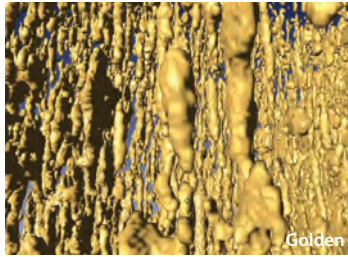
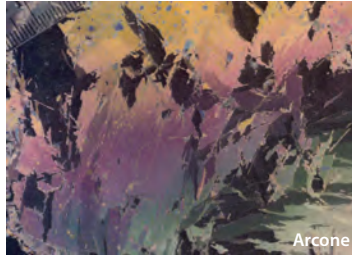


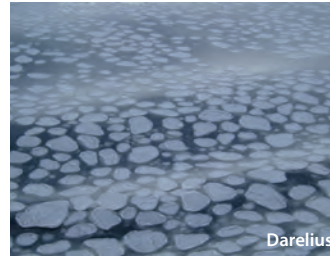
millimeters



centimeters



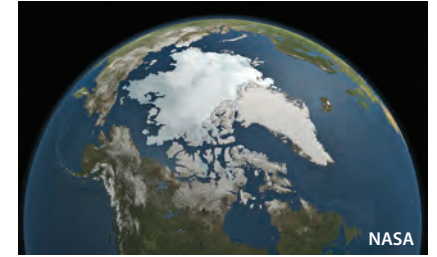
meters



kilometers



10^3 kilometers

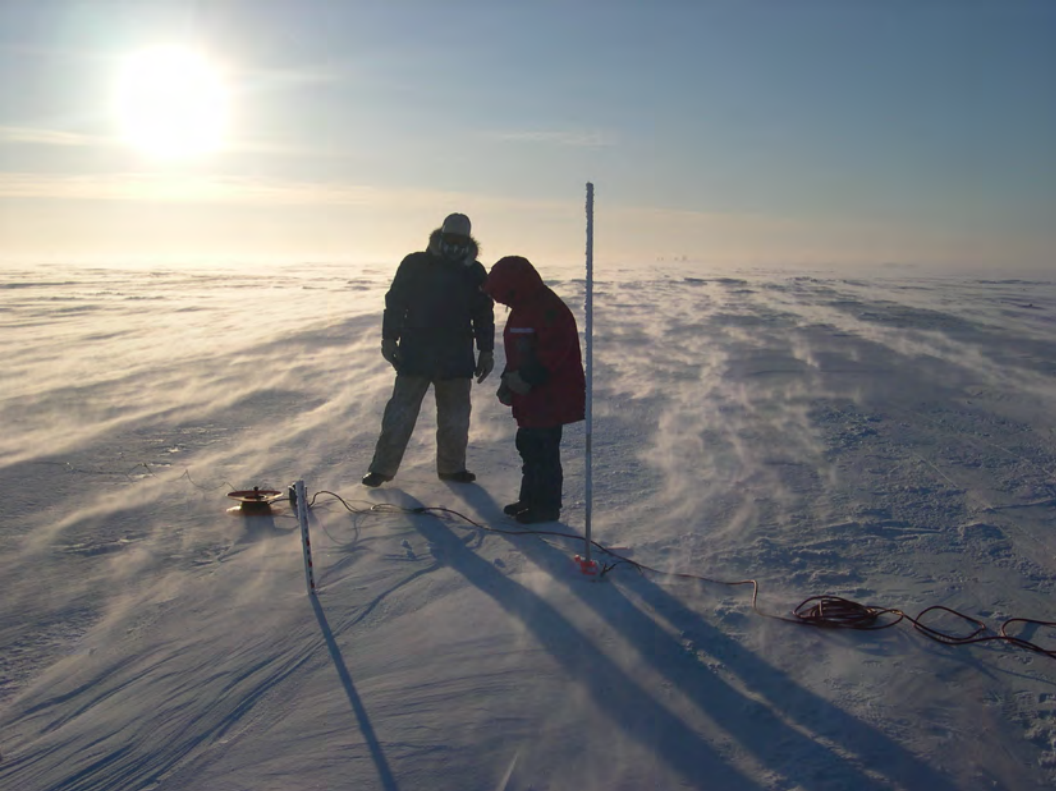


From Micro to Macro in the Fluid Dynamics of Sea Ice

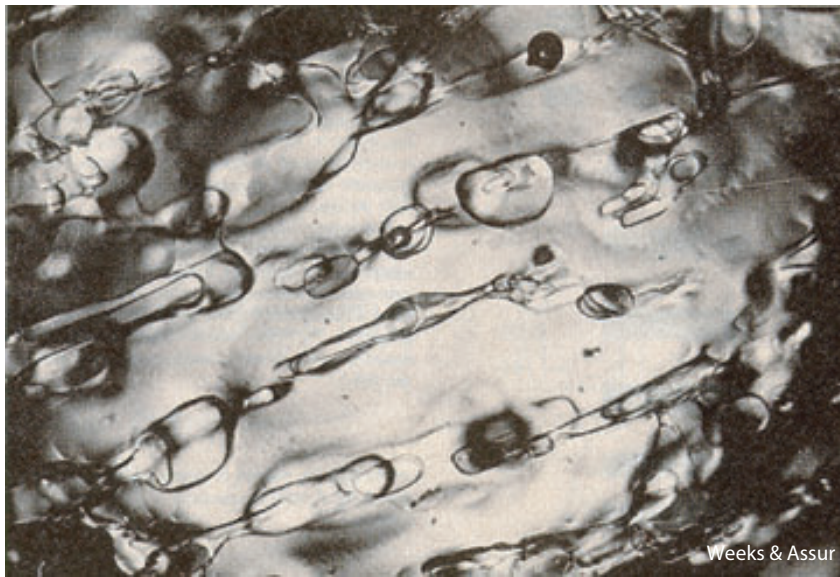
Kenneth M. Golden
Dept. of Mathematics, Univ. of Utah



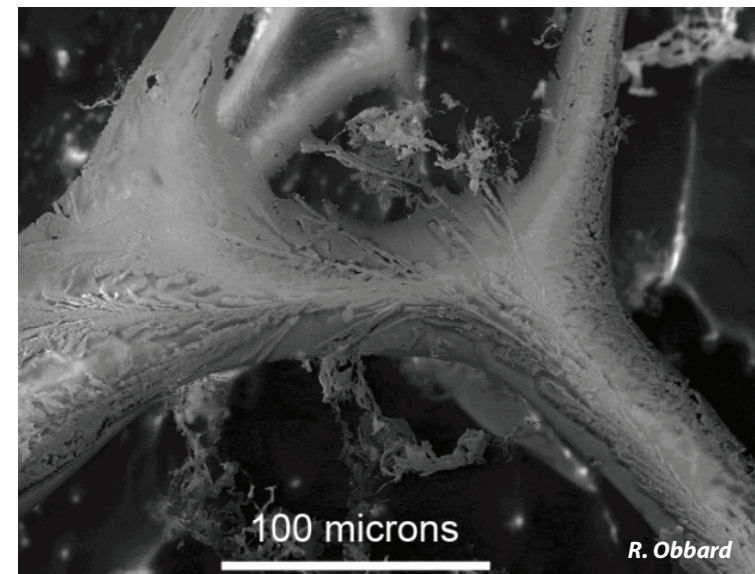
Oberwolfach Workshop on
Mathematical Advances in
Geophysical Fluid Dynamics
November 13 - 19, 2022



*sea ice may appear to be a
barren, impermeable cap ...*



brine inclusions in sea ice (mm)



micro - brine channel (SEM)

***sea ice is a
porous composite***

pure ice with brine, air, and salt inclusions

brine channels (cm)



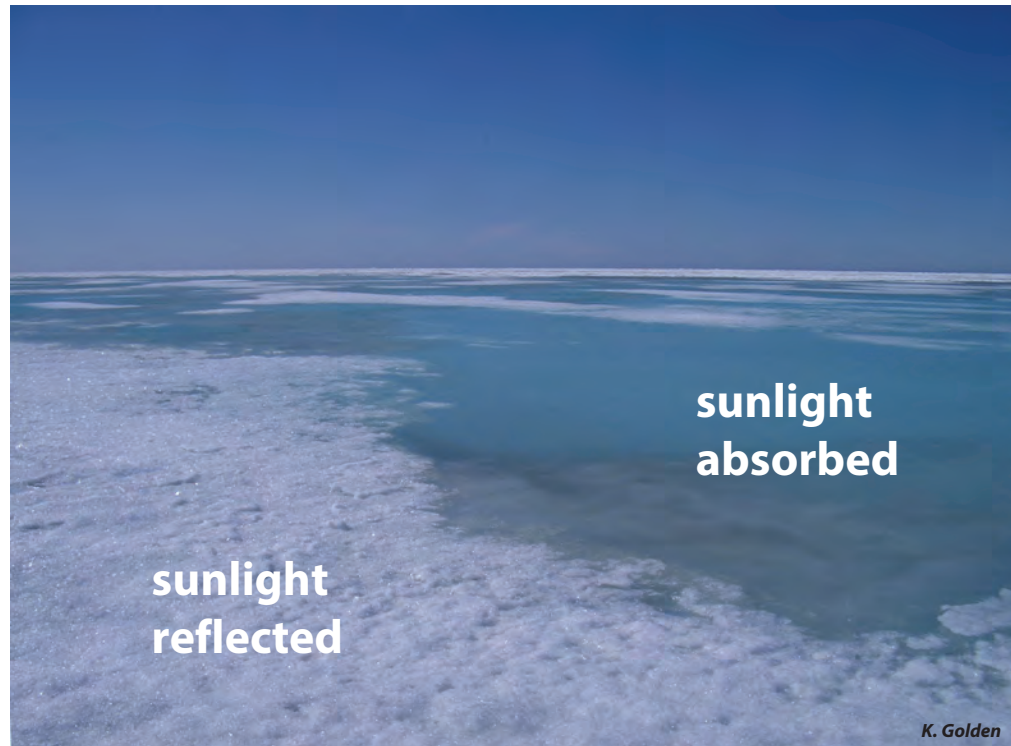
horizontal section



vertical section

fluid flow through the porous microstructure of sea ice governs key processes in polar climate and ecosystems

*evolution of Arctic melt ponds and sea ice **albedo***



nutrient flux for algal communities



***Antarctic surface flooding
and snow-ice formation***

September
snow-ice
estimates

- *evolution of salinity profiles*
- *ocean-ice-air exchanges of heat, CO₂*

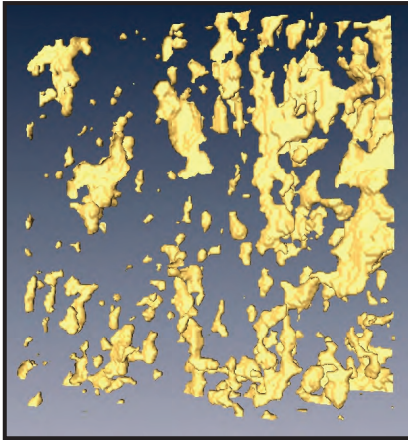
Sea Ice is a Multiscale Composite Material

microscale

brine inclusions

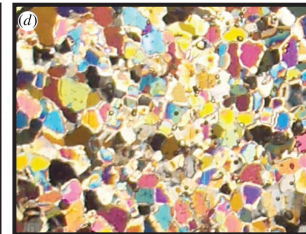
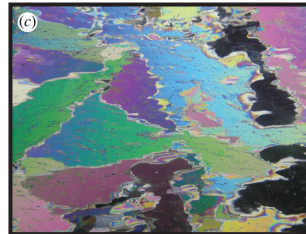
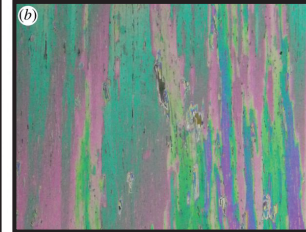


Weeks & Assur 1969



H. Eicken
Golden et al. GRL 2007

polycrystals

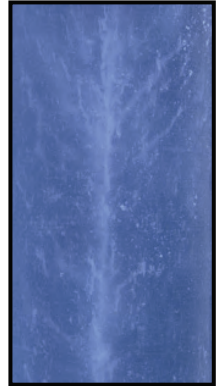


Gully et al. Proc. Roy. Soc. A 2015

brine channels



D. Cole



K. Golden

millimeters

centimeters

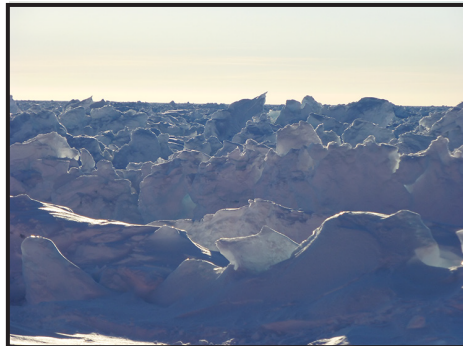
mesoscale

Arctic melt ponds



K. Frey

Antarctic pressure ridges



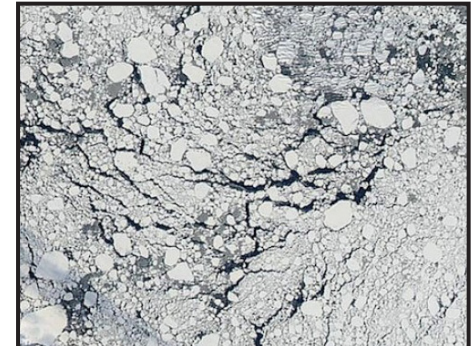
K. Golden

sea ice floes



J. Weller

sea ice pack



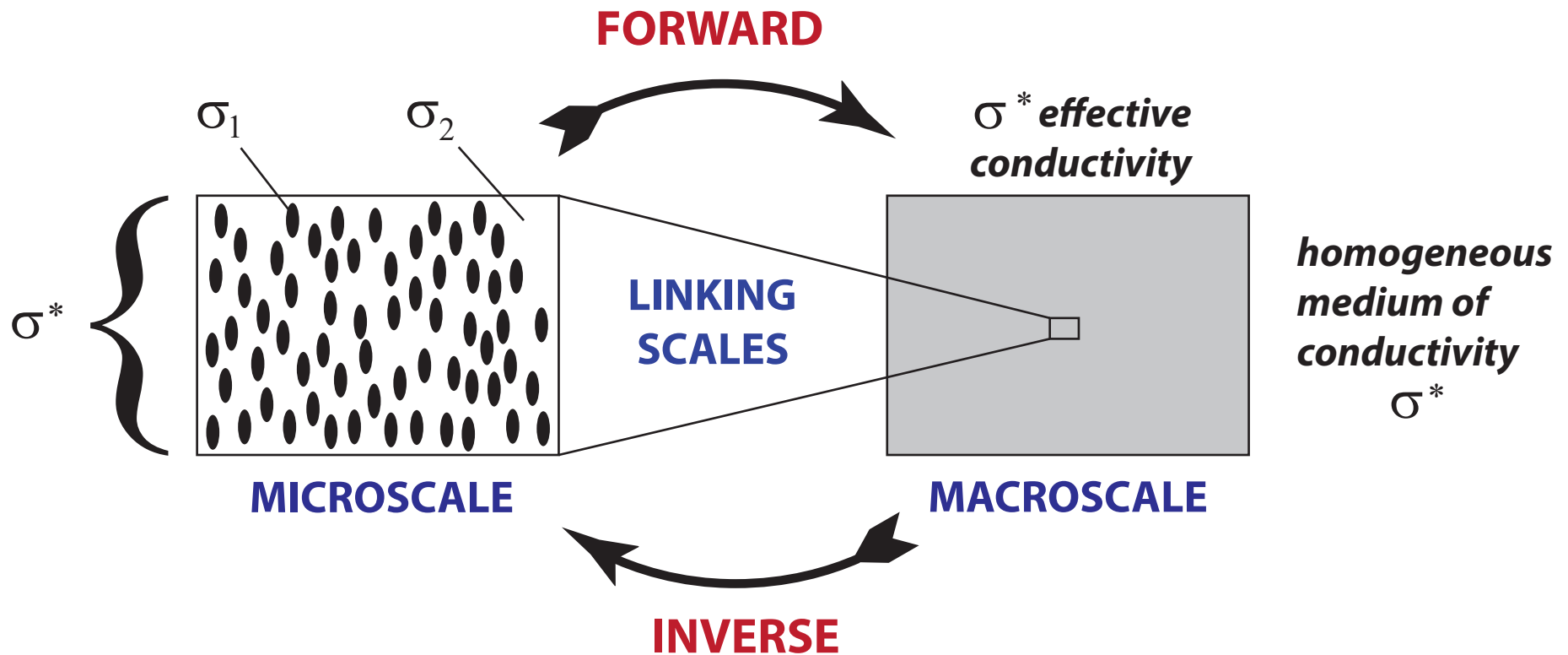
NASA

meters

kilometers

macroscale

HOMOGENIZATION for Composite Materials



Maxwell 1873 : effective conductivity of a dilute suspension of spheres

Einstein 1906 : effective viscosity of a dilute suspension of rigid spheres in a fluid

*Wiener 1912 : arithmetic and harmonic mean **bounds** on effective conductivity*

*Hashin and Shtrikman 1962 : variational **bounds** on effective conductivity*

widespread use of composites in late 20th century due in large part to advances in mathematically predicting their effective properties

What is this talk about?

A tour of recent results on multiscale modelling of fluid and composite behaviour in the sea ice system, with a focus on novel mathematics.

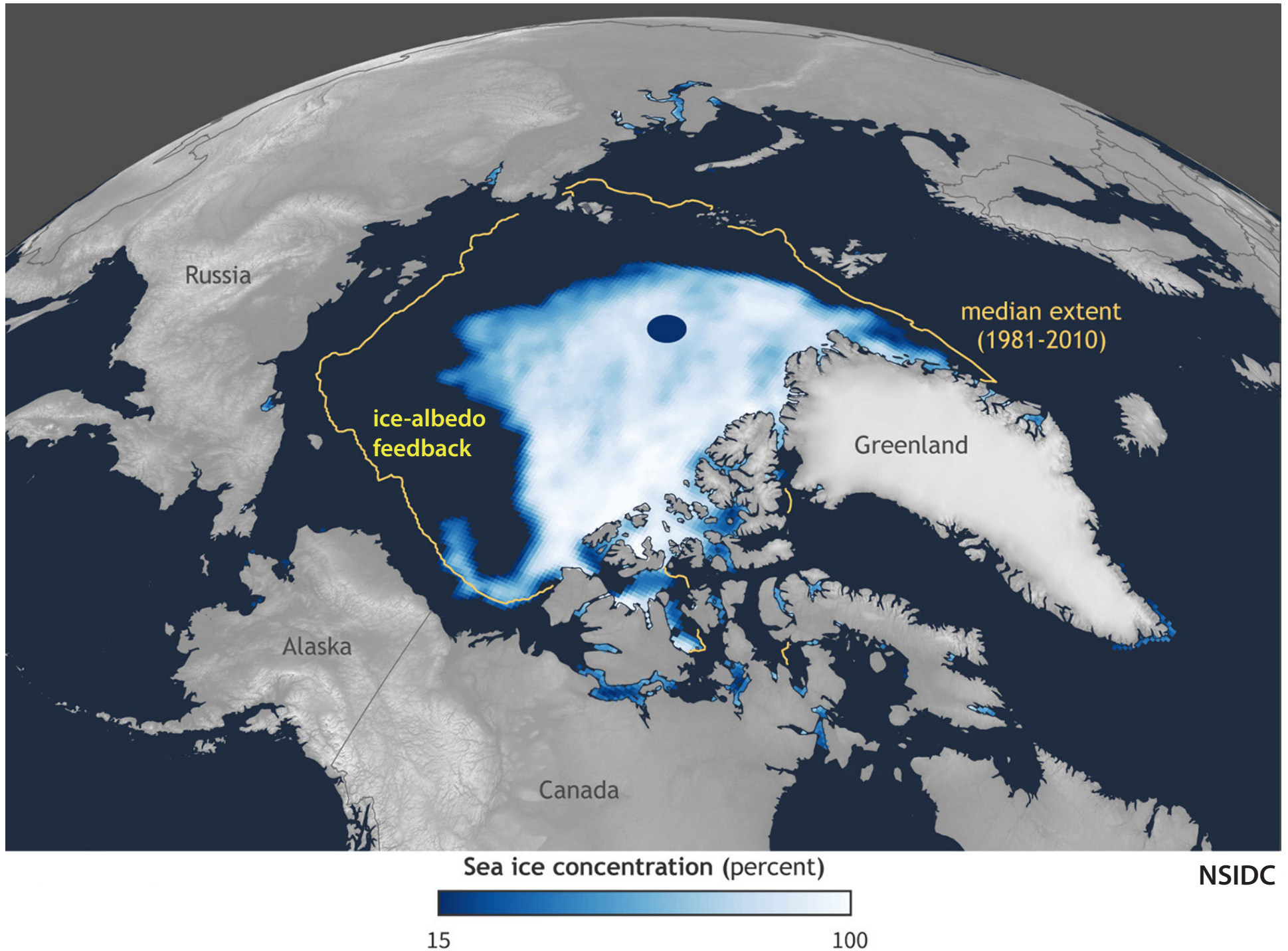
microscale

mesoscale

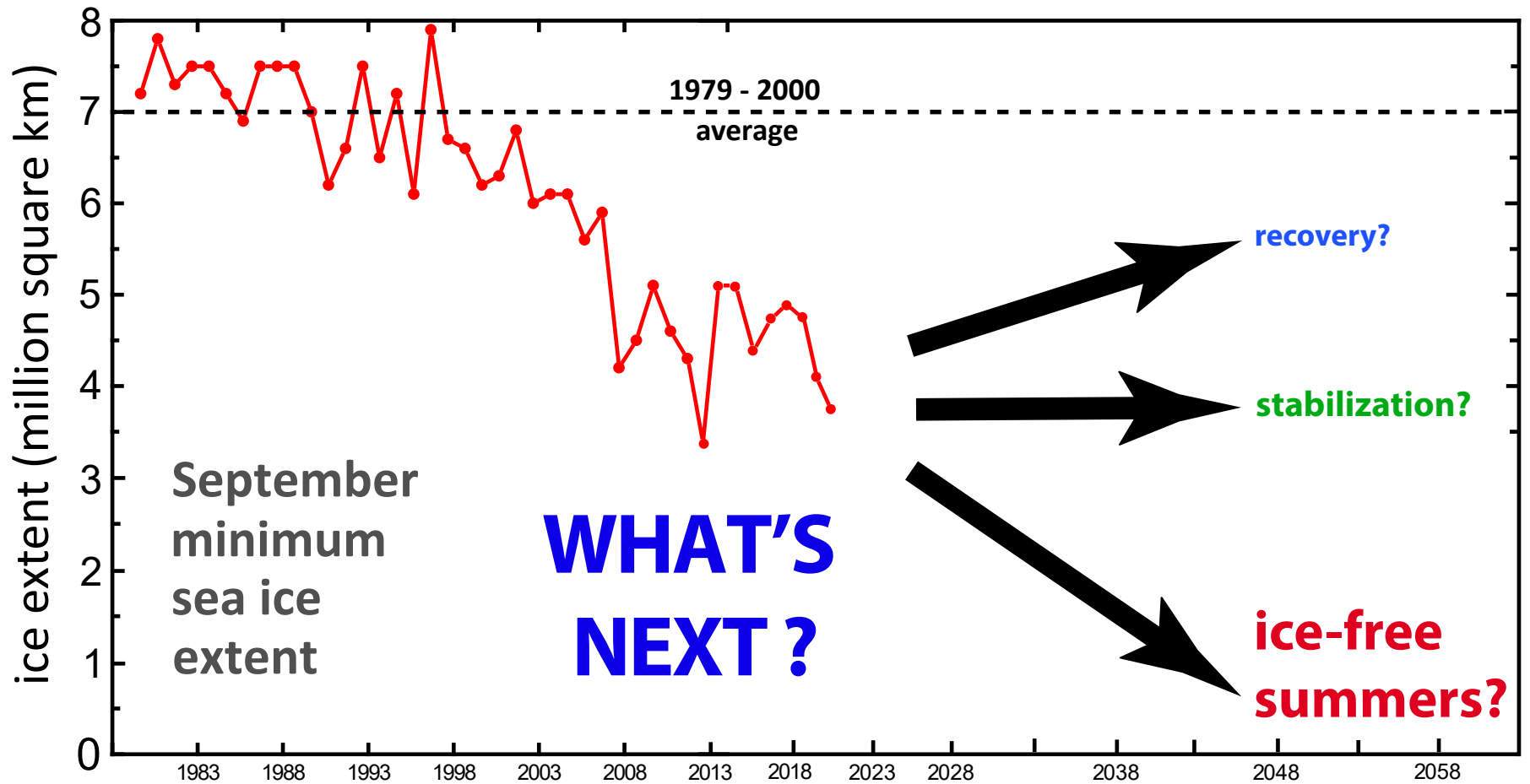
macroscale

Arctic sea ice extent

September 15, 2020

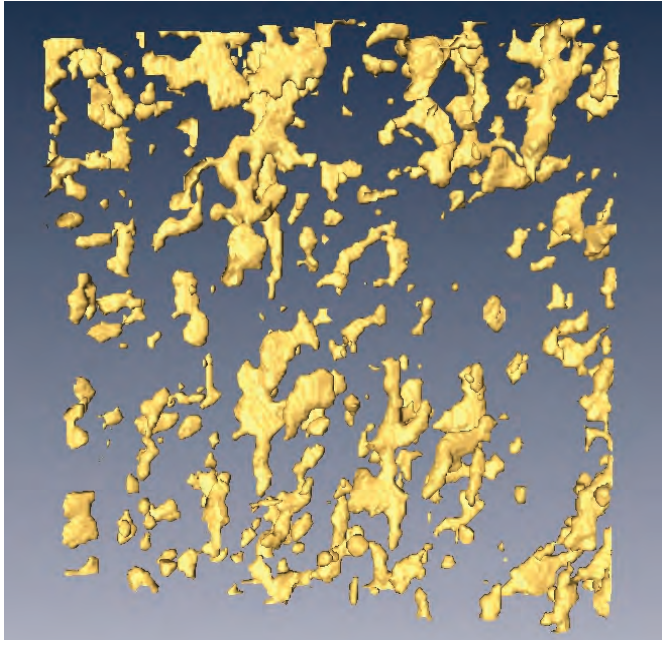


Predicting what may come next requires lots of math modeling.

