delaney Mosier  
Daniel Hallman

**Undergraduate Students:** Kenzie McLean, Jacqueline Cinella Rich,  
Dane Gollero, Samir Suthar, Anna Hyde,  
Kitsel Lusted, Ruby Bowers, Kimball Johnston,  
Jerry Zhang, Nash Ward, David Gluckman

**High School Students:** Jeremiah Chapman, Titus Quah, Dylan Webb

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**Sea Ice Ecology Group**    Postdoc Jody Reimer, Grad Student Julie Sherman,  
Undergraduates Kayla Stewart, Nicole Forrester





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# Notices

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MATHEMATICAL  
SOCIETY

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*The cover is based on "Modeling Sea Ice,"  
page 1535.*



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# Modeling Sea Ice



*Kenneth M. Golden, Luke G. Bennetts,  
Elena Cherkaev, Ian Eisenman, Daniel Feltham,  
Christopher Horvat, Elizabeth Hunke,  
Christopher Jones, Donald K. Perovich,  
Pedro Ponte-Castañeda, Courtenay Strong,  
Deborah Sulsky, and Andrew J. Wells*

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*Communicated by Notices Associate Editor Reza Malek-Madani.*

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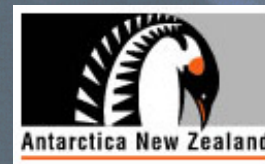
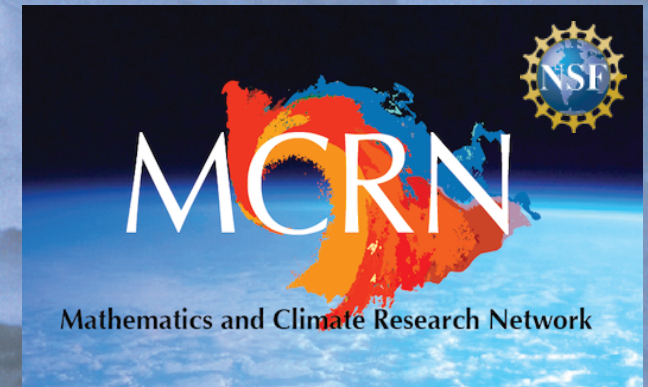
# THANK YOU

## Office of Naval Research

Applied and Computational Analysis Program  
Arctic and Global Prediction Program

## National Science Foundation

Division of Mathematical Sciences  
Division of Polar Programs

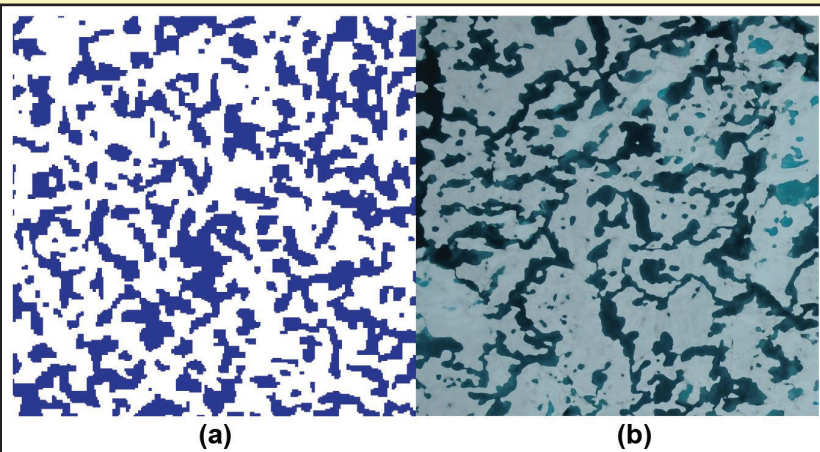


***Buchanan Bay, Antarctica    Mertz Glacier Polynya Experiment    July 1999***



## Special Issue on the Mathematics of Planet Earth

Read about the application of mathematics and computational science to issues concerning invasive populations, Arctic sea ice, insect flight, and more in this Planet Earth **special issue**!