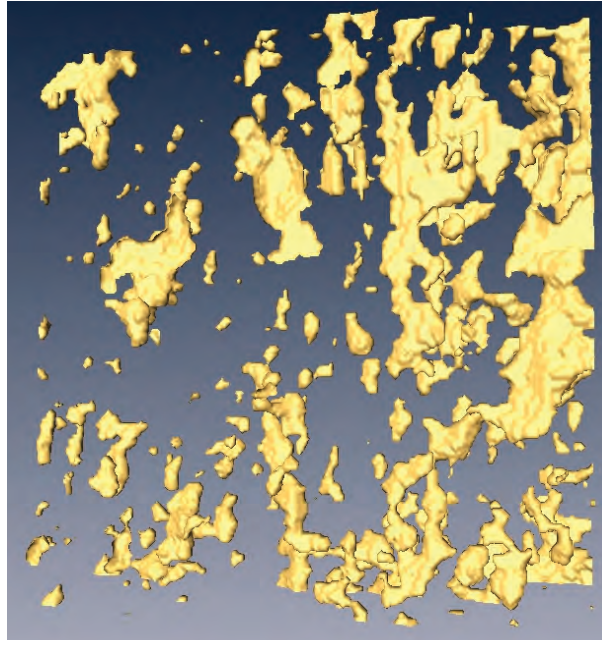


microscale

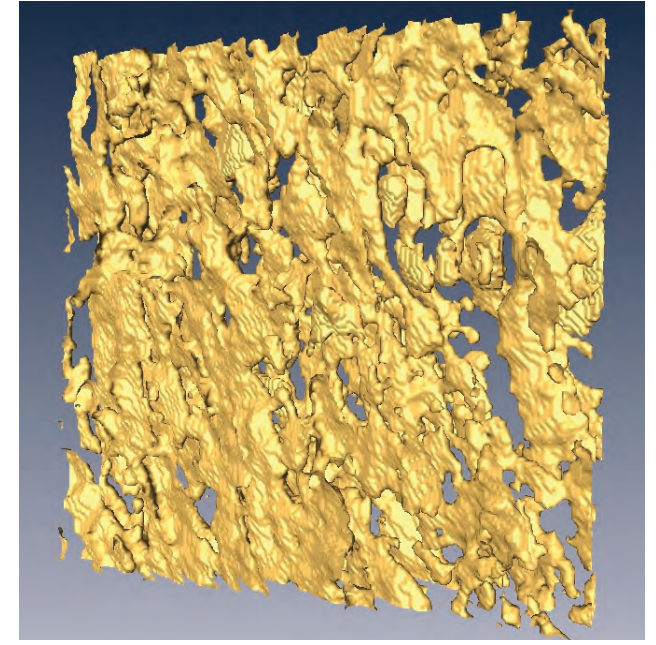
brine volume fraction and **connectivity** increase with temperature



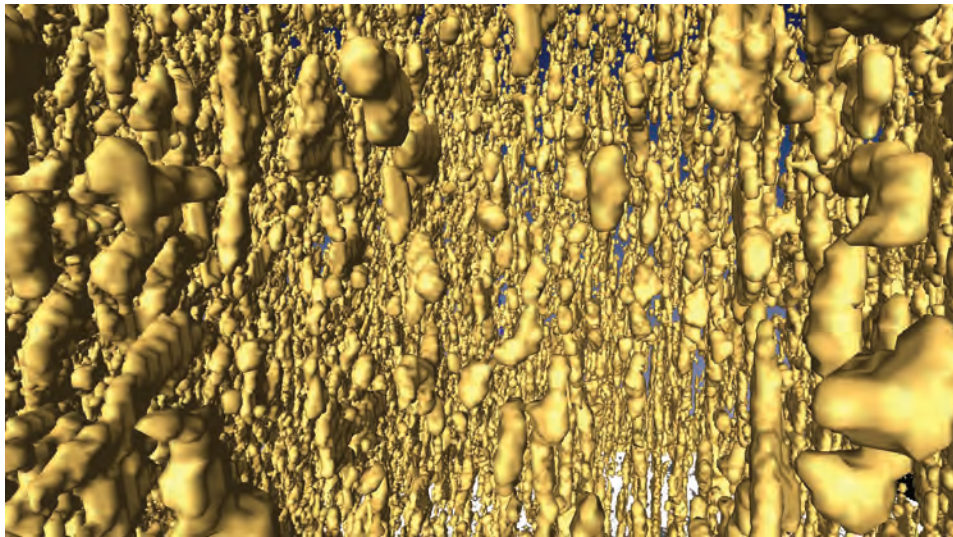
$T = -15\text{ }^{\circ}\text{C}$, $\phi = 0.033$



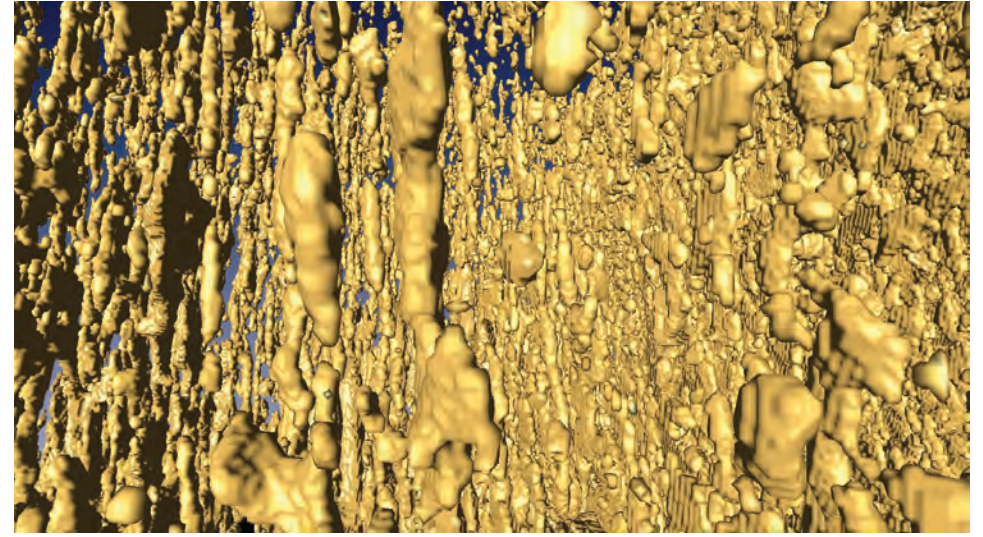
$T = -6\text{ }^{\circ}\text{C}$, $\phi = 0.075$



$T = -3\text{ }^{\circ}\text{C}$, $\phi = 0.143$



$T = -8\text{ }^{\circ}\text{C}$, $\phi = 0.057$



$T = -4\text{ }^{\circ}\text{C}$, $\phi = 0.113$

X-ray tomography for brine in sea ice

Golden et al., *Geophysical Research Letters*, 2007

Critical behavior of fluid transport in sea ice

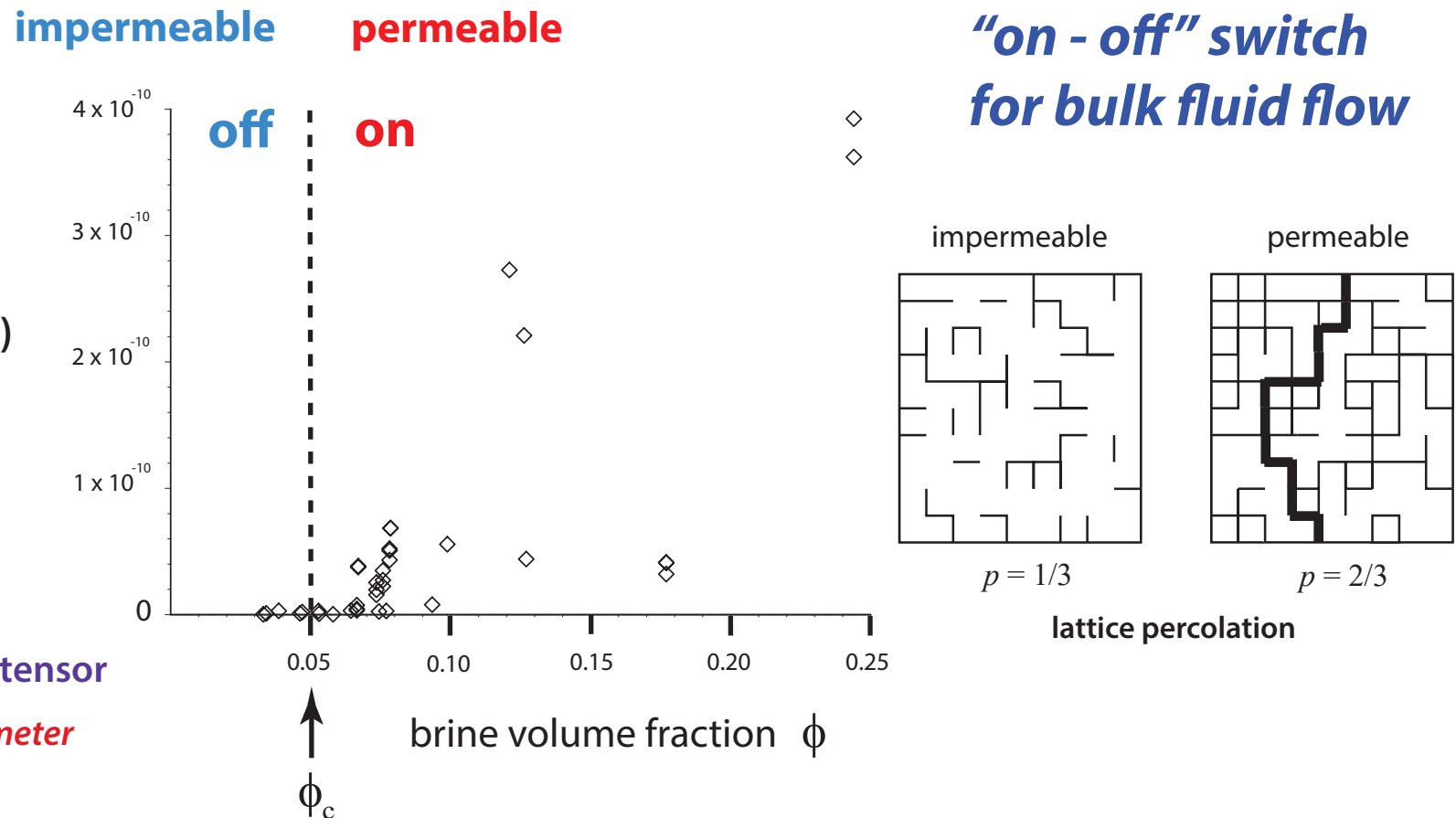
Arctic field data

vertical fluid permeability k (m^2)

Darcy's Law

$$\mathbf{v} = -\frac{\mathbf{k}}{\eta} \nabla p$$

\mathbf{k} = fluid permeability tensor
homogenized parameter



PERCOLATION THRESHOLD $\phi_c \approx 5\% \longleftrightarrow T_c \approx -5^\circ \text{C}, S \approx 5 \text{ ppt}$

RULE OF FIVES

Golden, Ackley, Lytle *Science* 1998

Golden, Eicken, Heaton, Miner, Pringle, Zhu *GRL* 2007

Pringle, Miner, Eicken, Golden *J. Geophys. Res.* 2009



sea ice algal communities

D. Thomas 2004

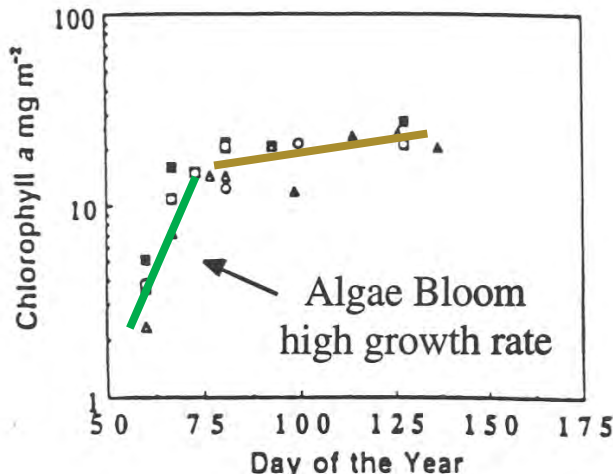
nutrient replenishment
controlled by ice permeability

biological activity turns on
or off according to
rule of fives

Golden, Ackley, Lytle Science 1998

Fritsen, Lytle, Ackley, Sullivan Science 1994

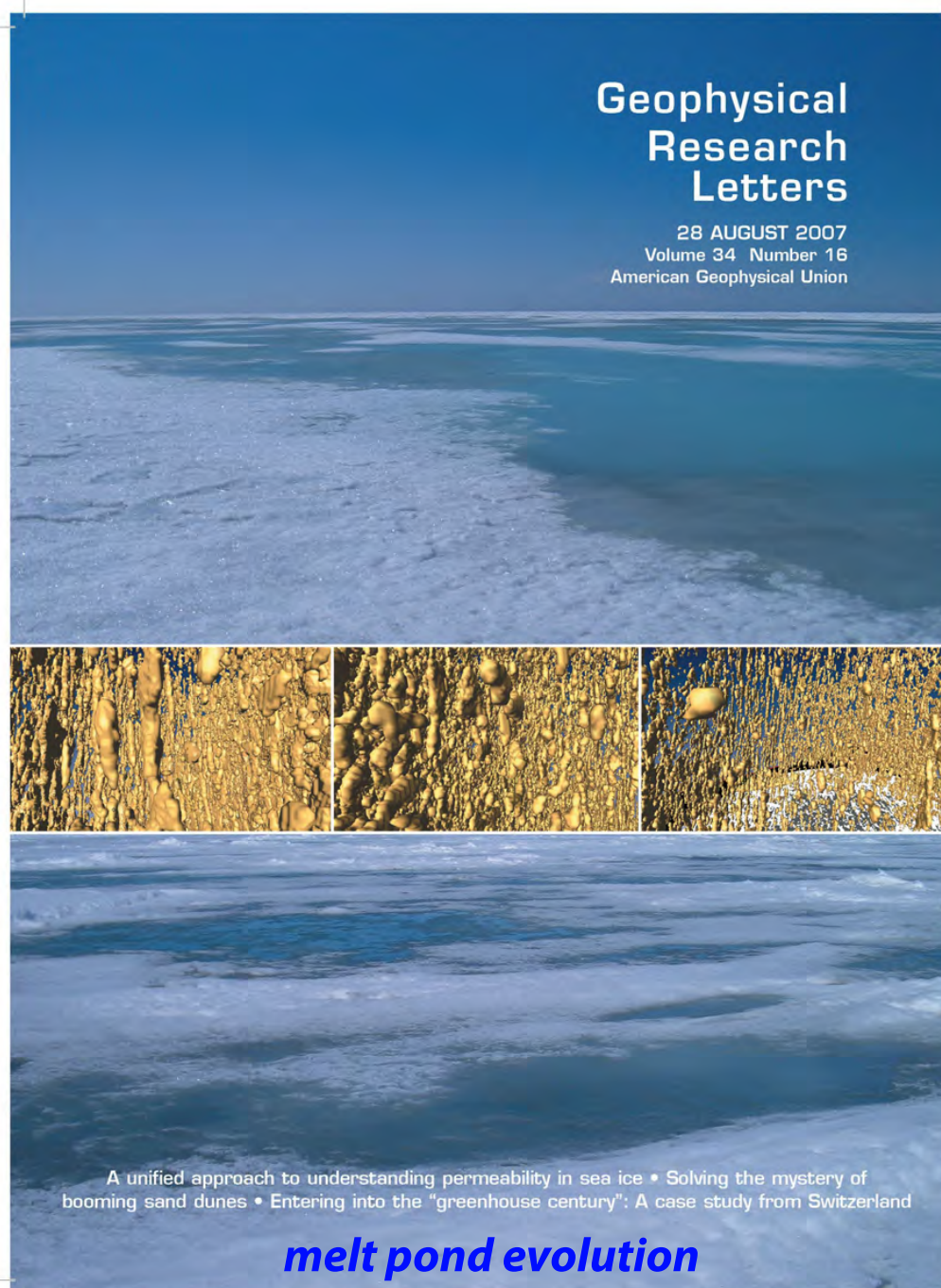
critical behavior of microbial activity



Convection-fueled algae bloom
Ice Station Weddell

Thermal evolution of permeability and microstructure in sea ice

Golden, Eicken, Heaton, Miner, Pringle, Zhu, *Geophysical Research Letters* 2007



percolation theory
for fluid permeability

$$k(\phi) = k_0 (\phi - 0.05)^2$$

critical exponent t

$$k_0 = 3 \times 10^{-8} \text{ m}^2$$

from critical path analysis
in hopping conduction

hierarchical model

rock physics

network model

rigorous bounds

X-ray tomography for
brine inclusions

confirms rule of fives

brine percolation threshold
of $\phi = 5\%$ for bulk fluid flow

Pringle, Miner, Eicken, Golden
J. Geophys. Res. 2009

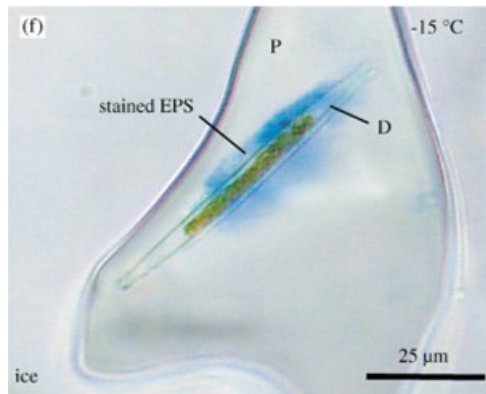
theories agree closely
with field data

microscale
governs
mesoscale
processes

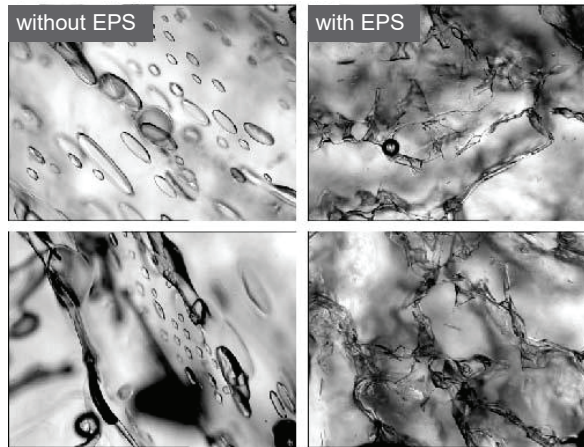
melt pond evolution

Sea ice algae secrete extracellular polymeric substances (EPS) affecting evolution of brine microstructure.

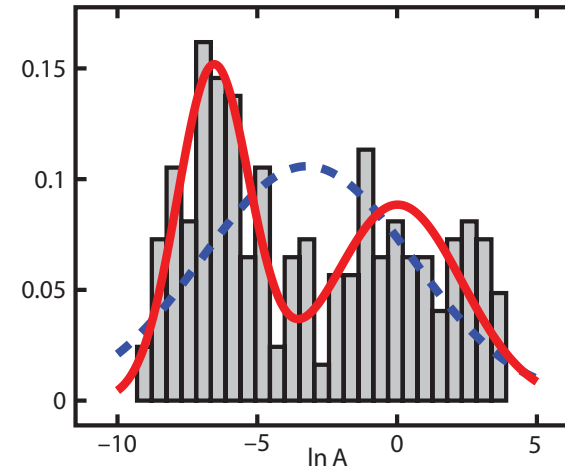
How does EPS affect fluid transport? How does the biology affect the physics?



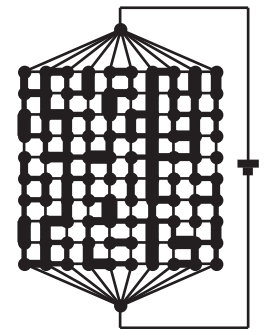
Krembs



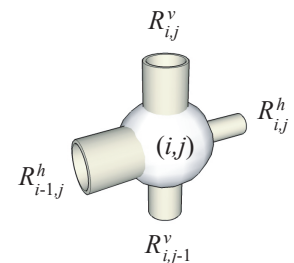
Krembs, Eicken, Deming, PNAS 2011



**RANDOM
PIPE
MODEL**



- 2D random pipe model with bimodal distribution of pipe radii
- Rigorous bound on permeability k ; results predict observed drop in k

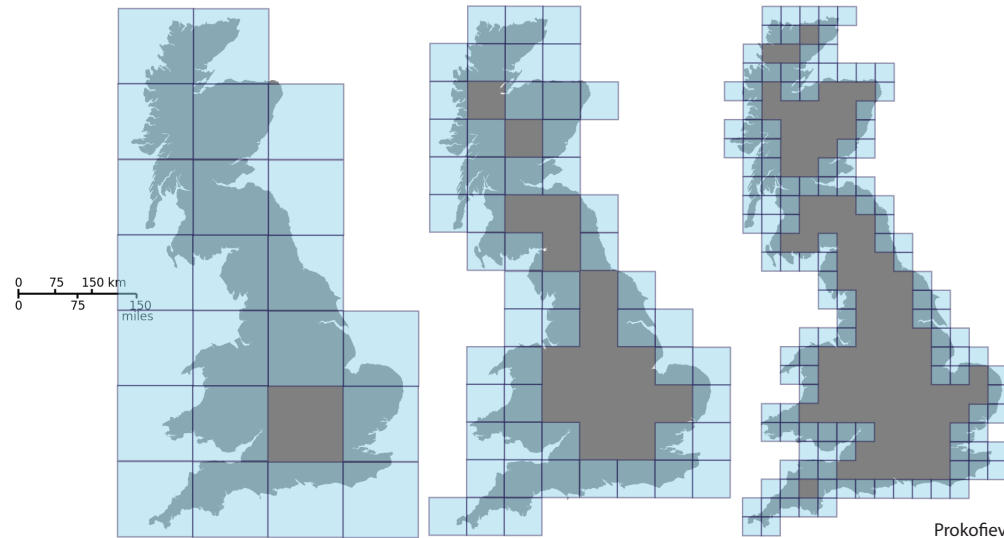


Steffen, Epshteyn, Zhu, Bowler, Deming, Golden
Multiscale Modeling and Simulation, 2018

Zhu, Jabini, Golden,
Eicken, Morris
Ann. Glac. 2006

Thermal Evolution of Brine Fractal Geometry in Sea Ice

Nash Ward, Daniel Hallman, Benjamin Murphy, Jody Reimer,
Marc Oggier, Megan O'Sadnick, Elena Cherkaev and Kenneth Golden, 2022

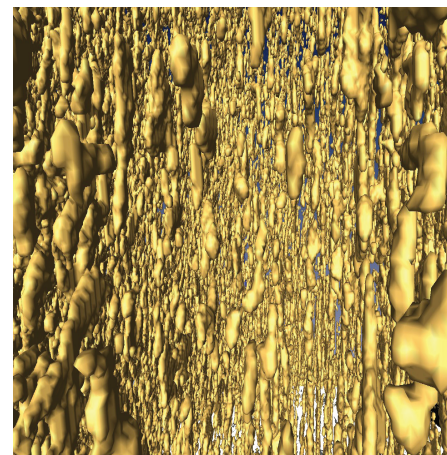


fractal dimension of the
British coastline by
box counting

$T = -12^{\circ} \text{C}$, $\phi = 0.033$



$T = -8^{\circ} \text{C}$, $\phi = 0.057$



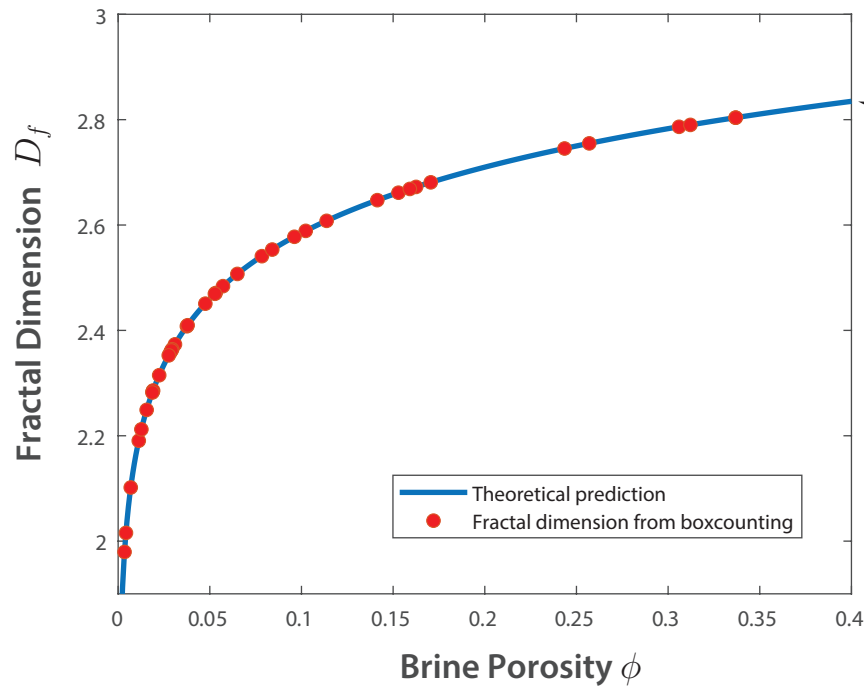
brine channels and
inclusions “look”
like fractals
(from 30 yrs ago)

X-ray computed
tomography of
brine in sea ice

columnar and granular

Golden, Eicken, et al. *GRL*, 2007

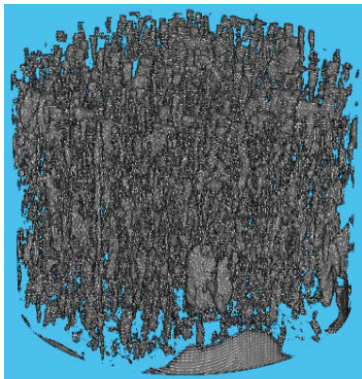
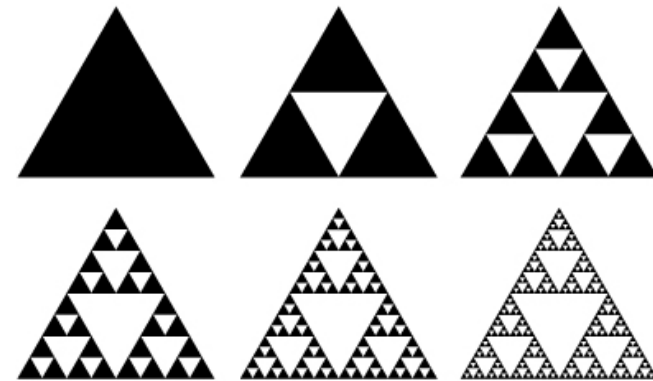
The first comprehensive, quantitative study of the fractal dimension of brine in sea ice and its strong dependence on temperature and porosity.



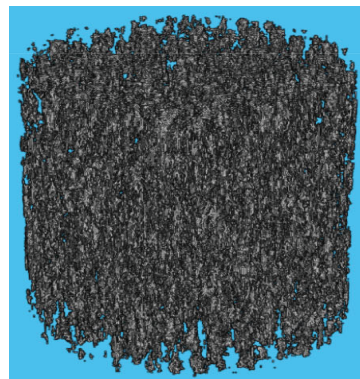
$$D_f = 3 - \frac{\ln \phi}{\ln(\lambda_{min}/\lambda_{max})}$$

The blue curve is exact for the Sierpinski gasket (an exactly self-similar geometry); discovered for sandstones - statistically self-similar porous media.

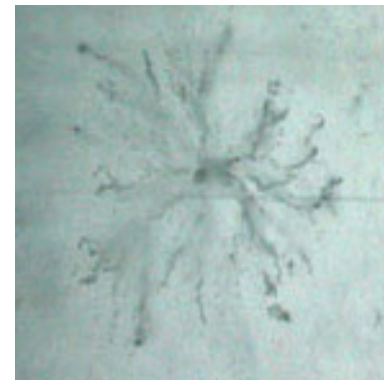
Katz and Thompson, 1985
Yu and Li, 2001



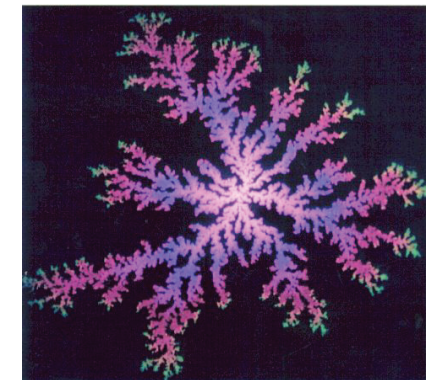
X-ray tomography



DLA model

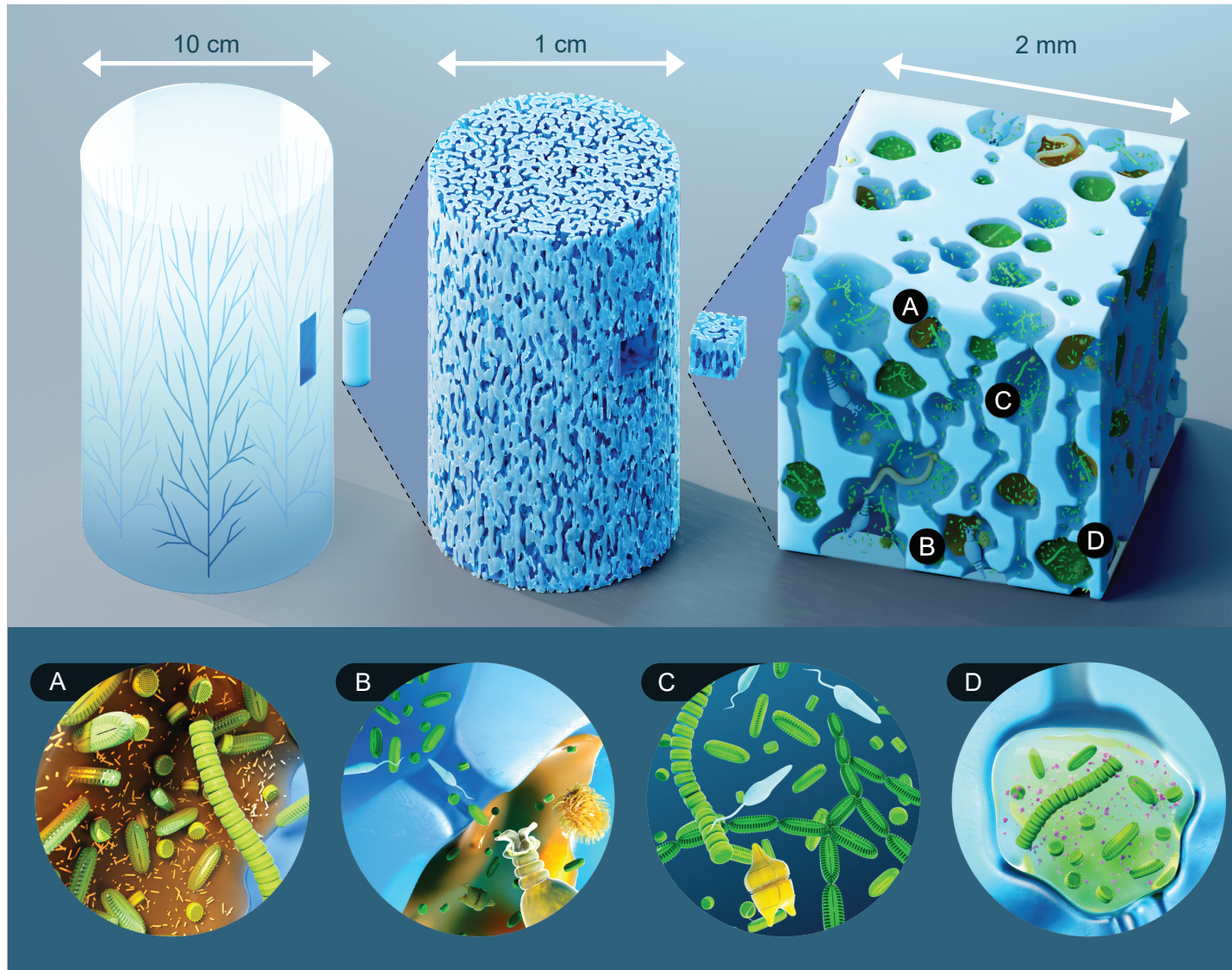


brine channel
in sea ice



diffusion limited
aggregation

Implications of brine fractal geometry on sea ice ecology and biogeochemistry



Brine inclusions are home to ice endemic organisms, e.g., bacteria, diatoms, flagellates, rotifers, nematodes.

The habitability of sea ice for these organisms is inextricably linked to its complex brine geometry.

- (A) Many sea ice organisms attach themselves to inclusion walls; inclusions with a higher fractal dimension have greater surface area for colonization.
- (B) Narrow channels prevent the passage of larger organisms, leading to refuges where smaller organisms can multiply without being grazed, as in (C).
- (D) Ice algae secrete extracellular polymeric substances (EPS) which alter inclusion geometry and may further increase the fractal dimension.

Arctic and Antarctic field experiments

*develop electromagnetic methods
of monitoring fluid transport and
microstructural transitions*

extensive measurements of fluid and
electrical transport properties of sea ice:

2007 Antarctic SIPEX

2010 Antarctic McMurdo Sound

2011 Arctic Barrow AK

2012 Arctic Barrow AK

2012 Antarctic SIPEX II

2013 Arctic Barrow AK

2014 Arctic Chukchi Sea



Notices

of the American Mathematical Society

May 2009

Volume 56, Number 5

Climate Change and
the Mathematics of
Transport in Sea Ice

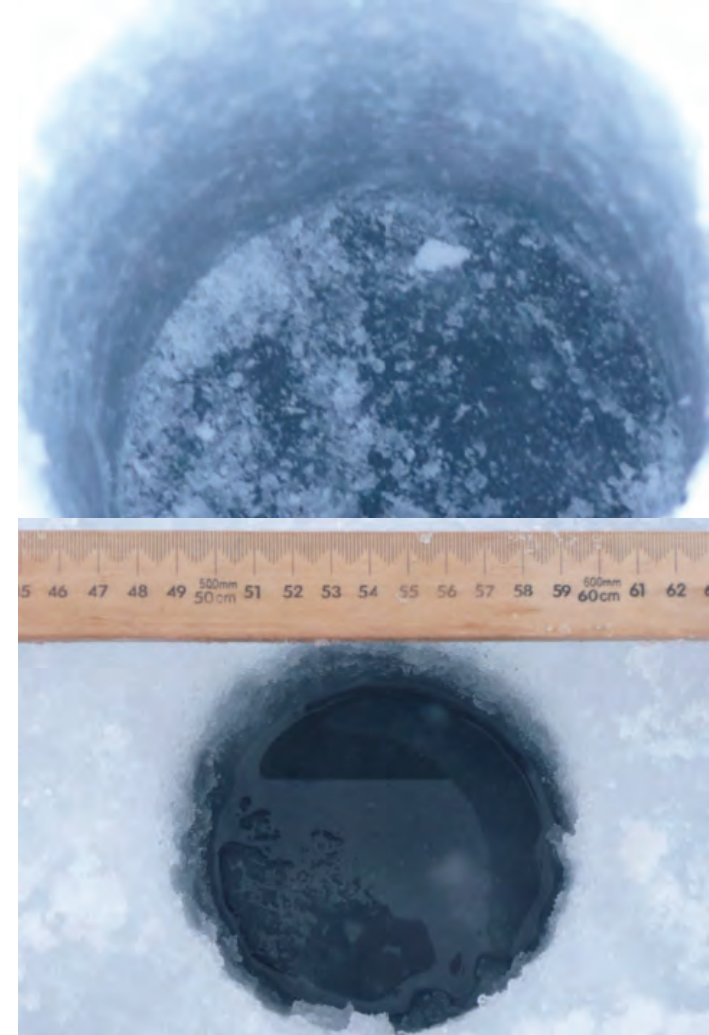
page 562

Mathematics and the
Internet: A Source of
Enormous Confusion
and Great Potential

page 586

photo by Jan Lieser

Real analysis in polar coordinates (see page 613)



***measuring
fluid permeability
of Antarctic sea ice***

SIPEX 2007



Remote sensing of sea ice



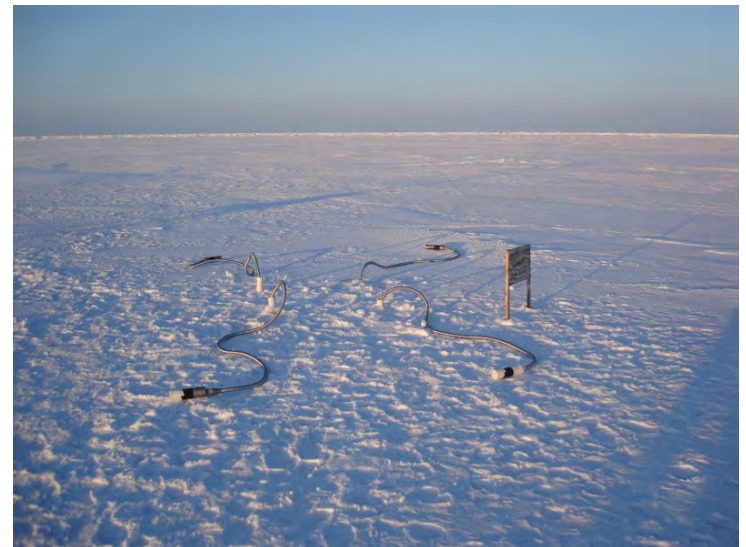
sea ice thickness
ice concentration

INVERSE PROBLEM

Recover sea ice
properties from
electromagnetic
(EM) data

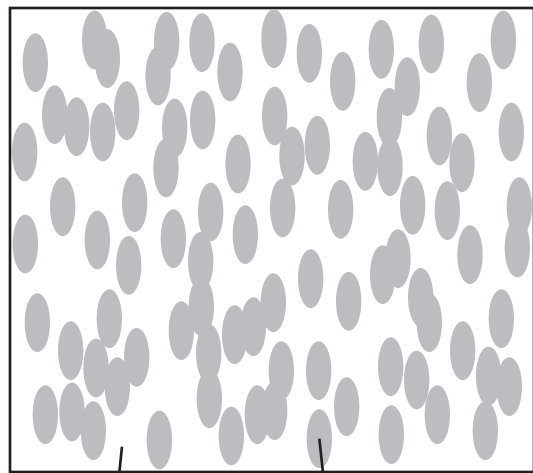
$$\epsilon^*$$

effective complex permittivity
(dielectric constant, conductivity)



brine volume fraction
brine inclusion connectivity

Effective complex permittivity of a two phase composite in the quasistatic (long wavelength) limit



ϵ_1

ϵ_2



ϵ^*

$$D = \epsilon E$$

$$\nabla \cdot D = 0$$

$$\nabla \times E = 0$$

$$\langle D \rangle = \epsilon^* \langle E \rangle$$

p_1, p_2 = volume fractions of
the components

$$\epsilon^* = \epsilon^* \left(\frac{\epsilon_1}{\epsilon_2}, \text{ composite geometry} \right)$$

**What are the effective propagation characteristics
of an EM wave (radar, microwaves) in the medium?**

Analytic Continuation Method for Homogenization

Bergman (1978), Milton (1979), Golden and Papanicolaou (1983), Theory of Composites, Milton (2002)

Stieltjes integral representation for homogenized parameter

separates geometry from parameters

$$F(s) = 1 - \frac{\epsilon^*}{\epsilon_2} = \int_0^1 \frac{d\mu(z)}{s - z}$$

← geometry

← material parameters

$$s = \frac{1}{1 - \epsilon_1 / \epsilon_2}$$

μ

- spectral measure of self adjoint operator $\Gamma\chi$
- mass = p_1
- higher moments depend on n -point correlations

$$\Gamma = \nabla(-\Delta)^{-1}\nabla.$$

χ = characteristic function of the brine phase

$$E = s (s + \Gamma\chi)^{-1} e_k$$

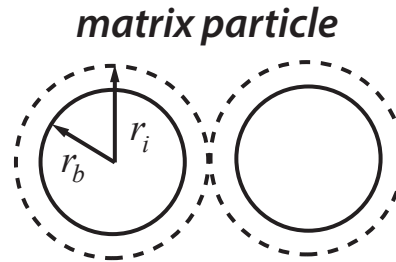
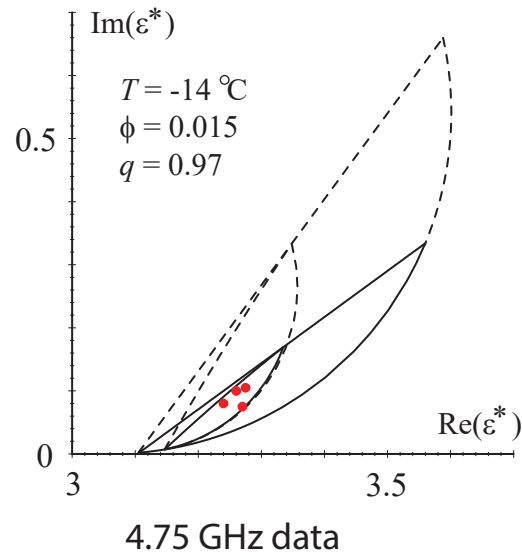
$\Gamma\chi$: microscale \rightarrow macroscale

$\Gamma\chi$ *links scales*

This representation distills the complexities of mixture geometry into the spectral properties of an operator like the Hamiltonian in physics.

forward and inverse bounds on the complex permittivity of sea ice

forward bounds

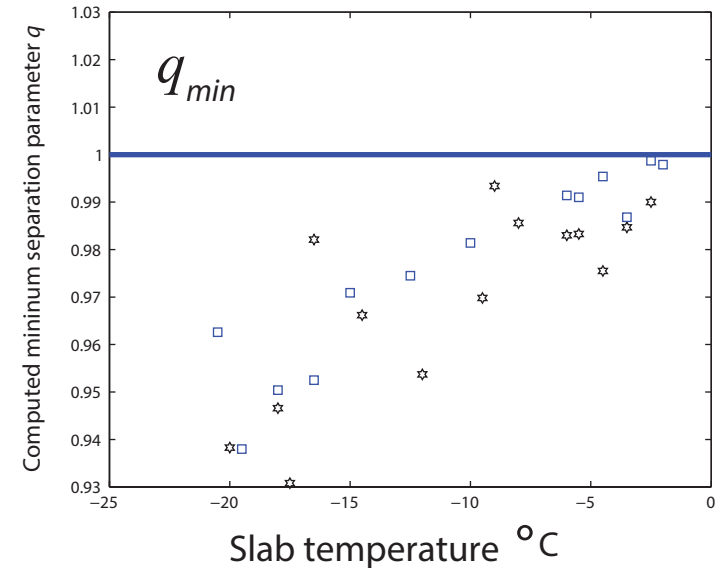


$$q = r_b / r_i$$

$$0 < q < 1$$

Golden 1995, 1997

inverse bounds



Inverse Homogenization

Cherkaev and Golden (1998), Day and Thorpe (1999), Cherkaev (2001), McPhedran, McKenzie, Milton (1982), *Theory of Composites*, Milton (2002)