**Figure 3.** Comparison of real Arctic melt ponds with metastable equilibria in our melt pond Ising model. **3a.** Ising model simulation. **3b.** Real melt pond photo. Figure 3a courtesy of Yiping Ma, 3b courtesy of Donald Perovich.

Vast labyrinthine ponds on the surface of melting Arctic sea ice are key players in the polar climate system and upper ocean ecology. Researchers have adapted the Ising model, which was originally developed to understand magnetic materials, to study the geometry of meltwater's distribution over the sea ice surface. In an article on page 5, Kenneth Golden, Yiping Ma, Courtenay Strong, and Ivan Sudakov explore model predictions.

## Controlling Invasive Populations in Rivers

By Yu Jin and Suzanne Lenhart

Flow regimes can change significantly over time and space and strongly impact all levels of river biodiversity, from the individual to the ecosystem. Invasive species in rivers—such as bighead and silver carp, as well as quagga and zebra mussels—continue to cause damage. Management of these species may include targeted adjustment of flow rates in rivers, based on recent research that examines the effects of river morphology and water flow on rivers' ecological statuses. While many previous methodologies rely on habitat suitability models or oversimplification of the hydrodynamics, few studies have focused on the integration of ecological dynamics into water flow assessments.

Earlier work yielded a hybrid modeling approach that directly links river hydrology with stream population models [3]. The hybrid model's hydrodynamic component is based on the water depth in a gradually varying river structure. The model derives the steady advective flow from this structure and relates it to flow features like water discharge, depth, velocity, cross-

sectional area, bottom roughness, bottom slope, and gravitational acceleration. This approach facilitates both theoretical understanding and the generation of quantitative predictions, thus providing a way for scientists to analyze the effects of river fluctuations on population processes.

When a population spreads longitudinally in a one-dimensional (1D) river with spatial heterogeneities in habitat and temporal fluctuations in discharge, the resulting hydrodynamic population model is

$$N_t = -A_t(x, t) \frac{N}{A(x, t)} + \frac{1}{A(x, t)} \left( D(x, t) A(x, t) N_x \right)_x - \frac{Q(t)}{A(x, t)} N_x + rN \left( 1 - \frac{N}{K} \right)$$

$$\begin{aligned} N(0, t) &= 0 & \text{on } (0, T), x = 0, \\ N_x(L, t) &= 0 & \text{on } (0, T), x = L, \\ N(x, 0) &= N_0(x) & \text{on } (0, L), t = 0 \end{aligned} \quad (1)$$

See **Invasive Populations** on page 4

## Modeling Resource Demands and Constraints for COVID-19 Intervention Strategies

By Erin C.S. Acquesta, Walt Beyeler, Pat Finley, Katherine Klise, Monear Makvandi, and Emma Stanislawski

As the world desperately attempts to control the spread of COVID-19, the need for a model that accounts for realistic trade-offs between time, resources, and corresponding epidemiological implications is apparent. Some early mathematical models of the outbreak compared trade-offs for non-pharmaceutical interventions [3], while others derived the necessary level of test coverage for case-based interventions [4] and demonstrated the value of prioritized testing for close contacts [7].

Isolated analyses provide valuable insights, but real-world intervention strategies are interconnected. Contact tracing is the lynchpin of infection control [6] and forms the basis of prioritized testing. Therefore, quantifying the effectiveness of contact tracing is crucial to understanding the real-life implications of disease control strategies.

### Contact Tracing Demands

Contact tracers are skilled, culturally competent interviewers who apply their knowledge of disease and risk factors when notifying people who have come into contact with COVID-19-infected individuals. They also continue to monitor the situation after case investigations [1].