ϵ^* \longrightarrow composite geometry
(spectral measure μ)

inverse bounds and recovery of brine porosity

**Gully, Backstrom, Eicken, Golden
Physica B, 2007**

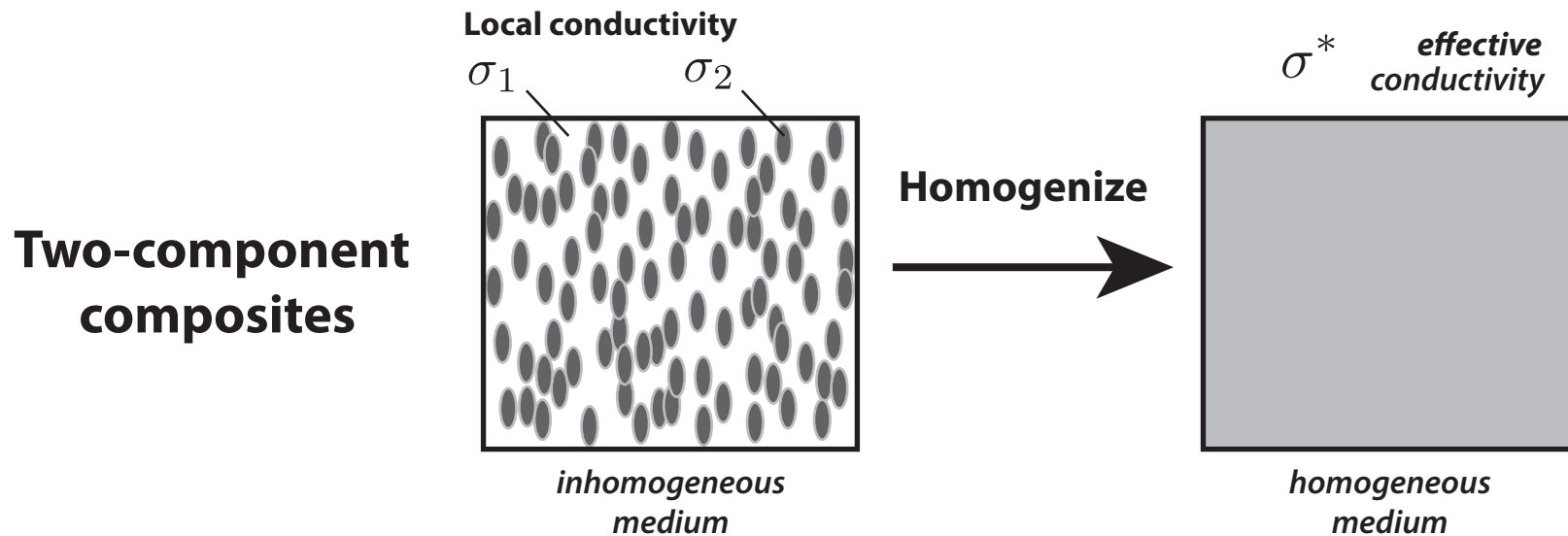
inversion for brine inclusion separations in sea ice from measurements of effective complex permittivity ϵ^*

rigorous inverse bound on spectral gap

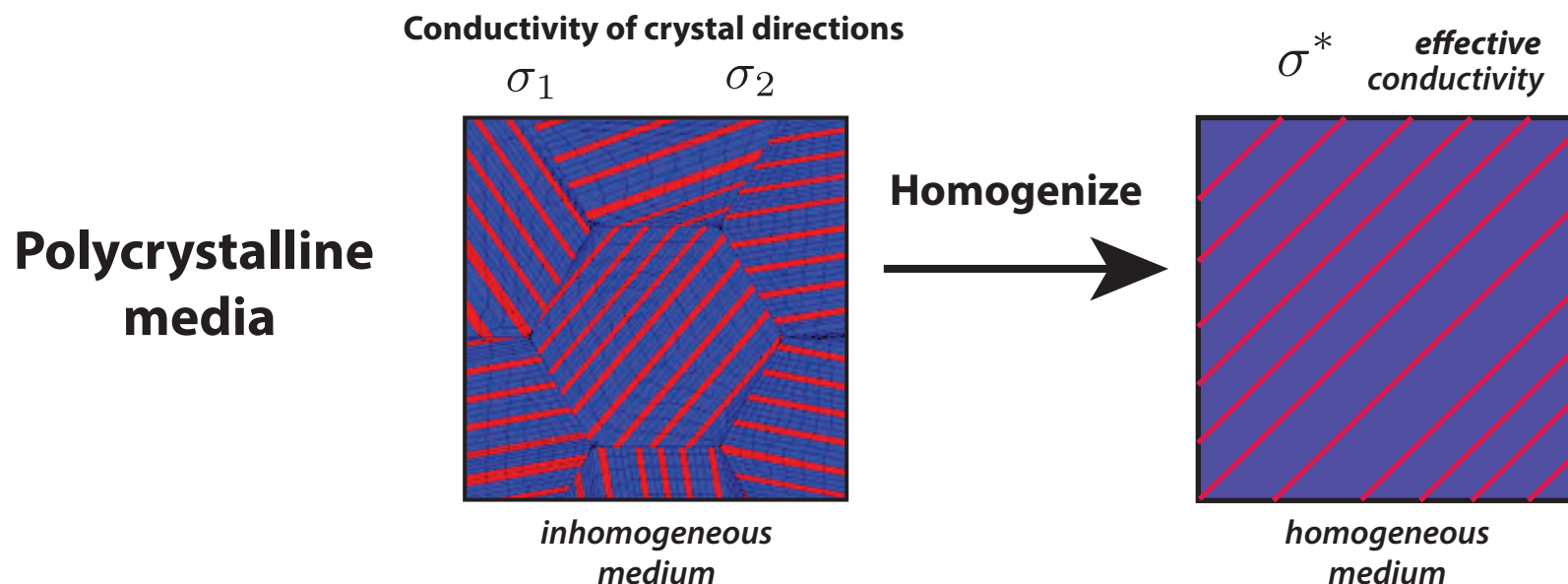
construct algebraic curves which bound admissible region in (p, q) -space

**Orum, Cherkaev, Golden
Proc. Roy. Soc. A, 2012**

Homogenization for polycrystalline materials



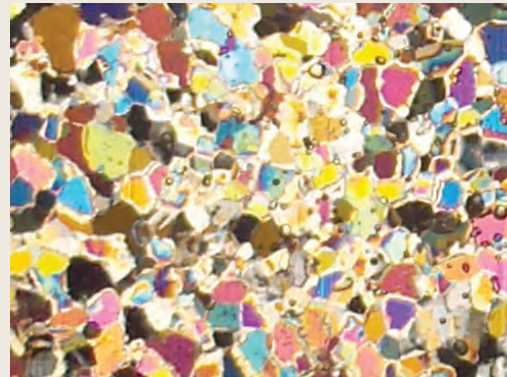
Find the homogeneous medium which behaves macroscopically the same as the inhomogeneous medium



Bounds on the complex permittivity of polycrystalline materials by analytic continuation

Adam Gully, Joyce Lin,
Elena Cherkaev, Ken Golden

- **Stieltjes integral representation for effective complex permittivity**
Milton (1981, 2002), Barabash and Stroud (1999), ...
- **Forward and inverse bounds**
orientation statistics
- **Applied to sea ice using two-scale homogenization**
- **Inverse bounds give method for distinguishing ice types using remote sensing techniques**



PROCEEDINGS A

350 YEARS
OF SCIENTIFIC
PUBLISHING

An invited review
commemorating 350 years
of scientific publishing at the
Royal Society

A method to distinguish
between different types
of sea ice using remote
sensing techniques

A computer model to
determine how a human
should walk so as to expend
the least energy



THE
ROYAL
SOCIETY
PUBLISHING

higher threshold for fluid flow in granular sea ice

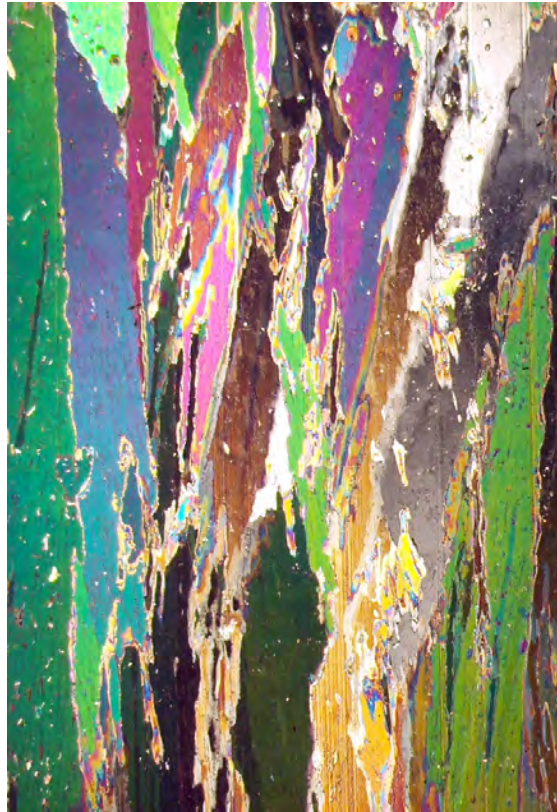
microscale details impact “mesoscale” processes

nutrient fluxes for microbes
melt pond drainage
snow-ice formation

columnar

granular

5%



10%



Golden, Sampson, Gully, Lubbers, Tison 2022

electromagnetically distinguishing ice types
Kitsel Lusted, Elena Cherkaev, Ken Golden

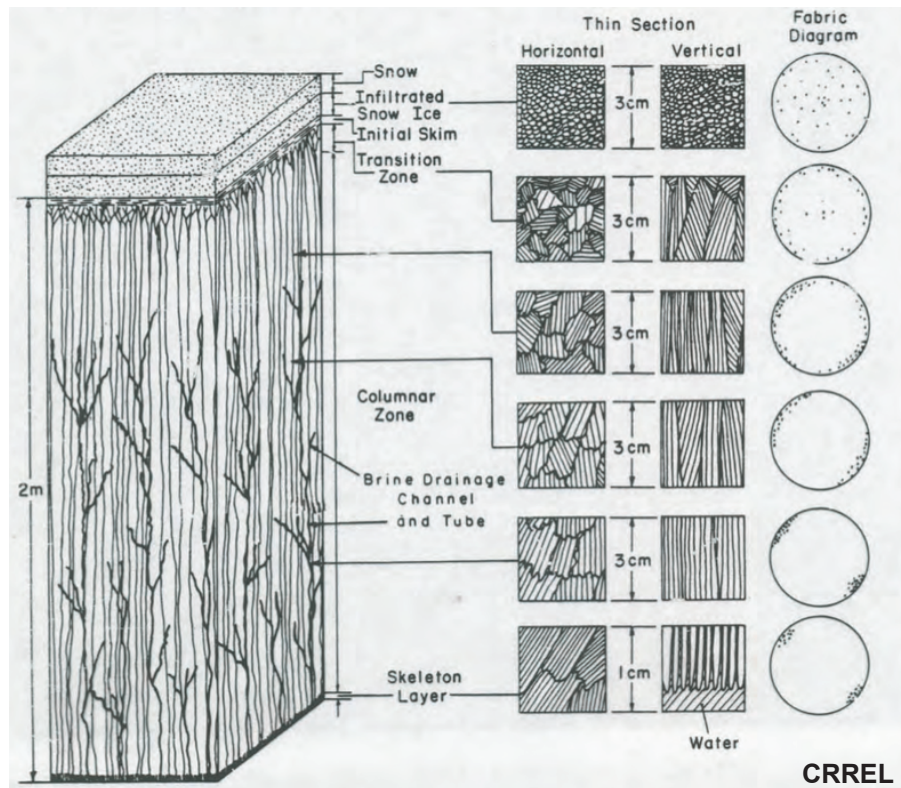
Rigorous bounds on the complex permittivity tensor of sea ice with polycrystalline anisotropy in the horizontal plane

Kenzie McLean, Elena Cherkaev, Ken Golden 2022

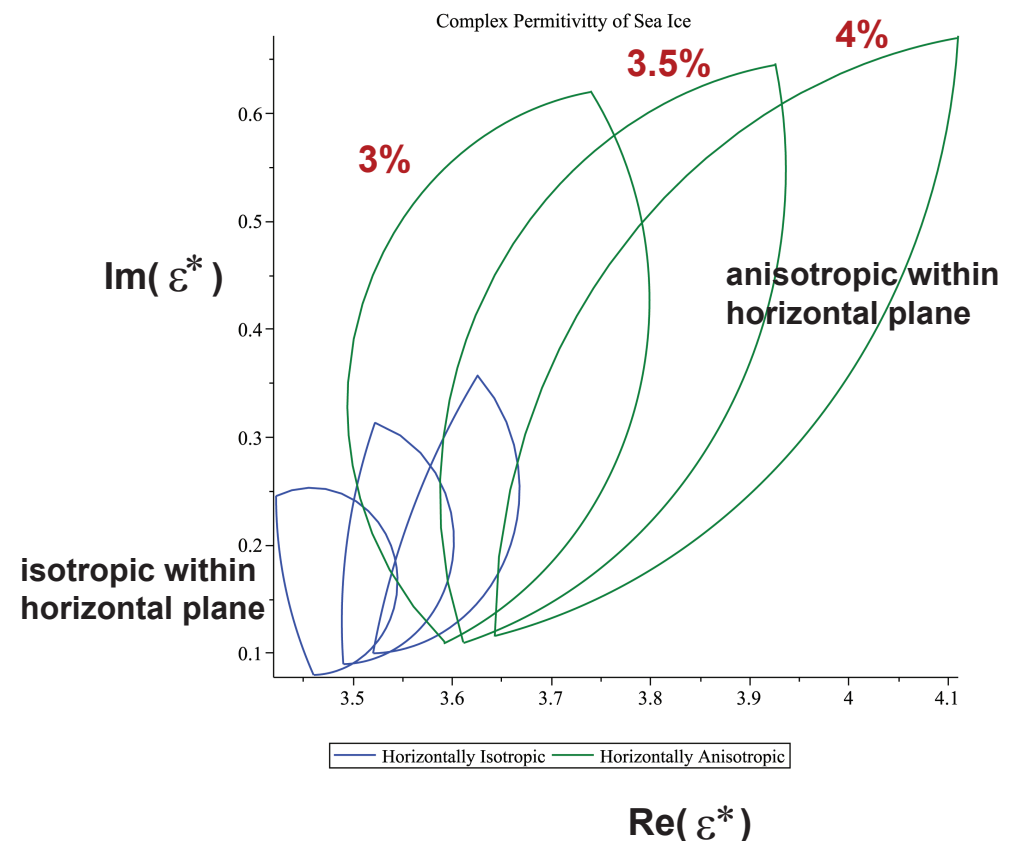
motivated by **Weeks and Gow, *JGR* 1979: c-axis alignment in Arctic fast ice off Barrow**

Golden and Ackley, *JGR* 1981: radar propagation model in aligned sea ice

input: orientation statistics



output: bounds



direct calculation of spectral measures

Murphy, Hohenegger, Cherkaev, Golden, *Comm. Math. Sci.* 2015

- depends only on the composite geometry
- discretization of microstructural image gives binary network
- fundamental operator becomes a random matrix
- spectral measure computed from eigenvalues and eigenvectors

**once we have the spectral measure μ it can be used in
Stieltjes integrals for other transport coefficients:**

***electrical and thermal conductivity, complex permittivity,
magnetic permeability, diffusion, fluid flow properties***

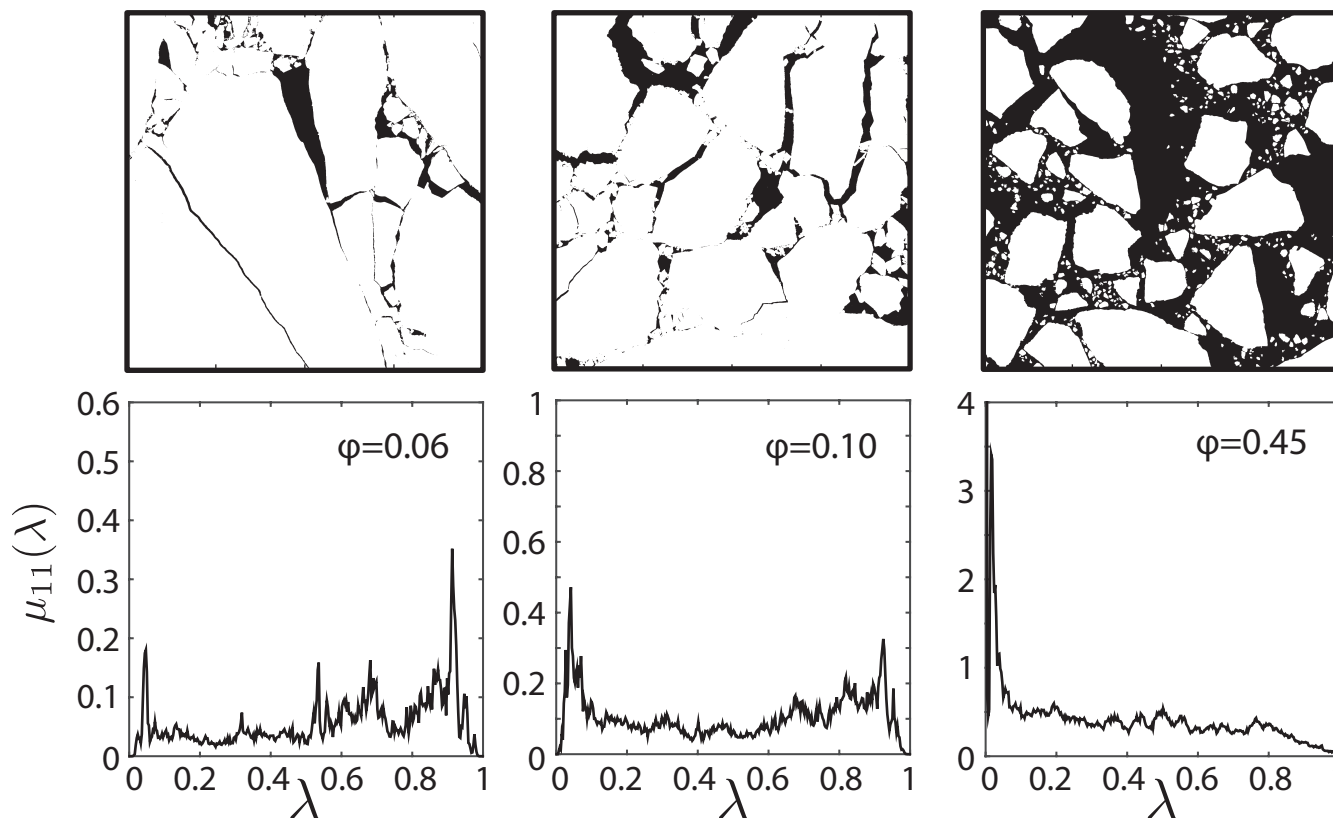
earlier studies of spectral measures

Day and Thorpe 1996

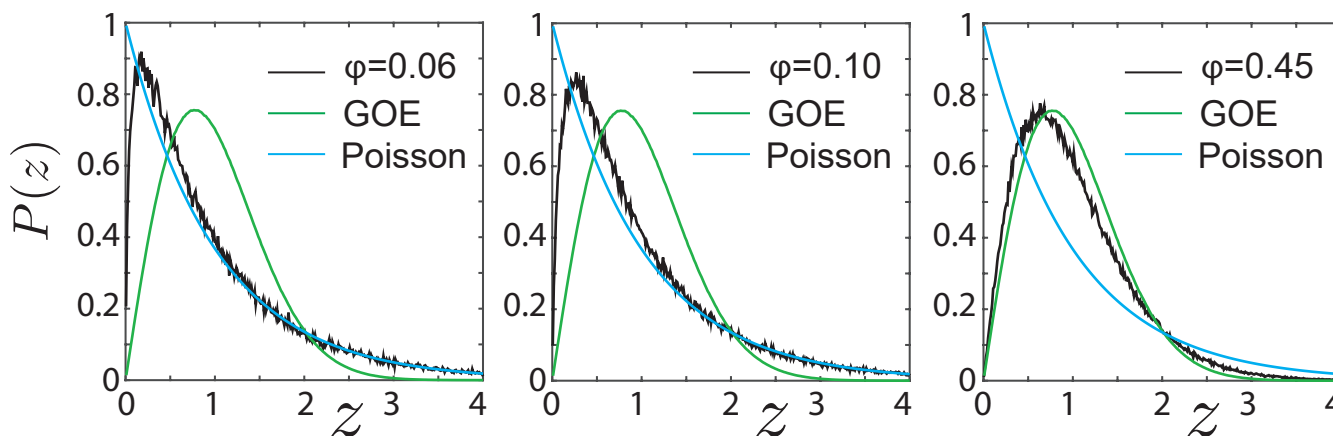
Helsing, McPhedran, Milton 2011

Spectral computations for sea ice floe configurations

spectral
measures



eigenvalue
spacing
distributions



uncorrelated



level repulsion

UNIVERSAL
Wigner-Dyson
distribution

Eigenvalue Statistics of Random Matrix Theory

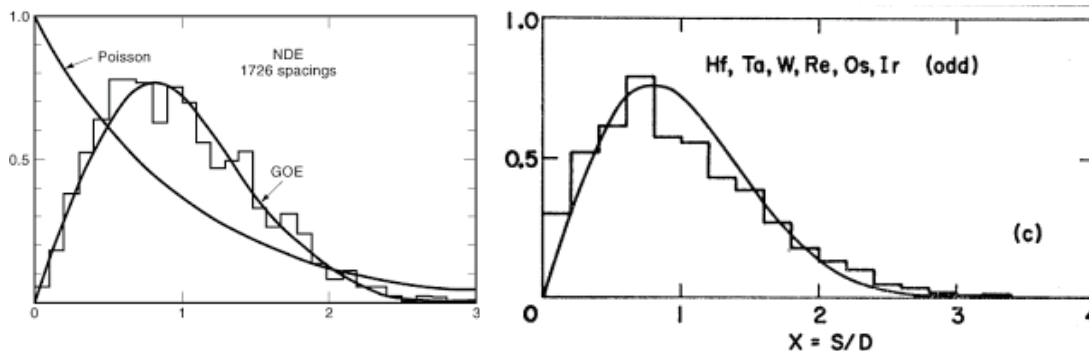
Wigner (1951) and Dyson (1953) first used random matrix theory (RMT) to describe quantized energy levels of heavy atomic nuclei.

$[N]_{ij} \sim N(0,1), \quad A = (N + N^T)/2 \quad \text{Gaussian orthogonal ensemble (GOE)}$

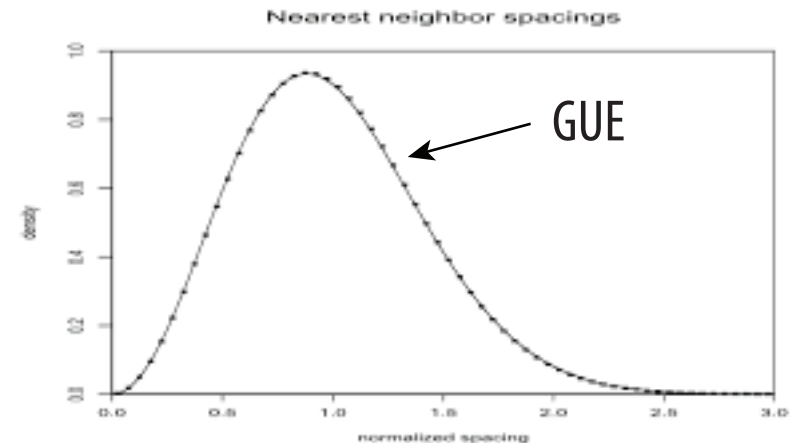
$[N]_{ij} \sim N(0,1) + iN(0,1), \quad A = (N + N^\dagger)/2 \quad \text{Gaussian unitary ensemble (GUE)}$

Short range and long range correlations of eigenvalues are measured by various eigenvalue statistics.

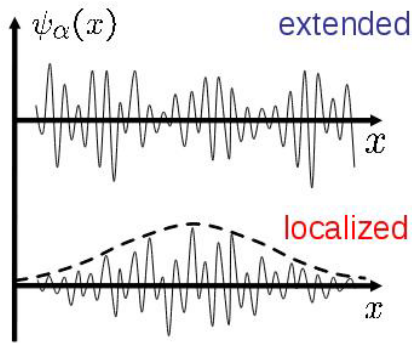
Spacing distributions of energy levels for heavy atomic nuclei



Spacing distributions of the first billion zeros of the Riemann zeta function



Universal eigenvalue statistics arise in a broad range of “unrelated” problems!



electronic transport in semiconductors

metal / insulator transition

localization

Anderson 1958
Mott 1949
Shklovshii et al 1993
Evangelou 1992

**Anderson transition in wave physics:
 quantum, optics, acoustics, water waves, ...**

from analysis of spectral measures for brine, melt ponds, ice floes

we find percolation-driven

Anderson transition for classical transport in composites

Murphy, Cherkaev, Golden Phys. Rev. Lett. 2017

**PERCOLATION
 TRANSITION**



**universal eigenvalue statistics (GOE)
 extended states, mobility edges**

-- but with NO wave interference or scattering effects ! --

Order to disorder in quasiperiodic composites

Morison, Murphy, Cherkhev, Golden, Comm. Phys. 2022

sea ice inspired - high tech spin off

tunable quasiperiodic composites with exotic properties

(optical, electrical, thermal, ...), Anderson localization; our Moiré patterned geometries are similar to **twisted bilayer graphene**

increasing twist angle between two lattices

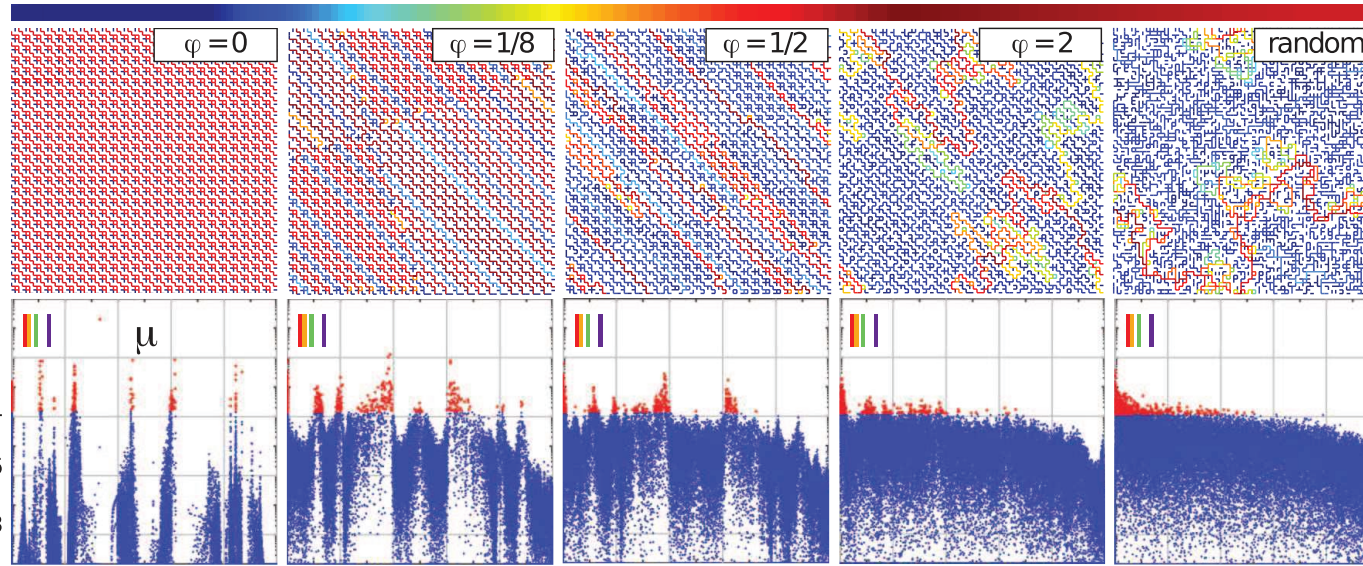
periodic

quasiperiodic

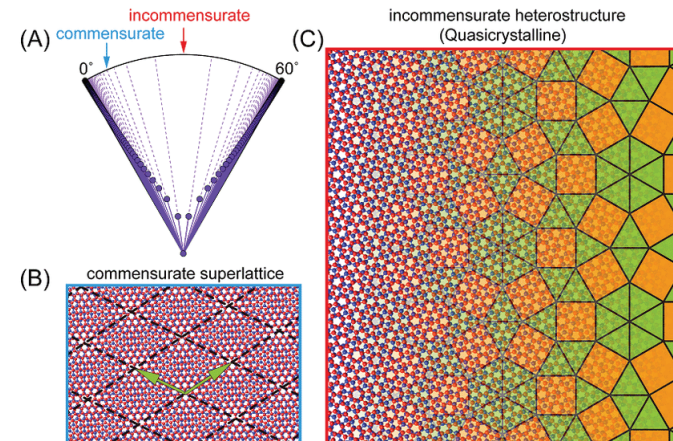
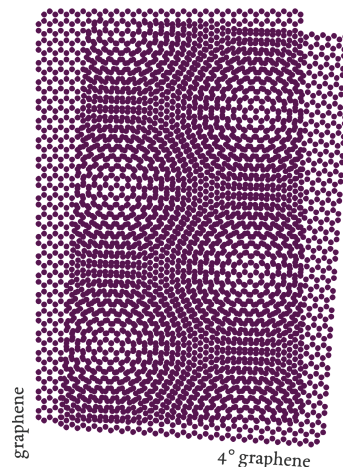
electric field strength

spectral measure

10^{-4}
 10^{-6}
 10^{-8}



twisted bilayer graphene



Yao et al., 2018

mesoscale

advection enhanced diffusion

effective diffusivity

nutrient and salt transport in sea ice
heat transport in sea ice with convection
sea ice floes in winds and ocean currents
tracers, buoys diffusing in ocean eddies
diffusion of pollutants in atmosphere

advection diffusion equation with a velocity field \vec{u}

$$\frac{\partial T}{\partial t} + \vec{u} \cdot \vec{\nabla} T = \kappa_0 \Delta T$$

$$\vec{\nabla} \cdot \vec{u} = 0$$



homogenize

$$\frac{\partial \bar{T}}{\partial t} = \kappa^* \Delta \bar{T}$$

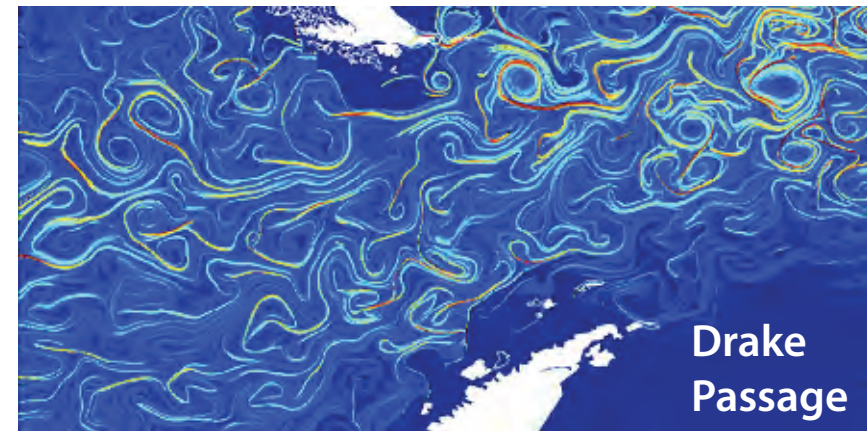
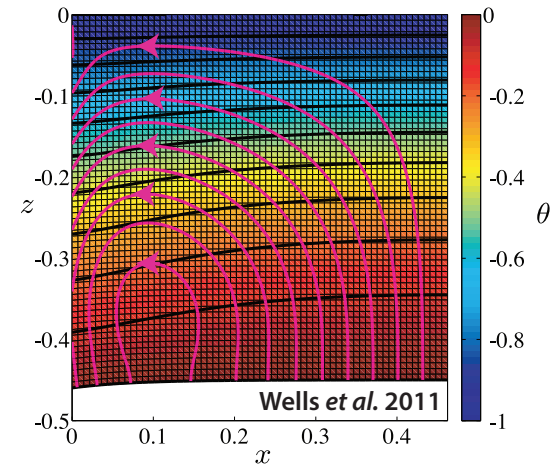
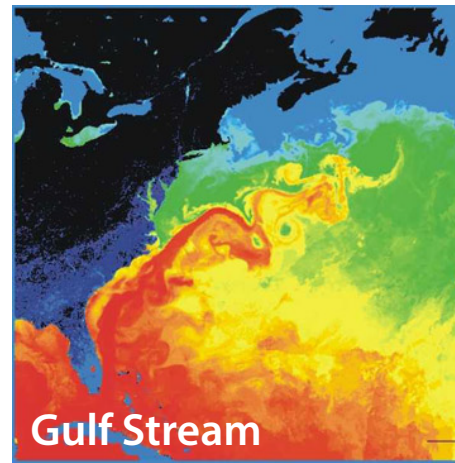
κ^* effective diffusivity

Stieltjes integral for κ^* with spectral measure

Avellaneda and Majda, PRL 89, CMP 91

Murphy, Cherkaev, Xin, Zhu, Golden, *Ann. Math. Sci. Appl.* 2017

Murphy, Cherkaev, Zhu, Xin, Golden, *J. Math. Phys.* 2020



tracers flowing through inverted sea ice blocks



Stieltjes Integral Representation for Advection Diffusion

Murphy, Cherkaev, Zhu, Xin, Golden, *J. Math. Phys.* 2020

$$\kappa^* = \kappa \left(1 + \int_{-\infty}^{\infty} \frac{d\mu(\tau)}{\kappa^2 + \tau^2} \right), \quad F(\kappa) = \int_{-\infty}^{\infty} \frac{d\mu(\tau)}{\kappa^2 + \tau^2}$$

- μ is a positive definite measure corresponding to the spectral resolution of the self-adjoint operator $i\Gamma H\Gamma$
- H = stream matrix , κ = local diffusivity
- $\Gamma := -\nabla(-\Delta)^{-1}\nabla$, Δ is the Laplace operator
- $i\Gamma H\Gamma$ is bounded for time independent flows
- $F(\kappa)$ is analytic off the spectral interval in the κ -plane

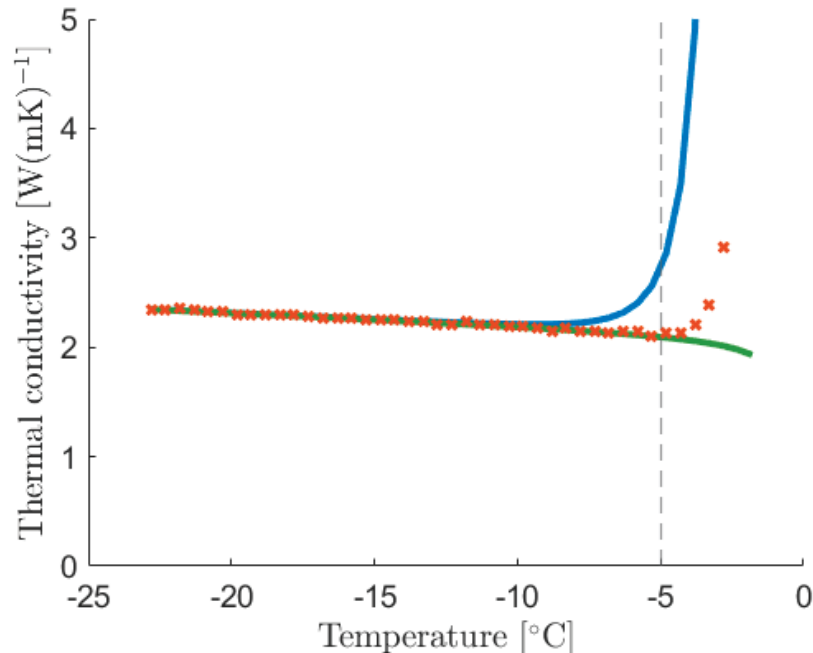
rigorous framework for numerical computations of spectral measures and effective diffusivity for model flows

new integral representations, theory of moment calculations

separation of material properties and flow field

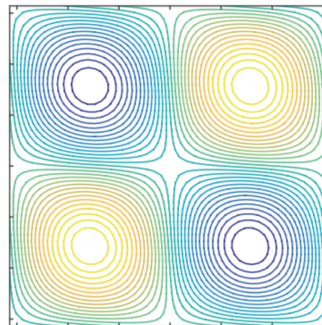
Bounds on Convection Enhanced Thermal Transport

simulations



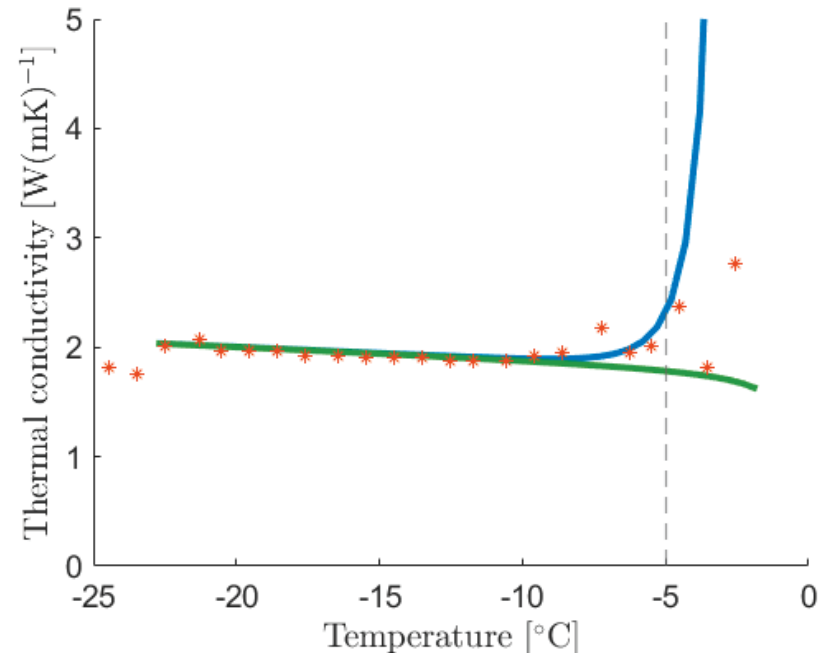
Monte-Carlo simulations of SDE with temperature dependent Péclet number P

strength of advection $B = \kappa P / 2\pi$
Euler-Maruyama and subsampling
methods for SDE



**cat's eye flow model for
brine convective flow**

data [Trodahl et al., 2001]



Rigorous Padé approximant bounds in terms of P using Stieltjes integral + analytic continuation method for the measure

Darcy velocity $v = 0.5$ $[\text{m/s}]$

wave propagation in the marginal ice zone (MIZ)

Stieltjes integral representation and bounds for the complex viscoelasticity of the ice - ocean layer

Sampson, Murphy, Cherkaev, Golden 2022

first theory of key parameter in wave-ice interactions only fitted to wave data before

Keller, 1998

Mosig, Montiel, Squire, 2015

Wang, Shen, 2012

Analytic Continuation Method

Bergman (78) - Milton (79)
integral representation for ϵ^*

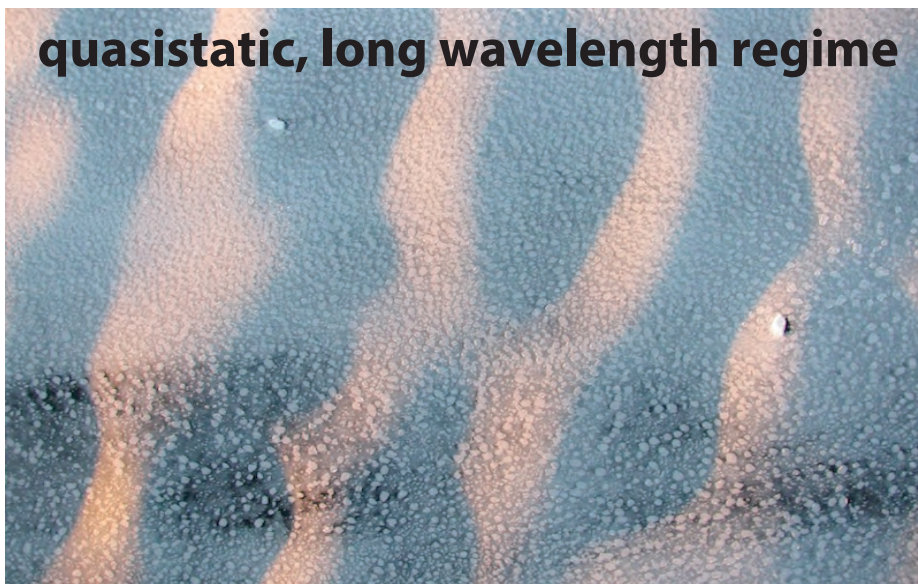
Golden and Papanicolaou (83)

Milton, *Theory of Composites* (02)

quasistatic, long wavelength regime

homogenized parameter depends on sea ice concentration and ice floe geometry

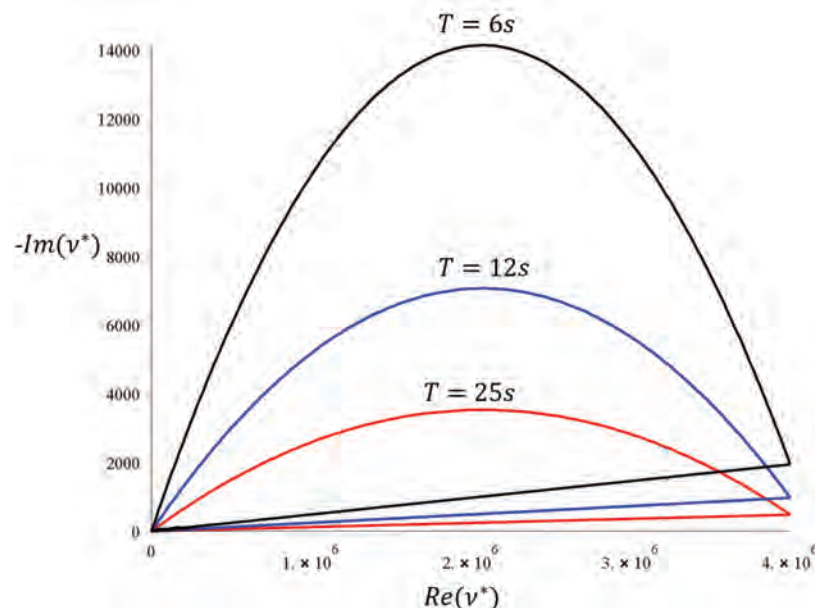
like EM waves



bounds on the effective complex viscoelasticity

$$\begin{aligned} V_1 &= 10^7 + i 4875 && \text{pancake ice} \\ V_2 &= 5 + i 0.0975 && \text{slush / frazil} \end{aligned}$$

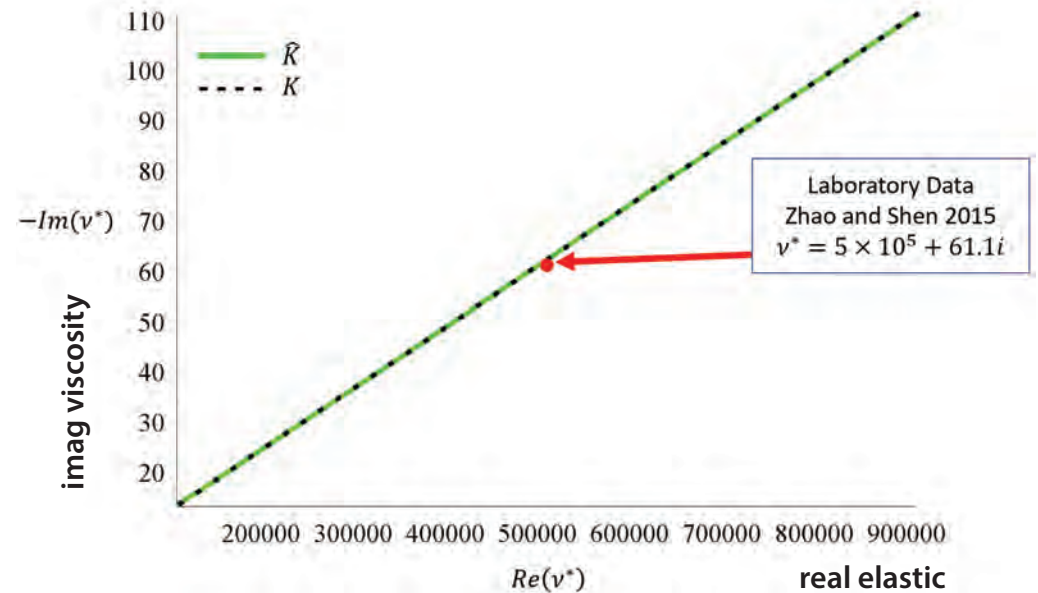
complex elementary bounds
(fixed area fraction of floes)



Elementary bounds for wave periods T .

high contrast

matrix-particle bounds



pancake ice

Golden

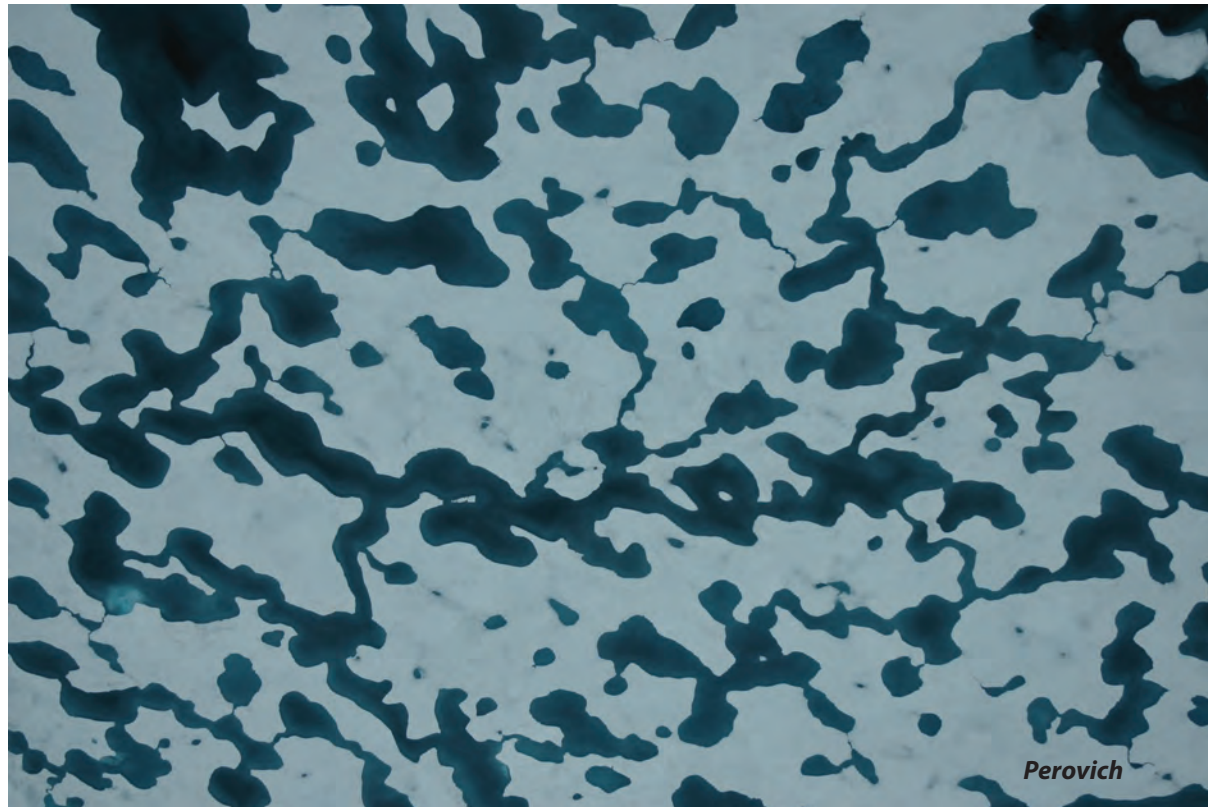
melt pond formation and albedo evolution:

- *major drivers in polar climate*
- *key challenge for global climate models*

numerical models of melt pond evolution, including topography, drainage (permeability), etc.

Lüthje, Feltham,
Taylor, Worster 2006
Flocco, Feltham 2007

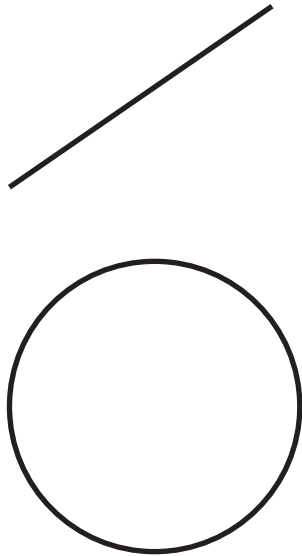
Skyllingstad, Paulson,
Perovich 2009
Flocco, Feltham,
Hunke 2012



Are there universal features of the evolution similar to phase transitions in statistical physics?