Case investigation consists of four steps:

1. Identify and notify cases
2. Interview cases
3. Locate and notify contacts
4. Monitor contacts.

Most health departments are implementing case investigation, contact identification, and quarantine to disrupt COVID-19 transmission. The timeliness of contact tracing is constrained by the length of the infectious period, the turn-around time for testing and result reporting, and the ability to successfully reach and interview patients and their contacts. The European Centre for Disease Prevention and Control approximates that contact tracers spend one to two hours conducting an interview [2]. Estimates regarding the timelines of other steps are limited to subject matter expert elicitation and can vary based on cases' access to phone service or willingness to participate in interviews.

### Bounded Exponential

The fundamental structure of our model follows traditional susceptible-exposed-infected-recovered (SEIR) compartmental modeling [5]. We add an asymptomatic population  $A$ , a hospitalized population  $H$ , and disease-related deaths  $D$ , as well as corresponding quarantine states. We define the states  $\{S_i, E_i, A_i, I_i, H, R, D\}_{i=0,1}$  for our compartments, such that  $i=0$  and  $i=1$

correspond to unquarantined and quarantined respectively. Rather than focus on the dynamics that are associated with the state transition diagram in Figure 1, we introduce a formulation for the real-time demands on contact tracers' time as a function of infection prevalence, while also respecting constraints on resources.

When the work that is required to investigate new cases and monitor existing contacts exceeds available resources, a backlog develops. To simulate this backlog, we introduce a new compartment  $C$  for tracking the dynamic states of cases:

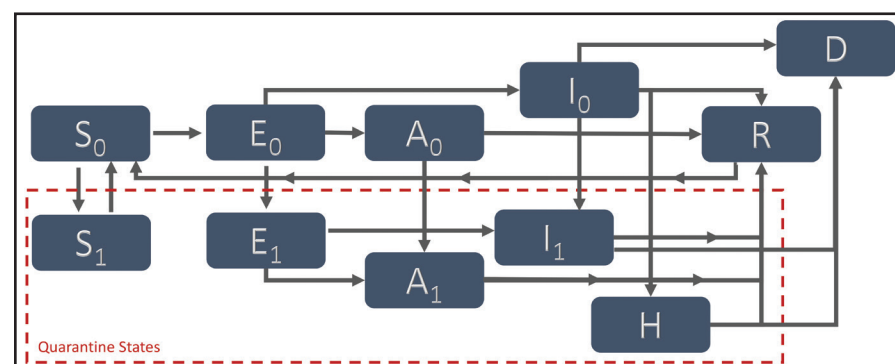
$$\frac{dC}{dt} = [flow_{in}] - [flow_{out}].$$

Flow into the backlog compartment, represented by  $[flow_{in}]$ , reflects case identification that is associated with the following transitions in the model:

- The rate of random testing:  $q_{rA}(t)A_0(t) \rightarrow A_1(t)$  and  $q_{rI}(t)I_0(t) \rightarrow I_1(t)$
- Testing triggered by contact tracing:  $q_{tA}(t)A_0(t) \rightarrow A_1(t)$ ,  $q_{tI}(t)I_0(t) \rightarrow I_1(t)$ , and  $q_{tE}(t)E_i(t) \rightarrow \{A_i(t), I_i(t)\}$
- The population that was missed by the non-pharmaceutical interventions that require hospitalization:  $\tau_{IH}(t)I_0(t) \rightarrow H(t)$ .

Here,  $q_{rs}(t)$  defines the time-dependent rate of random testing,  $q_{ts}(t)$  signifies the time-dependent rate of testing that is triggered by contact tracing, and  $\tau_{IH}$  is the inverse of the expected amount of time for which an infected individual is symptomatic before hospitalization. These terms collectively provide the simulated number of newly-identified positive COVID-19 cases. However, we also need the average number of contacts per case. We thus define function  $\mathcal{K}(\kappa, T_s, \phi_\kappa)$  that depends on the average number of contacts a day ( $\kappa$ ), the average number of days for which an individual is infectious before going into isolation ( $T_s$ ), and the likelihood that the individual

See **COVID-19 Intervention** on page 3



**Figure 1.** Disease state diagram for the compartmental infectious disease model. Figure courtesy of the authors.

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