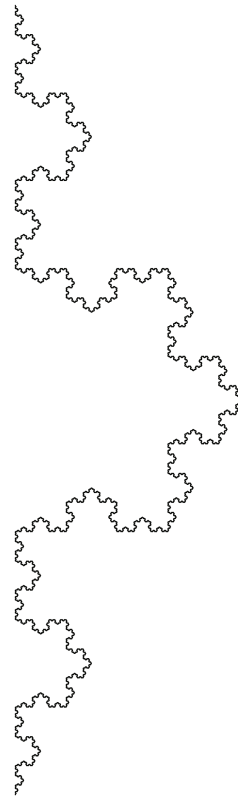
fractal curves in the plane

they wiggle so much that their dimension is >1



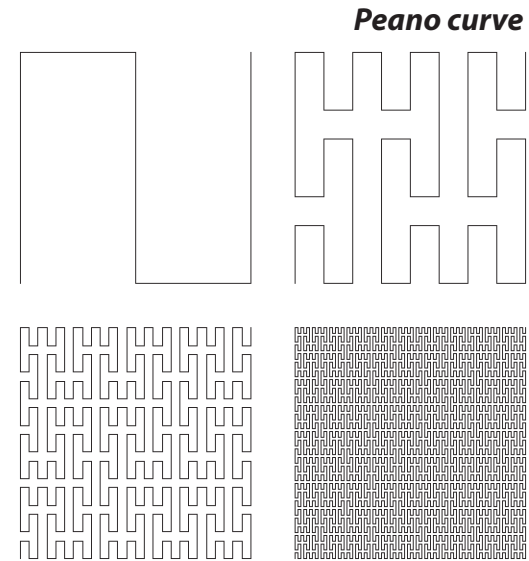
simple curves

$$D = 1$$



Koch snowflake

$$D = 1.26$$



Peano curve

Brownian motion

space filling curves

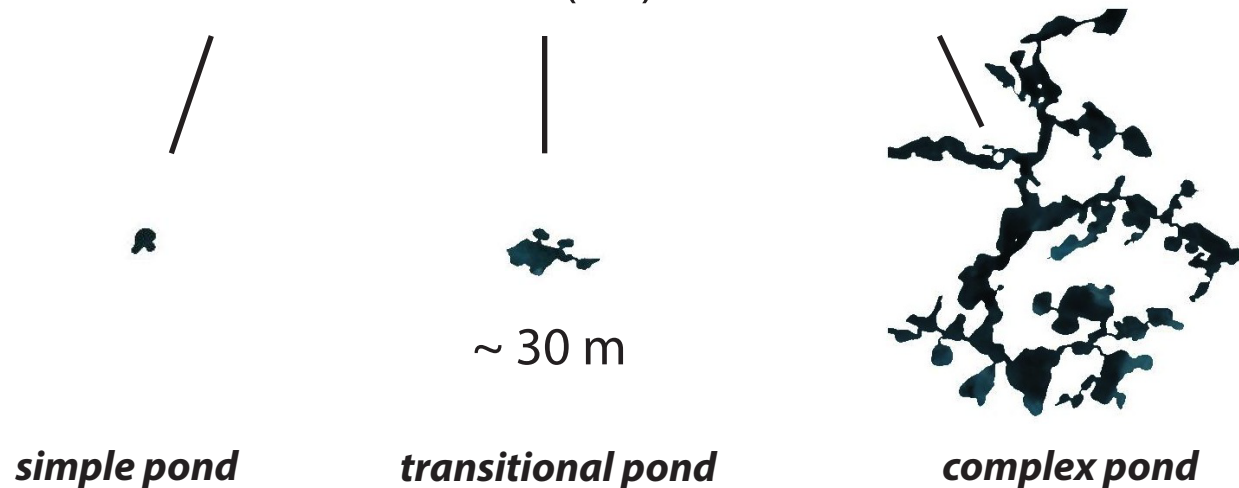
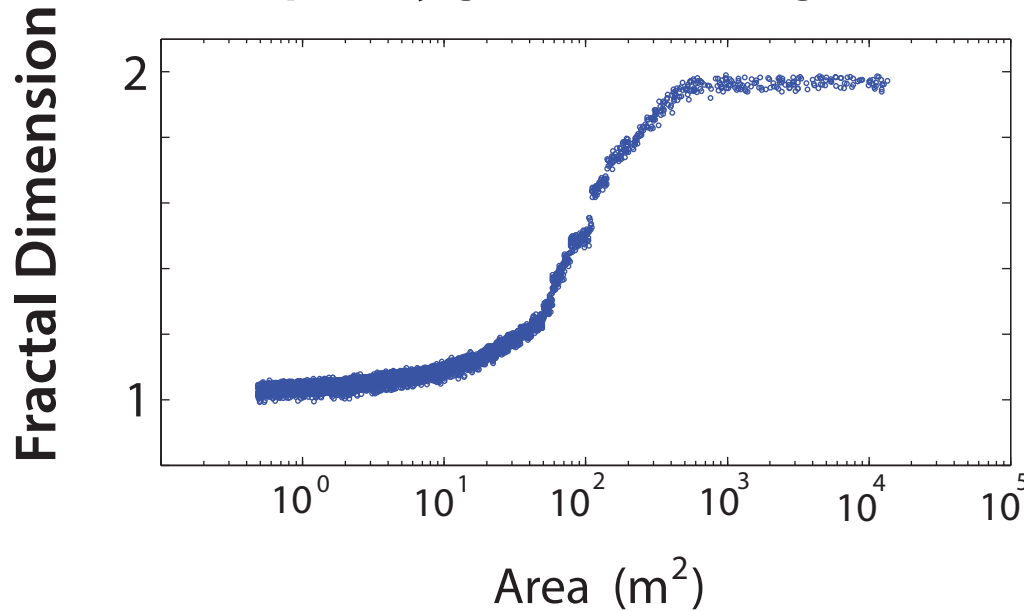
$$D = 2$$

Transition in the fractal geometry of Arctic melt ponds

Christel Hohenegger, Bacim Alali, Kyle Steffen, Don Perovich, Ken Golden

The Cryosphere, 2012

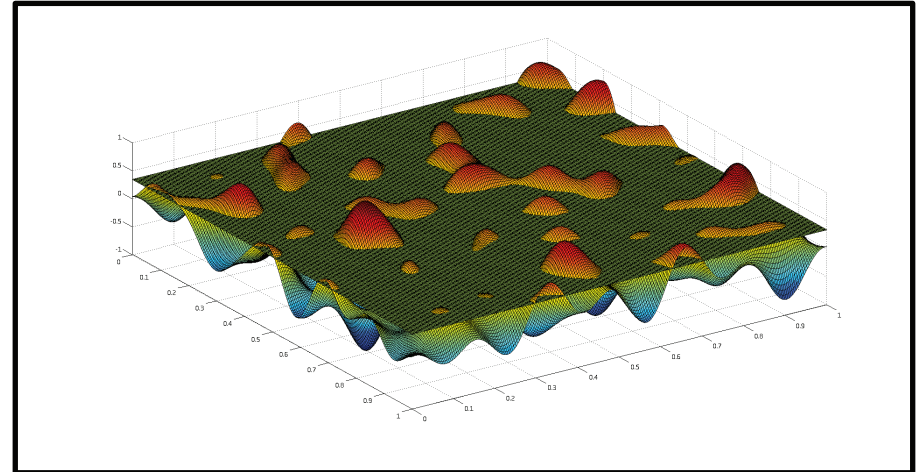
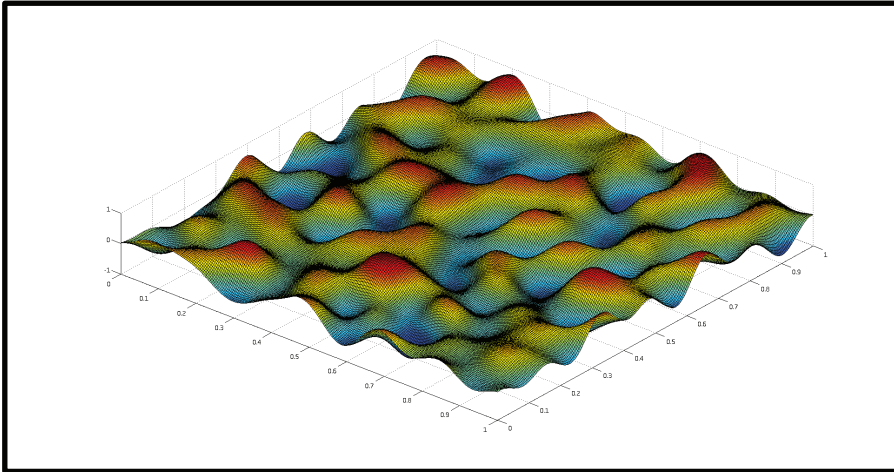
complexity grows with length scale



Continuum percolation model for melt pond evolution

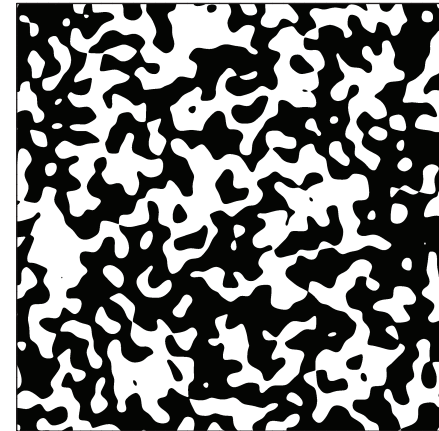
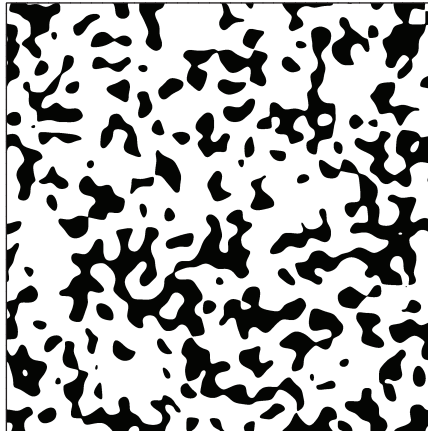
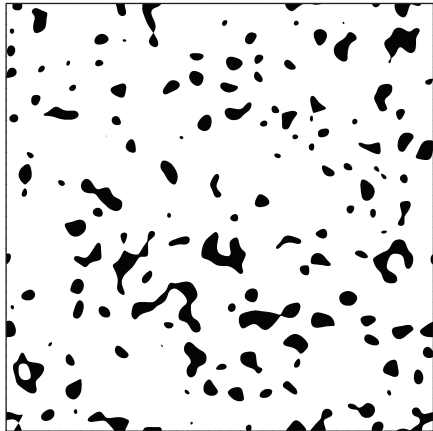
level sets of random surfaces

Brady Bowen, Court Strong, Ken Golden, J. Fractal Geometry 2018



random Fourier series representation of surface topography

intersections of a plane with the surface define melt ponds

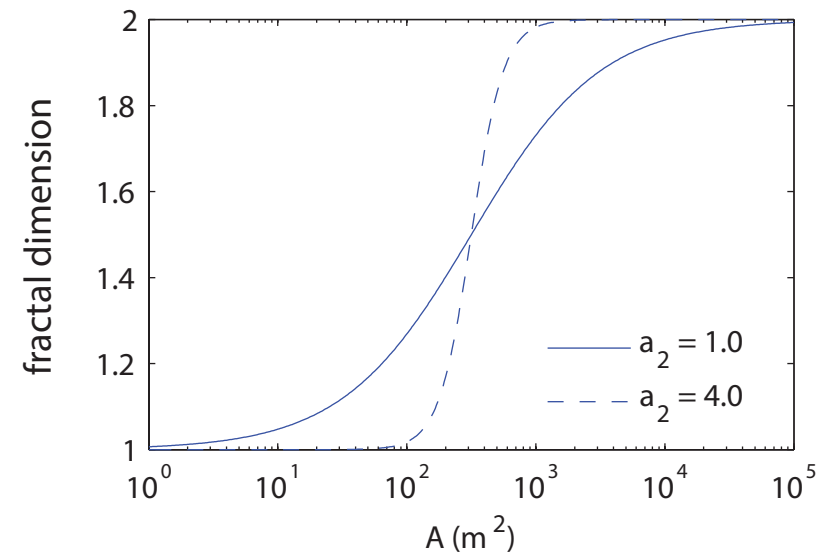
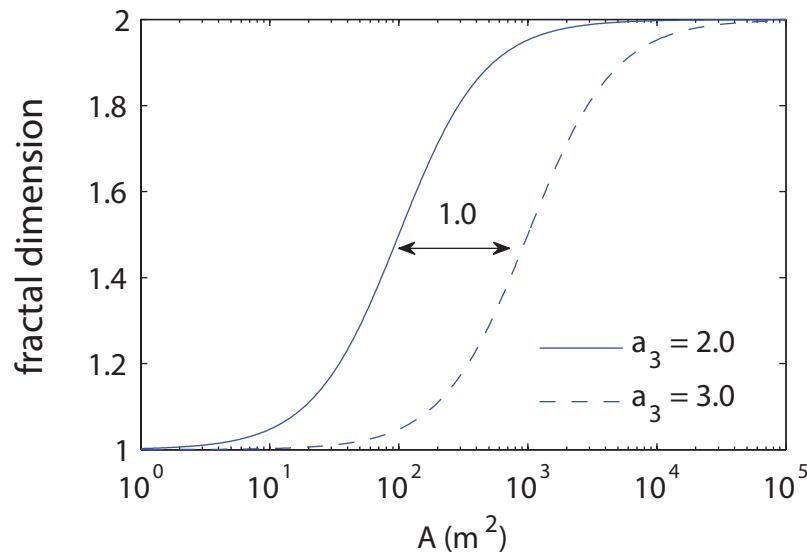


electronic transport in disordered media

diffusion in turbulent plasmas

Isichenko, Rev. Mod. Phys., 1992

fractal dimension curves depend on statistical parameters defining random surface



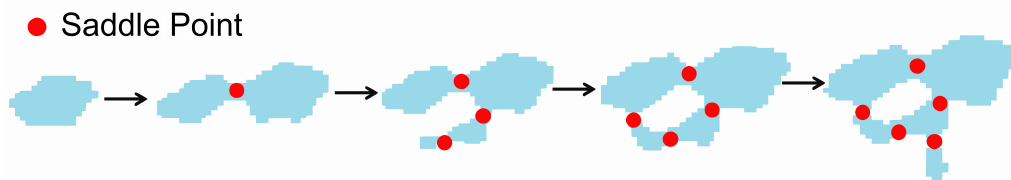
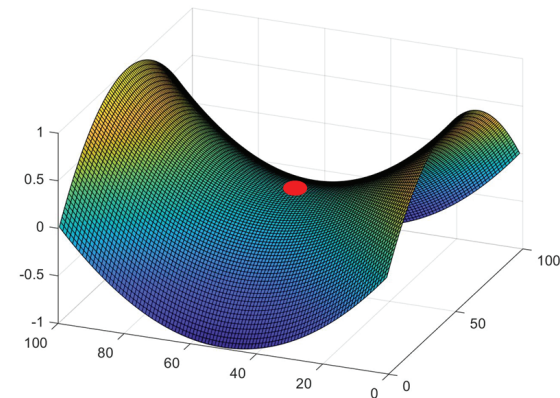
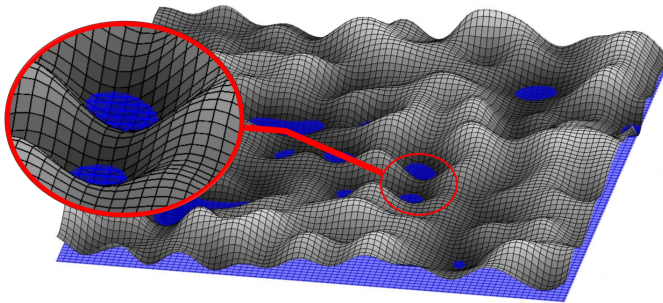
Topology of the sea ice surface and the fractal geometry of Arctic melt ponds

Physical Review Research (invited, under revision)

Ryleigh Moore, Jacob Jones, Dane Gollero,
Court Strong, Ken Golden

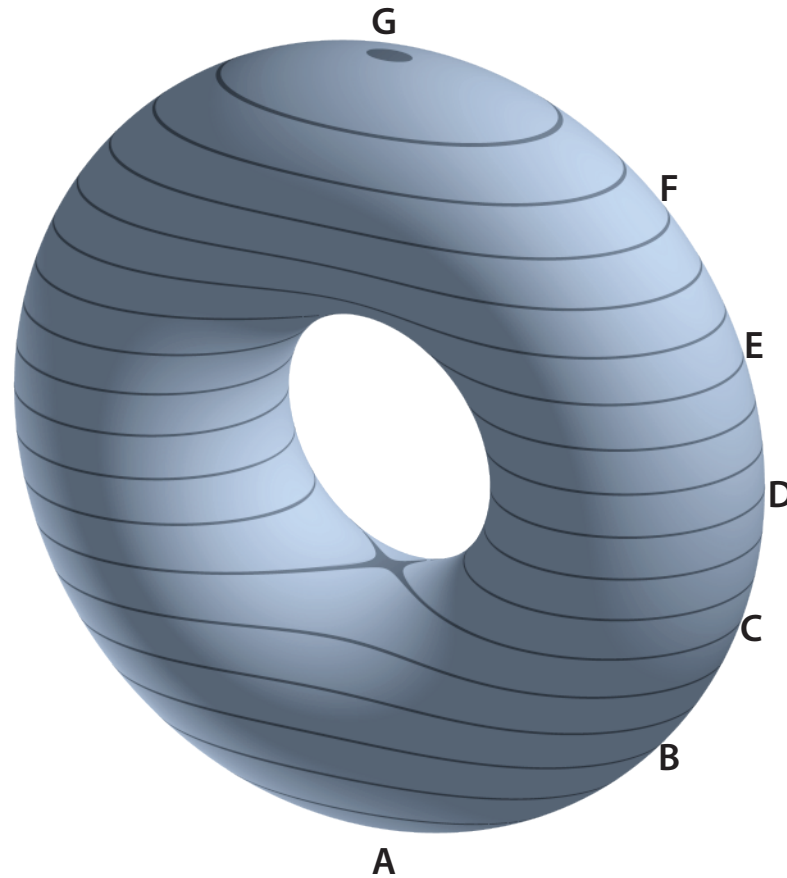
Several models replicate the transition in
fractal dimension, but none explain how it arises.

We use Morse theory applied to the random surface model
to show that **saddle points** play the critical role in the fractal transition.



ponds coalesce
(change topology) and
complexify at saddle points

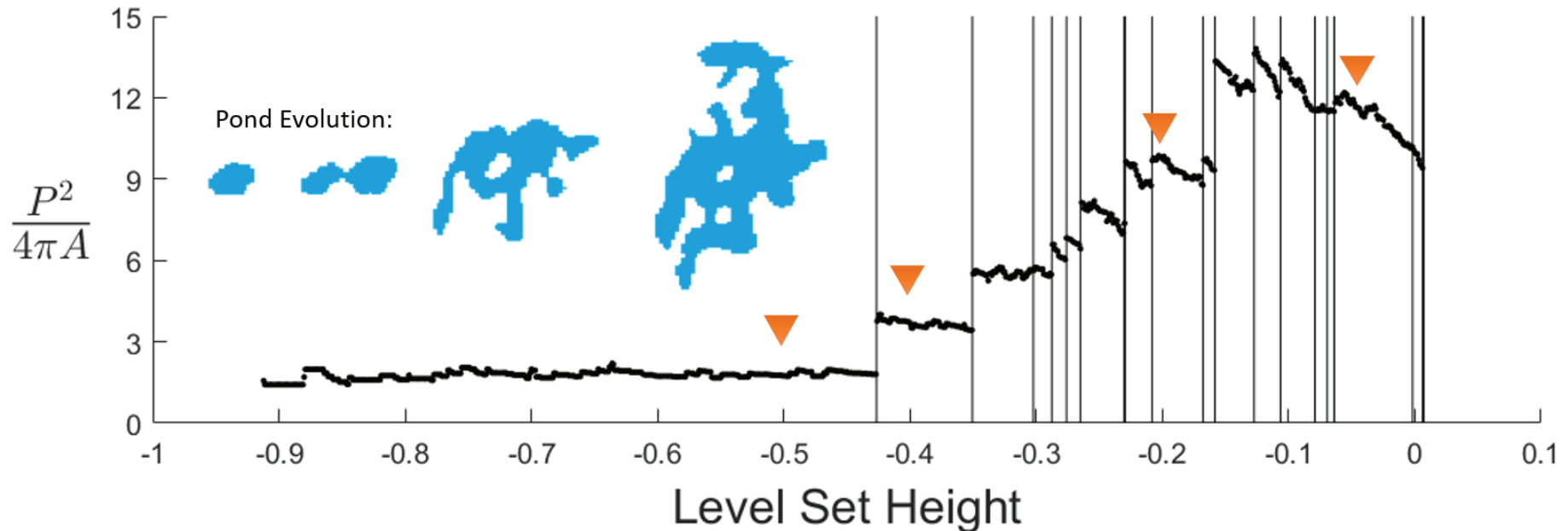
Morse theory



Morse theory tells us that changes in the topology of a surface occur at critical points of smooth functions on the surface: maxima, minima, and saddles.

Main results

Isoperimetric quotient - as a proxy for fractal dimension - increases in discrete jumps when ponds coalesce at saddle points.



Horizontal fluid permeability “controlled” by saddles ~ electronic transport in 2D random potential.

drainage processes, seal holes

melt pond evolution depends also on large-scale “pores” in ice cover



Melt pond connectivity enables vast expanses of melt water to drain down seal holes, thaw holes, and leads in the ice.

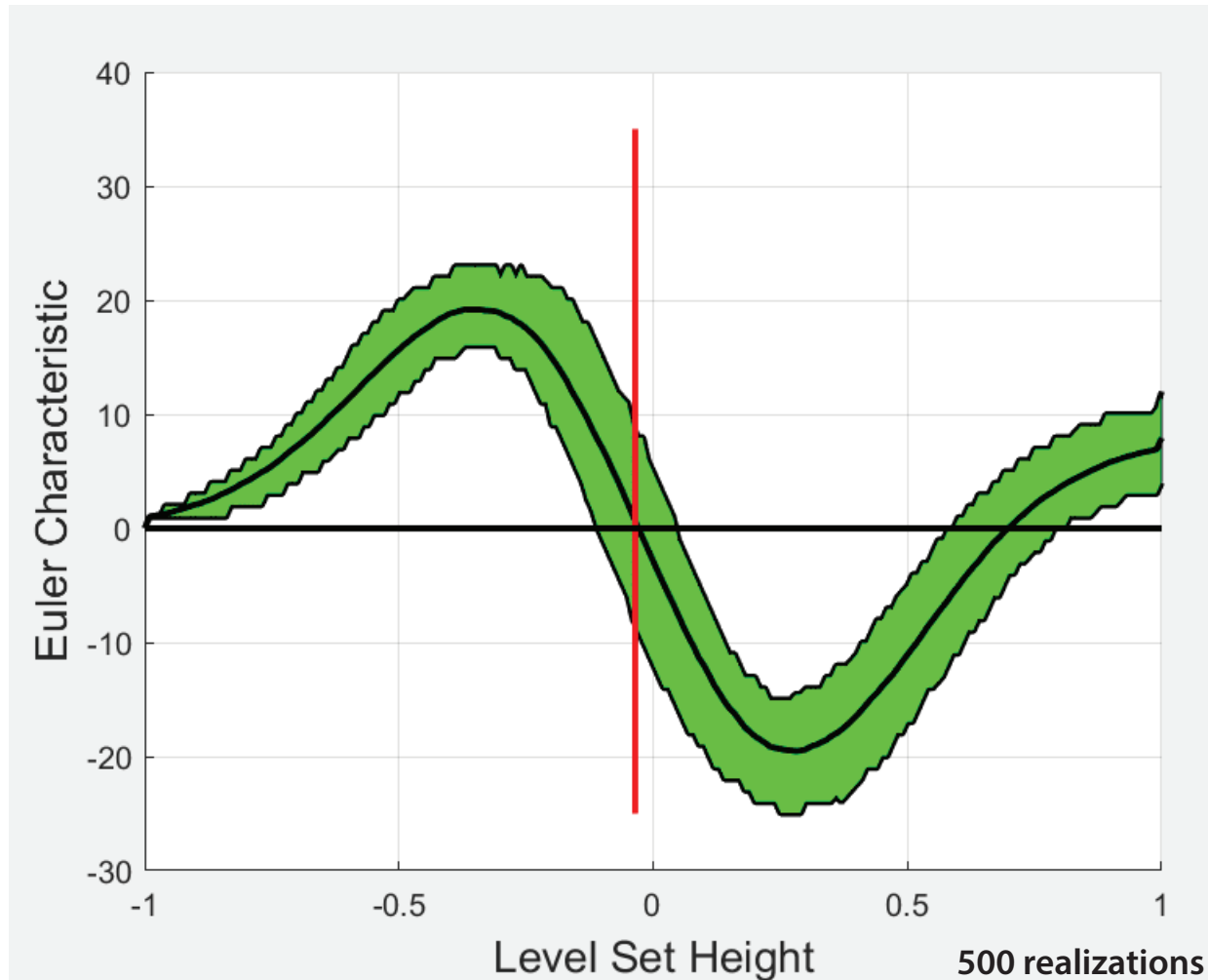
Topological Data Analysis

Euler characteristic = # maxima + # minima - # saddles

topological invariant

persistent homology

filtration - sequence of nested topological spaces, indexed by water level



Expected
Euler Characteristic Curve (ECC)

tracks the evolution of the EC of
the flooded surface as water rises

zero of ECC ~ percolation

percolation on a torus
creates a giant cycle

Bobrowski &
Skraba, 2020

Carlsson, 2009

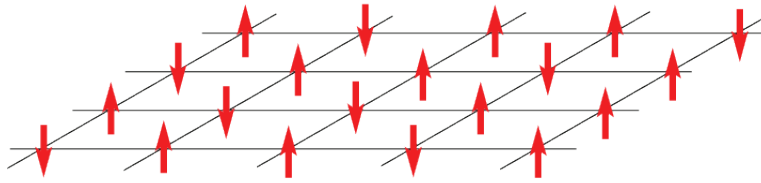
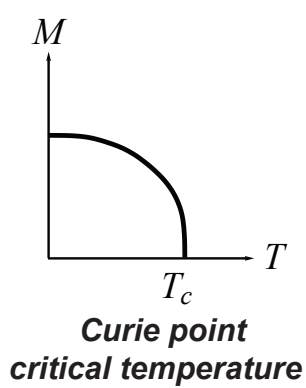
Vogel, 2002 GRF

porous media
cosmology
brain activity

melt pond donuts



Ising Model for a Ferromagnet



$$s_i = \begin{cases} +1 & \text{spin up} \\ -1 & \text{spin down} \end{cases} \quad \begin{matrix} \text{blue} \\ \text{white} \end{matrix}$$

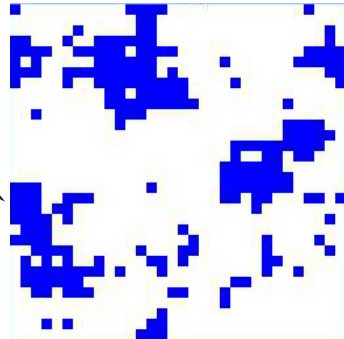
$$\mathcal{H} = -H \sum_i s_i - J \sum_{\langle i,j \rangle} s_i s_j$$

nearest neighbor Ising Hamiltonian

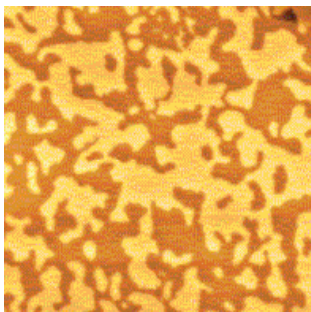
$$M(T, H) = \lim_{N \rightarrow \infty} \frac{1}{N} \left\langle \sum_j s_j \right\rangle$$

effective magnetization

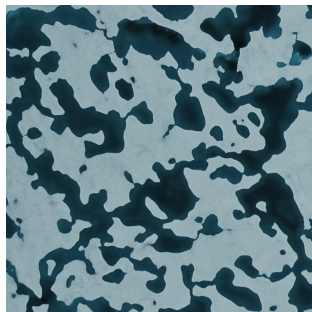
islands of like spins



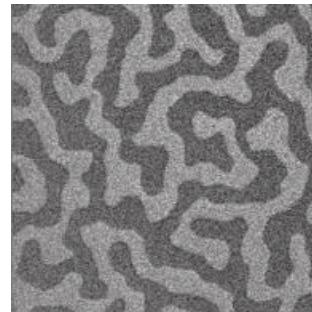
energy is lowered when nearby spins align with each other, forming **magnetic domains**



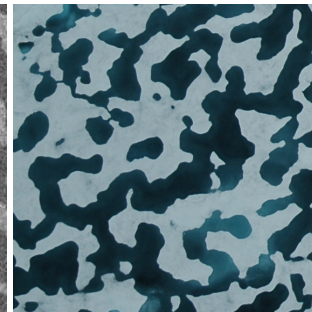
magnetic domains in cobalt



melt ponds (Perovich)



magnetic domains in cobalt-iron-boron



melt ponds (Perovich)

Ising model for ferromagnets \longrightarrow Ising model for melt ponds

Ma, Sudakov, Strong, Golden, *New J. Phys.*, 2019

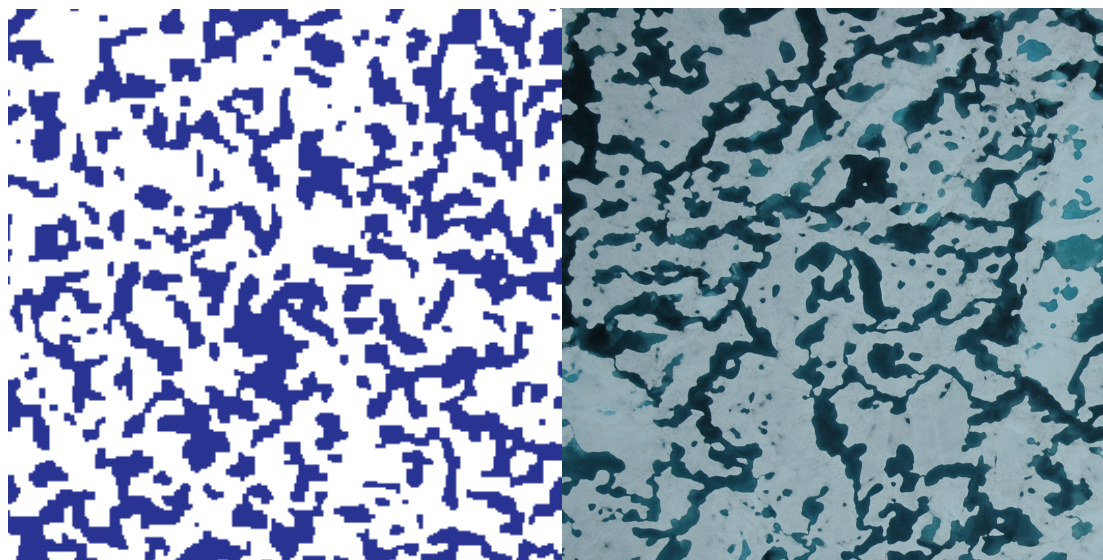
$$\mathcal{H} = - \sum_i^N H_i s_i - J \sum_{\langle i,j \rangle}^N s_i s_j \quad s_i = \begin{cases} \uparrow & +1 \text{ water (spin up)} \\ \downarrow & -1 \text{ ice (spin down)} \end{cases}$$

random magnetic field
represents snow topography

magnetization M pond area fraction $F = \frac{(M+1)}{2}$ only nearest neighbor patches interact
 \sim albedo

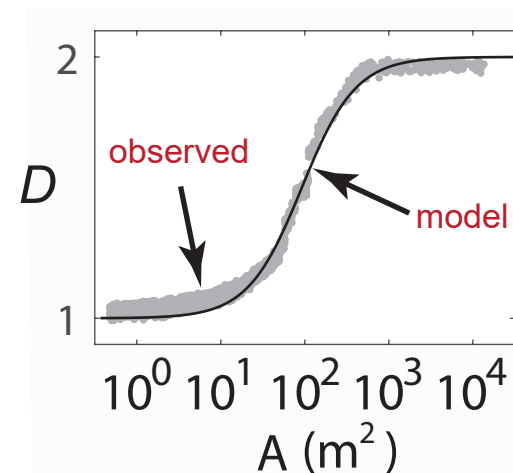
Starting with random initial configurations, as Hamiltonian energy is minimized by Glauber spin flip dynamics, system “flows” toward metastable equilibria.

Order from Disorder



Ising
model

melt pond
photo (Perovich)



pond size
distribution exponent

observed -1.5

(Perovich, et al. 2002)

model -1.58

*Scientific American
EOS, PhysicsWorld, ...*

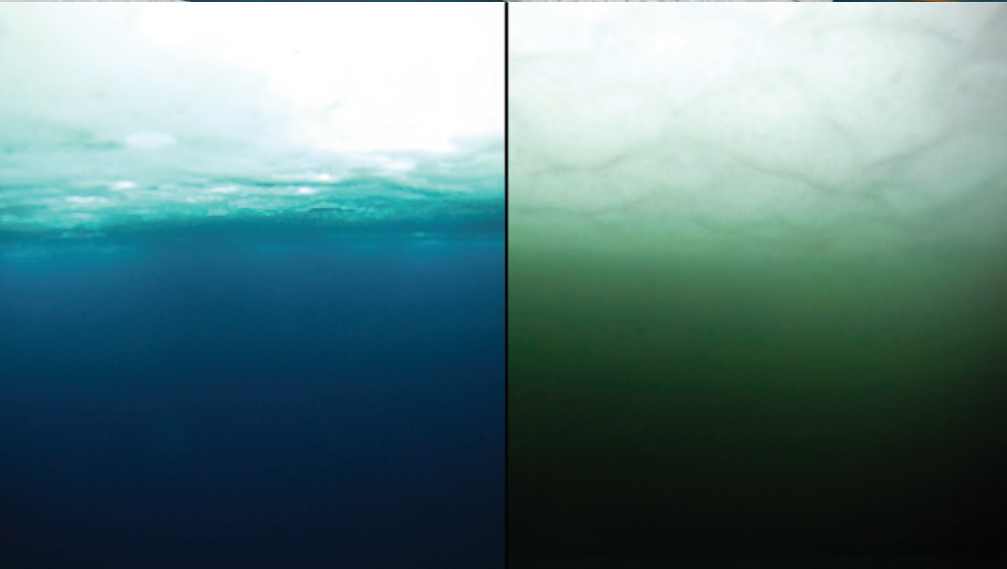
ONLY MEASURED INPUT = LENGTH SCALE (GRID SIZE) from snow topography data



Perovich

Melt ponds control transmittance of solar energy through sea ice, impacting upper ocean ecology.

WINDOWS



no bloom

bloom

massive under-ice **algal bloom**

Arrigo et al., *Science* 2012

Have we crossed into a new ecological regime?

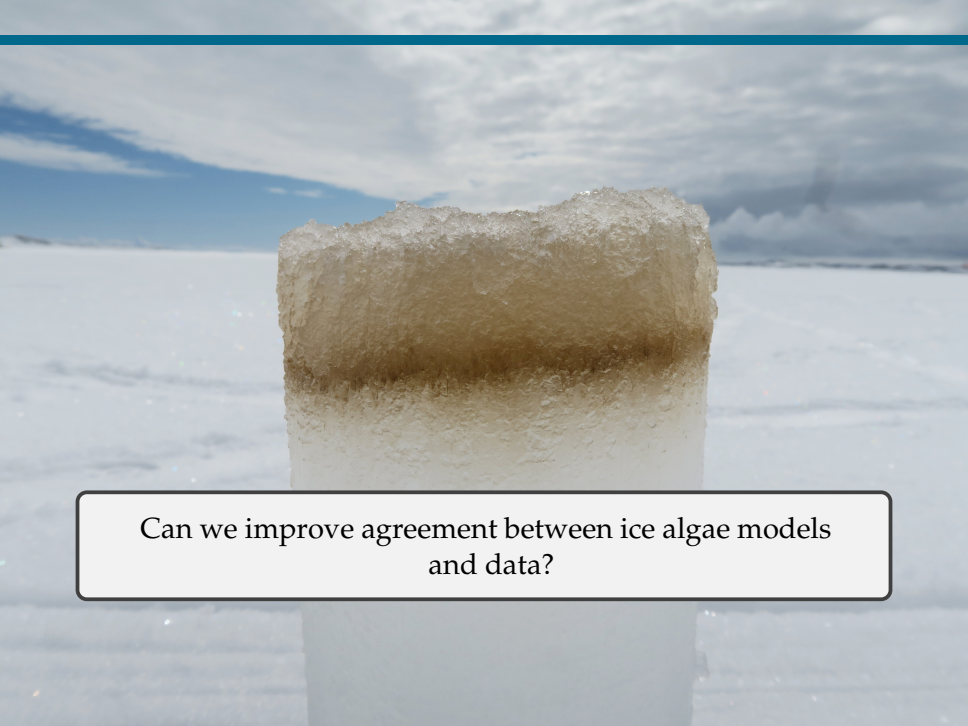
The frequency and extent of sub-ice phytoplankton blooms in the Arctic Ocean

Horvat, Rees Jones, Iams, Schroeder, Flocco, Feltham, *Science Advances* 2017

The effect of melt pond geometry on the distribution of solar energy under first year sea ice

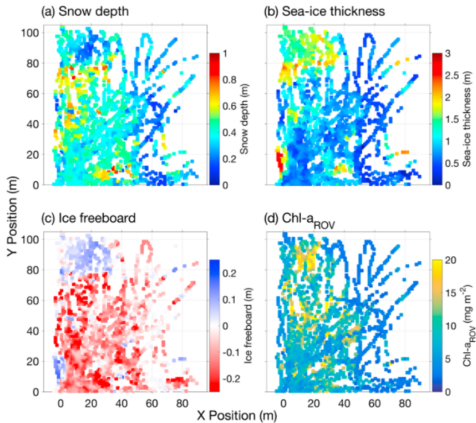
Horvat, Flocco, Rees Jones, Roach, Golden
Geophys. Res. Lett. 2019

(2015 AMS MRC)



Can we improve agreement between ice algae models
and data?

HETEROGENEITY



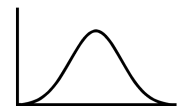
HETEROGENEITY IN INITIAL CONDITIONS

At each location within a larger region, we could consider

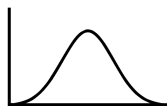
$$\text{Nutrients} \quad \frac{dN}{dt} = \alpha - \textcolor{brown}{B}NP - \eta N$$

$$\text{Algae} \quad \frac{dP}{dt} = \gamma \textcolor{brown}{B}NP - \delta P$$

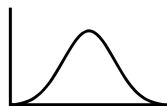
$$N(0) = \textcolor{brown}{N}_0, \quad P(0) = \textcolor{brown}{P}_0$$



growth rate, $\textcolor{brown}{B}$



Initial nutrients, $\textcolor{brown}{N}_0$



Initial algae, $\textcolor{brown}{P}_0$

METHOD

Uncertainty quantification for ecological models with random parameters

Jody R. Reimer^{1,2}  | Frederick R. Adler^{1,2}  | Kenneth M. Golden¹  | Akil Narayan^{1,3} 

¹Department of Mathematics, University of Utah, Salt Lake City, Utah, USA

²School of Biological Sciences, University of Utah, Salt Lake City, Utah, USA

³Scientific Computing and Imaging Institute, University of Utah, Salt Lake City, Utah, USA

Correspondences

Jody R. Reimer, Department of Mathematics and School of Biological Sciences, University of Utah, Salt Lake City, Utah, USA.

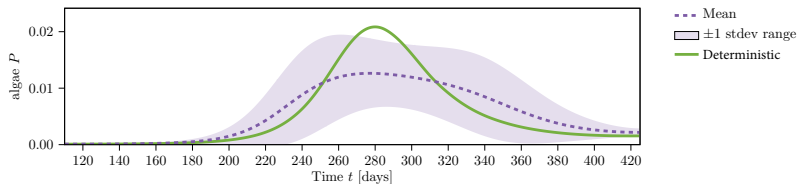
Email: reimer@math.utah.edu

Abstract

There is often considerable uncertainty in parameters in ecological models. This uncertainty can be incorporated into models by treating parameters as random variables with distributions, rather than fixed quantities. Recent advances in uncertainty quantification methods, such as polynomial chaos approaches, allow for the analysis of models with random parameters. We introduce these methods with a motivating case study of sea ice algal blooms in heterogeneous environments. We compare Monte Carlo methods with polynomial chaos techniques to help understand the dynamics of an algal bloom model with random parameters.

Introduce polynomial chaos approach to widely used ecological ODE models, but with random parameters.

ECOLOGICAL INSIGHTS



- lower peak bloom intensity
- longer bloom duration
- able to compare variance to data

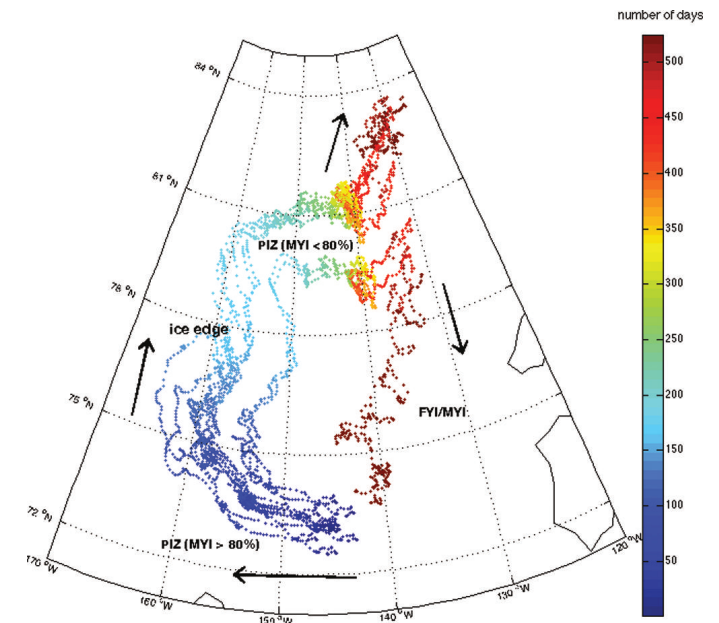
macroscale

Anomalous diffusion in sea ice dynamics

Ice floe diffusion in winds and currents

observations from GPS data:

Jennifer Lukovich, Jennifer Hutchings,
David Barber, *Ann. Glac.* 2015



- On short time scales floes observed (buoy data) to exhibit Brownian-like behavior, but they are also being advected by winds and currents.
- Effective behavior is purely diffusive, sub-diffusive or super-diffusive depending on ice pack and advective conditions - **Hurst exponent**.

modeling:

Huy Dinh, Ben Murphy, Elena Cherkaev,
Court Strong, Ken Golden 2022

floe scale model to analyze transport regimes in
terms of ice pack crowding, advective conditions

Delaney Mosier, Jennifer Hutchings, Jennifer Lukovich,
Marta D'Elia, George Karniadakis, Ken Golden 2022

learning fractional PDE
governing diffusion from data

Floe Scale Model of Anomalous Diffusion in Sea Ice Dynamics

Huy Dinh, Ben Murphy, Elena Cherkaev, Court Strong, Ken Golden 2022

$$\langle |\mathbf{x}(t) - \mathbf{x}(0) - \langle \mathbf{x}(t) - \mathbf{x}(0) \rangle|^2 \rangle \sim t^\alpha$$

α = Hurst exponent

diffusive $\alpha = 1$
sub-diffusive $\alpha < 1$
super-diffusive $\alpha > 1$

Model Approximations

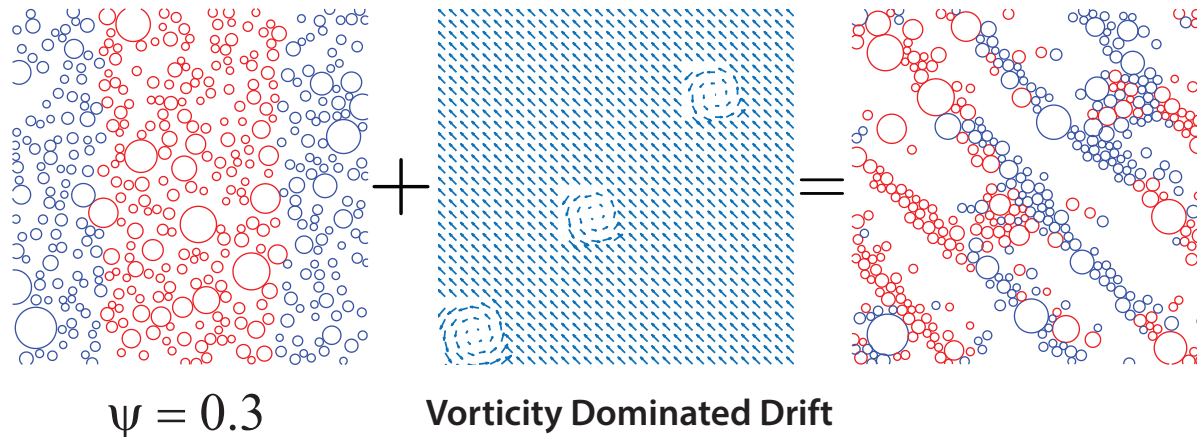
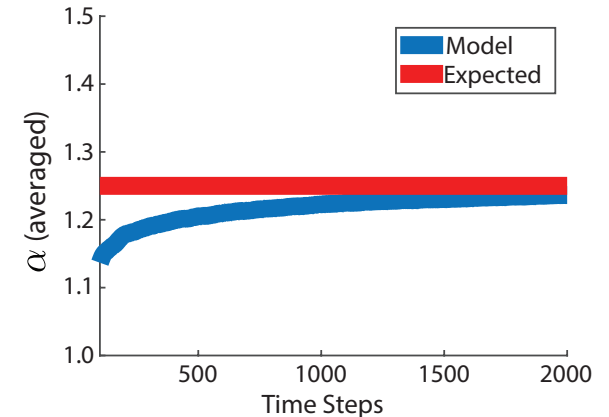
Power Law Size Distribution: $N(D) \sim D^{-k}$

D. A. Rothrock and A. S. Thorndike Journal of Geophysical Research 1984

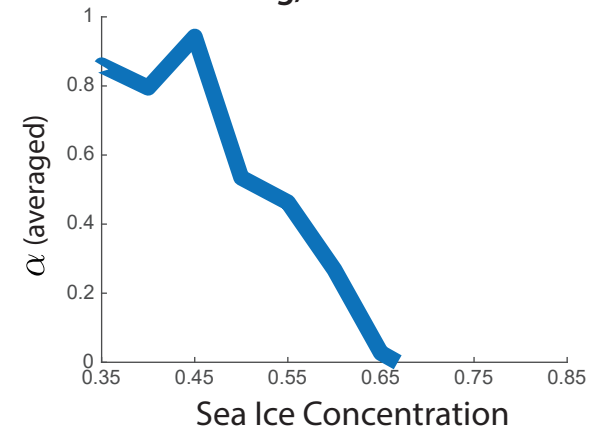
Floe-Floe Interactions: Linear Elastic Collisions

Advective Forcing: Passive, Linear Drag Law

Sparse Packing, Shear Dominated Drift



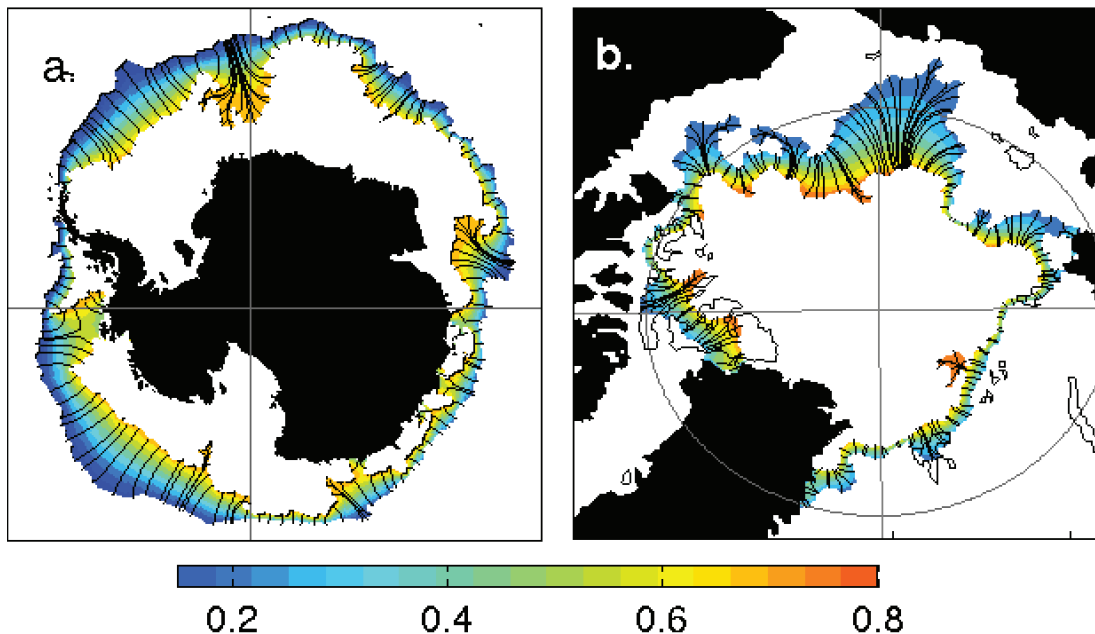
Crowding, Diffusive Drift



Marginal Ice Zone

MIZ

- biologically active region
- intense ocean-sea ice-atmosphere interactions
- region of significant wave-ice interactions



MIZ WIDTH

fundamental length scale of
ecological and climate dynamics

Strong, *Climate Dynamics* 2012

Strong and Rigor, *GRL* 2013

transitional region between
dense interior pack ($c > 80\%$)
sparse outer fringes ($c < 15\%$)

**How to objectively
measure the “width”
of this complex,
non-convex region?**

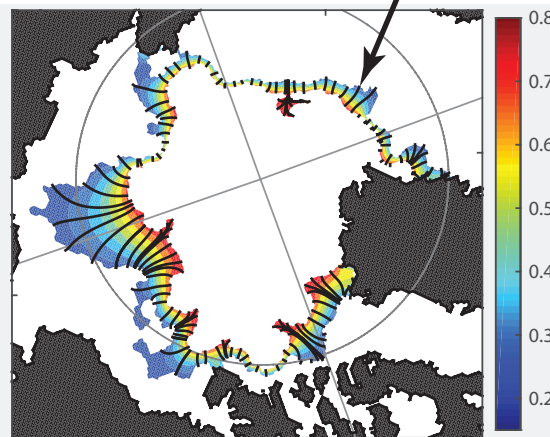
Objective method for measuring MIZ width motivated by medical imaging and diagnostics

Strong, *Climate Dynamics* 2012
Strong and Rigor, *GRL* 2013

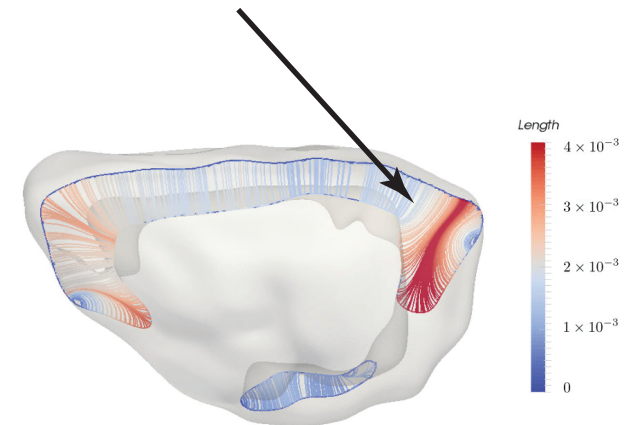
39% widening
1979 - 2012

“average” lengths of streamlines

streamlines of a solution
to Laplace’s equation



Arctic Marginal Ice Zone



**cross-section of the
cerebral cortex of a rodent brain**

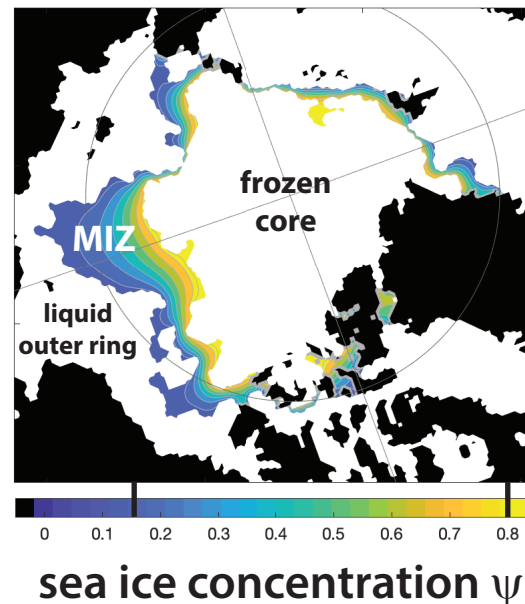
analysis of different MIZ WIDTH definitions

Strong, Foster, Cherkaev, Eisenman, Golden
J. Atmos. Oceanic Tech. 2017

Strong and Golden
Society for Industrial and Applied Mathematics News, April 2017

Model larger scale effective behavior
with partial differential equations that
homogenize complex local structure and dynamics.

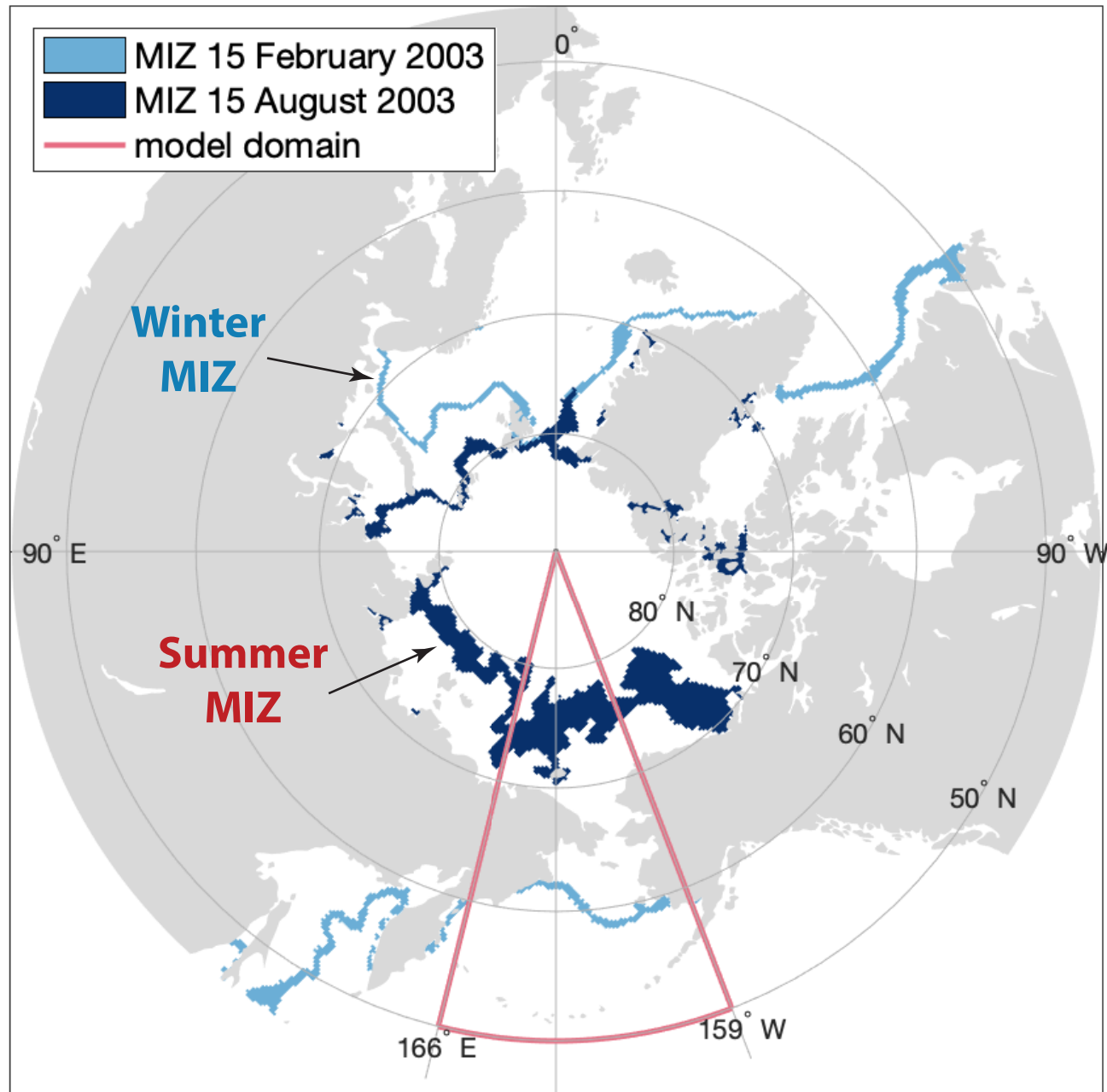
Arctic MIZ



Predict MIZ width and location with basin-scale phase change model.
dynamic transitional region - mushy layer - separating two “pure” phases
seasonal and long term trends

C. Strong, E. Cherkaev, and K. M. Golden,
Annual cycle of Arctic marginal ice zone location
and width explained by dynamic phase transition model, 2022

Observed Arctic MIZ



MIZ as a moving phase transition region

$$\rho c \frac{\partial T}{\partial t} = \nabla \cdot (k \nabla T) + S$$

$$S = [\rho(c_l - c_s)T + \rho L] \frac{\partial \psi}{\partial t}$$

$$\psi = 1 - \left(\frac{T - T_s}{T_l - T_s} \right)^\alpha$$

$$k_x = \left(\frac{\psi}{k_s} + \frac{1 - \psi}{k_l} \right)^{-1}$$

$$k_z = \psi k_s + (1 - \psi) k_l$$

homogenization

ρ effective density

T temperature

c specific heat

L latent heat of fusion

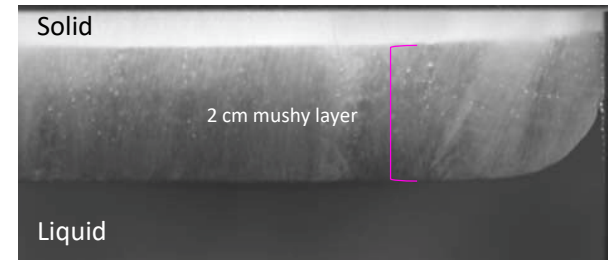
S models nonlinear phase change

ψ sea ice concentration

k effective diffusivity

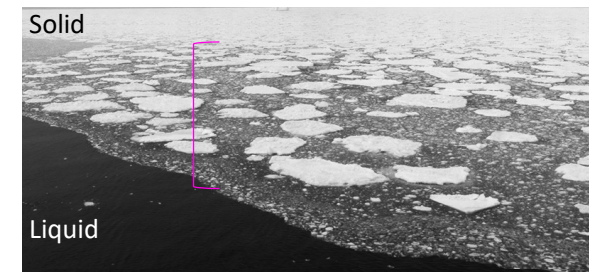
l liquid, s solid

Classical small-scale application



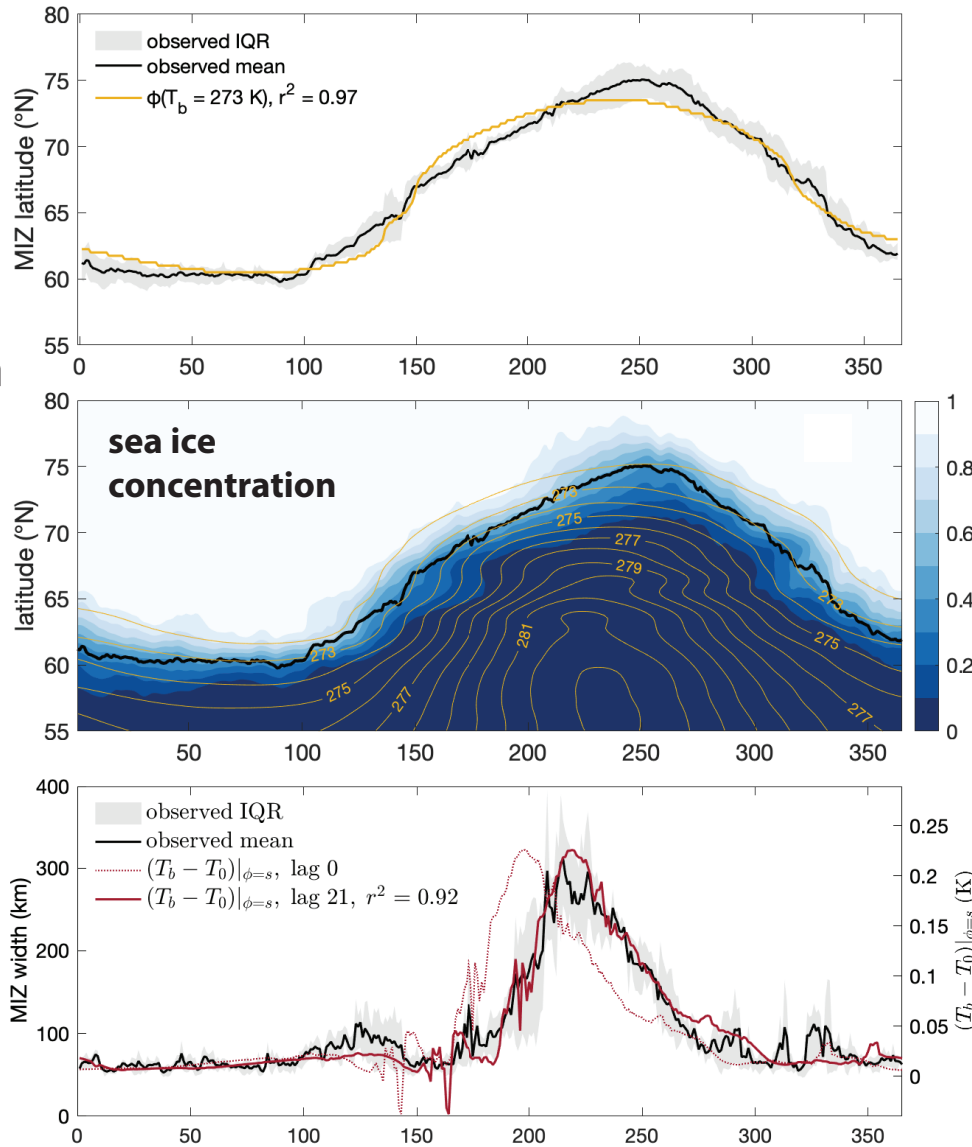
NaCl-H₂O in lab
(Peppin et al., 2007; J. Fluid Mech.)

Macroscale application

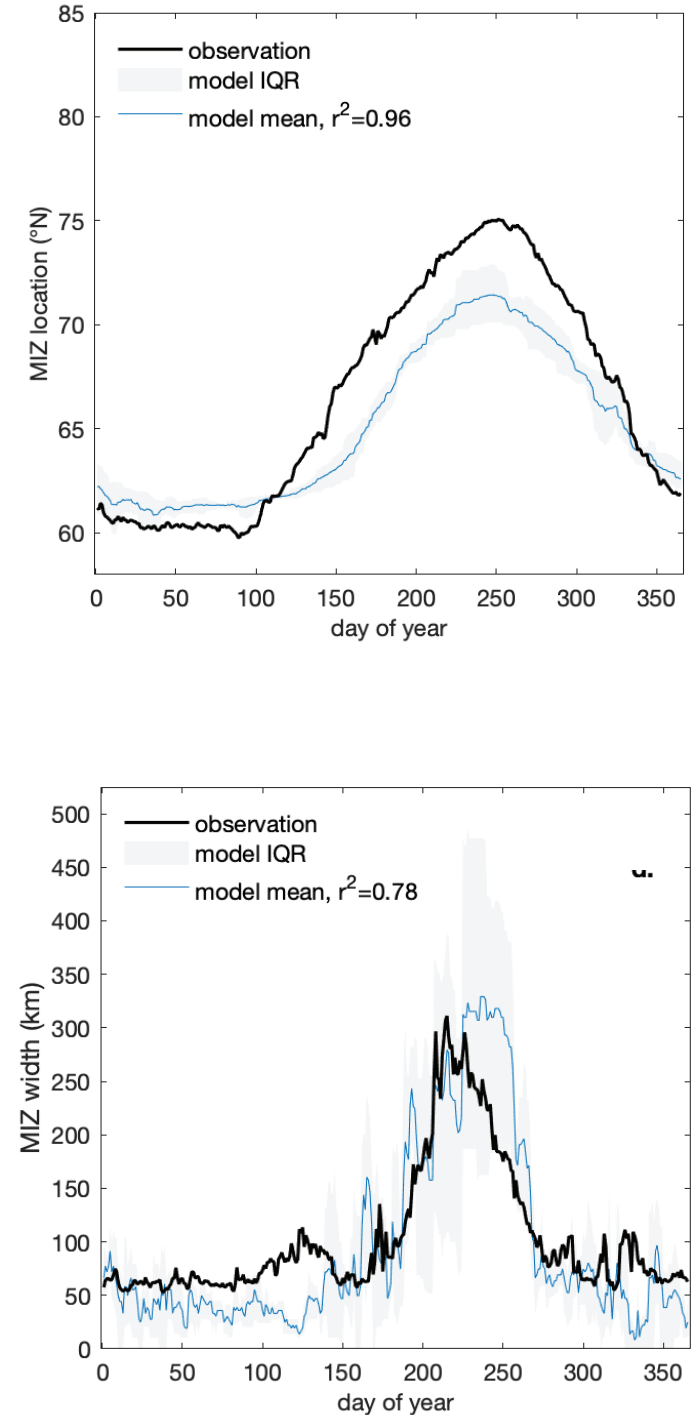


- Develop multiscale PDE model for simulating phase transition fronts to predict MIZ seasonal cycles and decadal trends
- Model simulates MIZ as a large-scale mushy layer with effective thermal conductivity derived from physics of composite materials

MIZ observations



MIZ model vs. observations



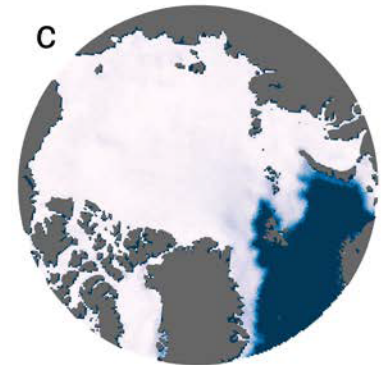
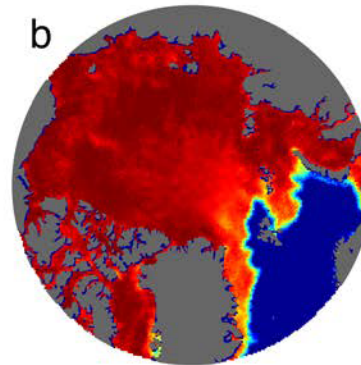
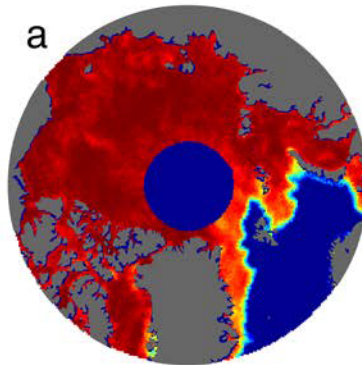
Model captures basic physics of MIZ dynamics.

Filling the polar data gap with partial differential equations

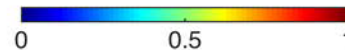
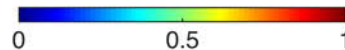
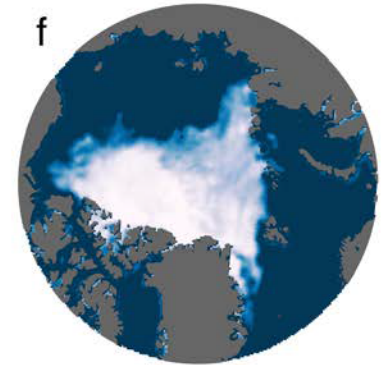
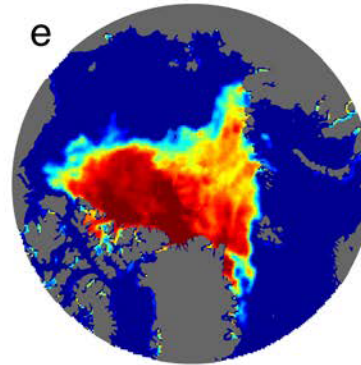
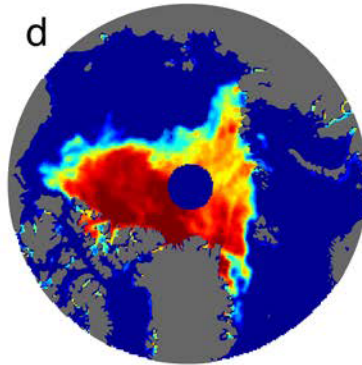
hole in satellite coverage
of sea ice concentration field

previously assumed
ice covered

Gap radius: 611 km
06 January 1985



Gap radius: 311 km
30 August 2007



$$\Delta\psi=0$$

fill = harmonic function with
learned stochastic term

Strong and Golden, *Remote Sensing* 2016
Strong and Golden, *SIAM News* 2017

NOAA/NSIDC Sea Ice Concentration CDR
product update will use our PDE method.

Conclusions

1. Sea ice is a fascinating multiscale composite/fluid structure with similarities to many other natural and man-made materials.
2. Mathematics developed for sea ice advances the theory of composites and other areas of science and engineering.
3. **Homogenization and statistical physics help *link scales in sea ice and composites***; provide rigorous methods for finding effective behavior; advance sea ice representations in climate models.
4. **Fluid flow** through sea ice mediates **melt pond evolution** and many processes important to climate change and polar ecosystems.
5. Field experiments are essential to developing relevant mathematics.
6. Our research is helping to **improve projections of climate change**, the fate of Earth's sea ice packs, and the ecosystems they support.



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Notices

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OPEN POSITIONS

MATHEMATICS of SEA ICE and CLIMATE

Postdoctoral Fellowships (3 year)

Graduate Fellowships (Ph.D.)

Undergraduate Research

NSF Research Training Grant (5 years)

ONR + NSF Research Grants

research and teaching

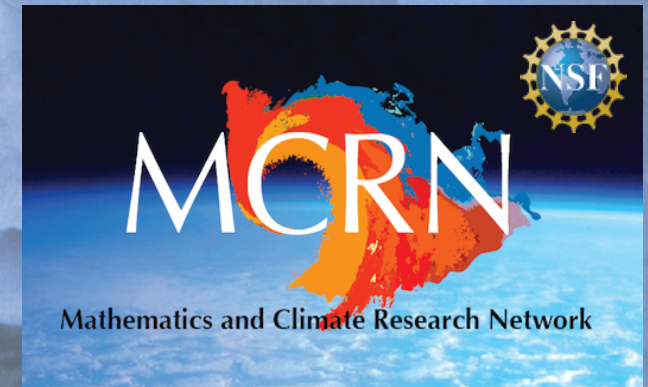
THANK YOU

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National Science Foundation

Division of Mathematical Sciences
Division of Polar Programs



Buchanan Bay, Antarctica Mertz Glacier Polynya Experiment July 1999

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Christian Sampson (now at UNC Chapel Hill with Chris Jones)
Huy Dinh (now a sea ice MURI Postdoc at NYU/Courant)
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David Morison (Physics Department)
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Kitsel Lusted, Ruby Bowers, Kimball Johnston,
Jerry Zhang, Nash Ward, David Gluckman

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Sea Ice Ecology Group Postdoc Jody Reimer, Grad Student Julie Sherman,
Undergraduates Kayla Stewart, Nicole Forrester

SEA ICE covers ~12% of Earth's ocean surface

- boundary between ocean and atmosphere
- mediates exchange of heat, gases, momentum
- global ocean circulation
- hosts rich ecosystem
- indicator of **climate change**



polar ice caps critical to global climate in reflecting incoming solar radiation



white snow and ice
reflect



dark water and land
absorb

$$\text{albedo } \alpha = \frac{\text{reflected sunlight}}{\text{incident sunlight}}$$